# Fluid Substitution with Dynamic Fluid Modulus: Facing the Challenges in Heterogeneous Rocks

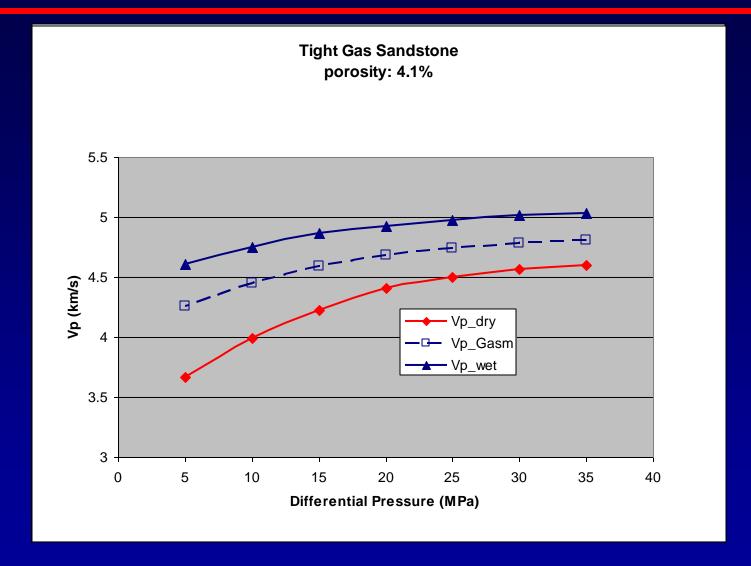
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#### **Observations: where does Gassmann break?**



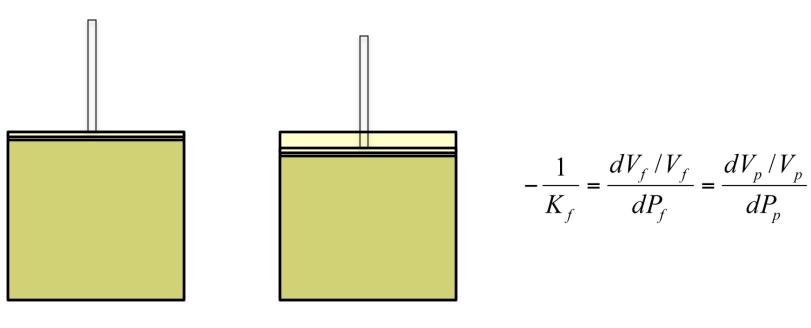


#### **Outline**

- 1. Fluid effect in closed system: Gassmann
- 2. Fluid effect in non- closed system: partial drainage and DFM
- 3. Wave induced internal flow and dispersion
- 4. Modeling Examples:
  - Mesoscopic heterogeneity
  - Microscopic heterogeneity (crack-pore system)
- 5. DFM from real measured data

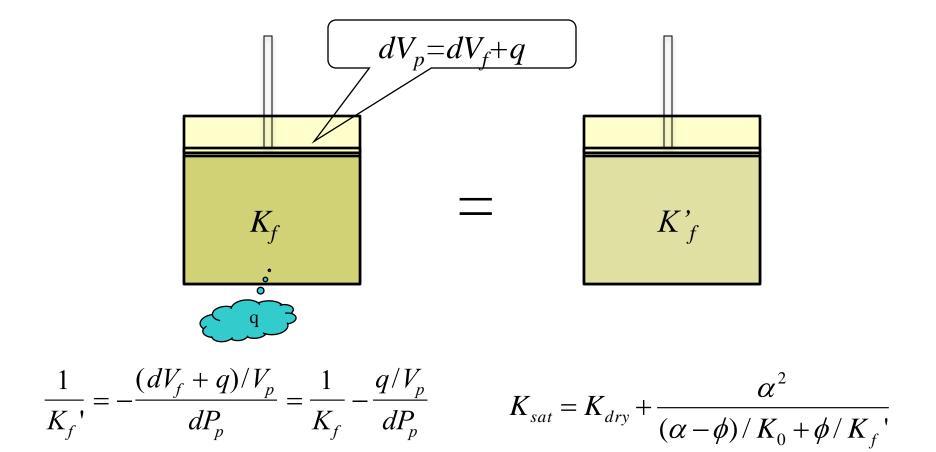


### Fluid effect in closed system: Gassmann



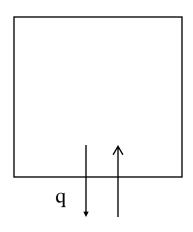
$$K_{sat} = K_{dry} + \frac{\alpha^2}{(\alpha - \phi)/K_0 + \phi/K_f}$$

#### Fluid effect in non-closed system: partial drainage and DFM

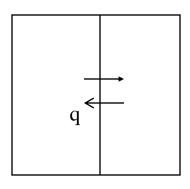


q: "+" for incoming flow, "-" for outgoing flow

# What about closed system with internal flow?

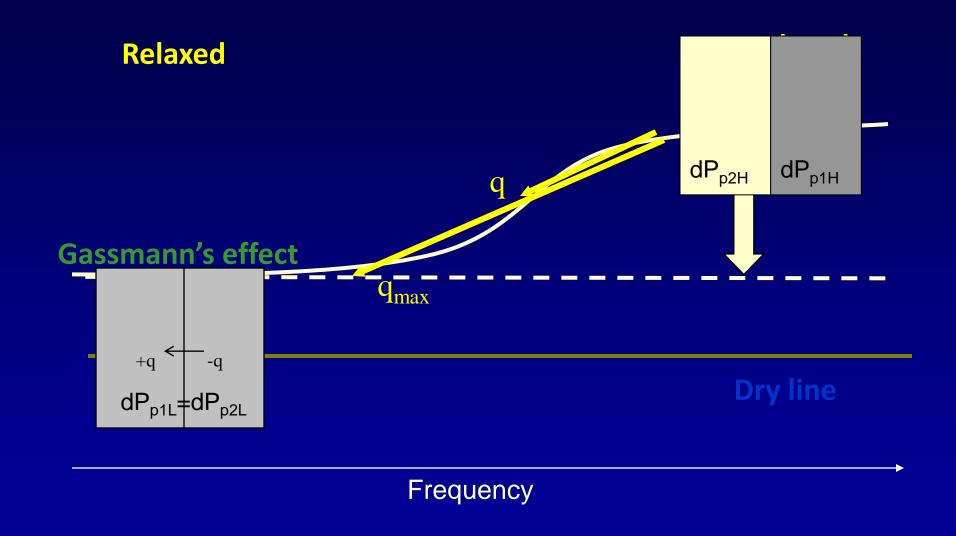






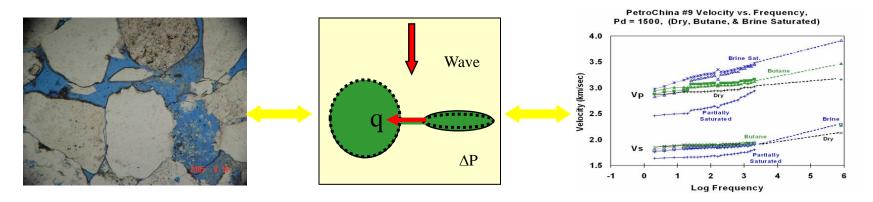
Internal flow

# What's the role of fluid flow?





#### What's the role of fluid flow?



How much fluid moved?

$$q_{\text{max}} = \frac{V_{p2}}{K_f} (dP_{p2H} - dP_{p1H})$$

How much more deformed due to relaxation?

$$\delta V_L - \delta V_H = -\beta_f V_{p2} (\delta P_{p1H} - \delta P_{p2H})$$

$$\delta V_L - \delta V_H = q_{\text{max}}$$

#### How to calculate q=q(f)?

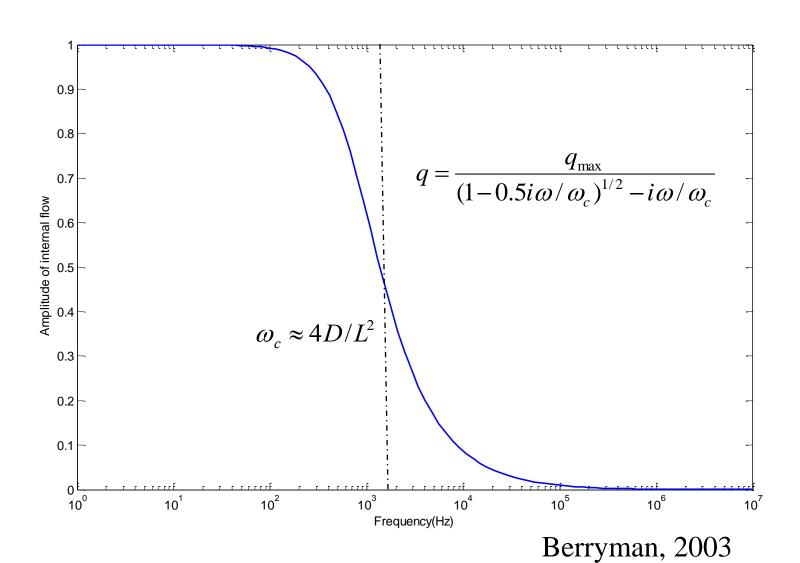
- Need geometry info on heterogeneity
- \* Navier-Stokes equation with proper boundary condition
- \* Need proper approximation
- \* Mesoscale: general format with characteristic frequency  $\omega_c$

$$q = \frac{q_{\text{max}}}{(1 - iP\omega/\omega_c)^{1/2} - i\omega/\omega_c}$$

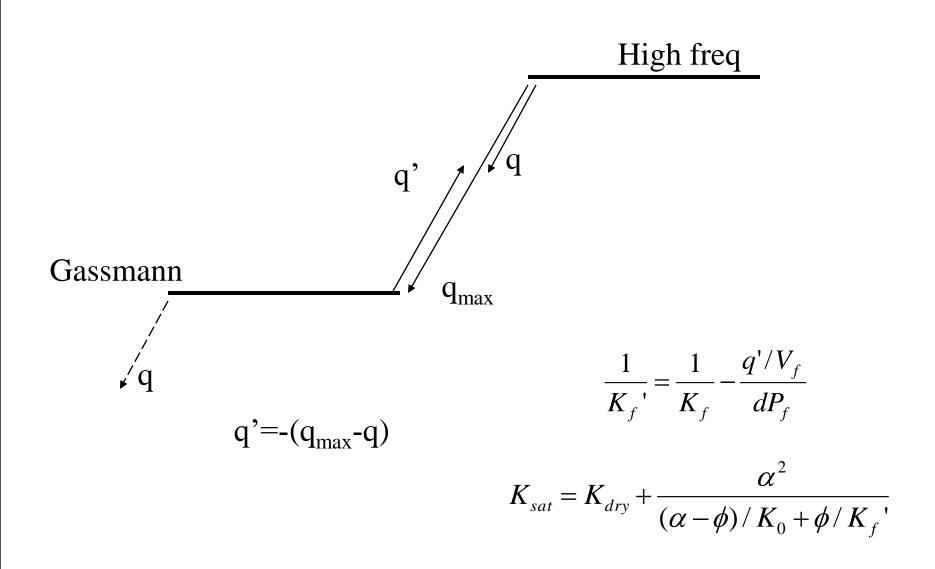
- \* Microscale:
  - **Analytical solution**
  - **Numerical solution**



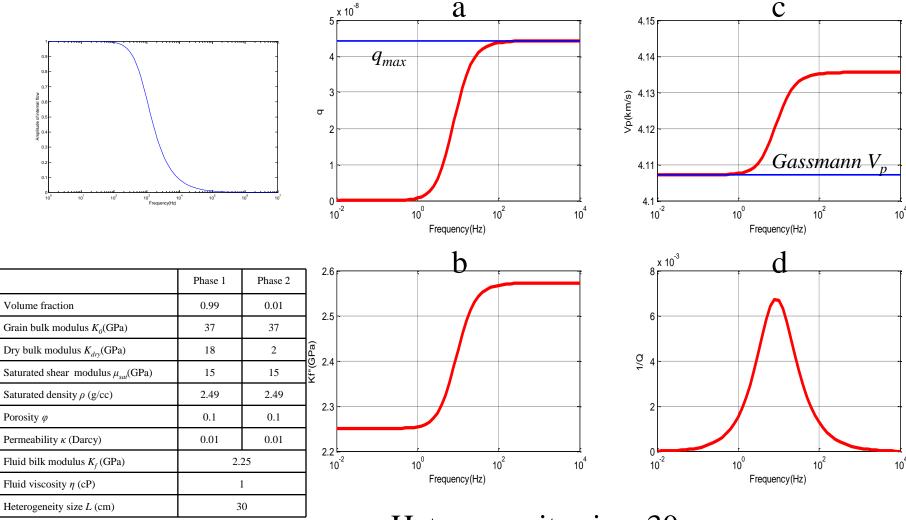
### Frequency dependency of fluid flow: q=q(f)



#### Where to start from: ∞ Hz or 0 Hz?



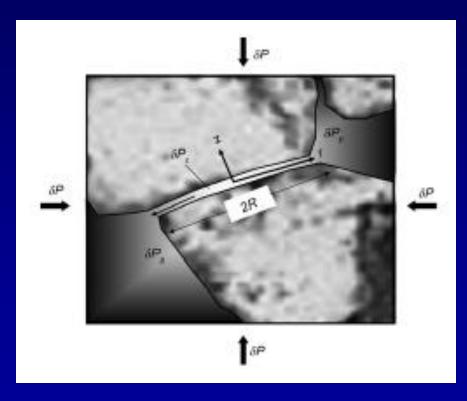
# q=q(f) in mesoscopic heterogeneity?



Heterogeneity size: 30cm

#### q=q(f) in microscopic heterogeneity?

$$\frac{q}{\delta P_{f}V_{f}} = \frac{8\pi\varepsilon(1-\nu)}{3\phi\mu} f(\zeta) \left[ \frac{\frac{1}{K_{d}} - \frac{1}{K_{s}}}{\frac{1}{K_{d}} - \frac{1}{K}} - f(\zeta) \right] / \left\{ 1 + \frac{4(1-\nu)K_{f}}{3\mu\gamma} \left[ 1 - f(\zeta) \right] \right\}$$



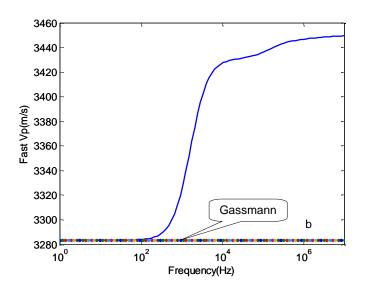
$$f(\zeta) = \frac{2J_1(\zeta)}{\zeta J_0(\zeta)}$$

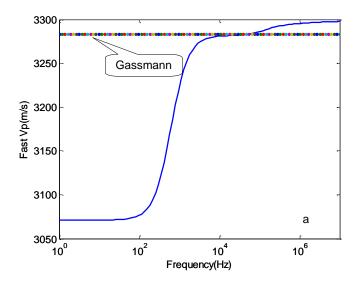
$$\zeta = \sqrt{3i\omega\eta/K_f}/\gamma$$

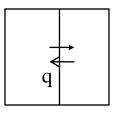
Tang 2011



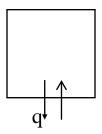
### Squirt flow dispersion, by DFM



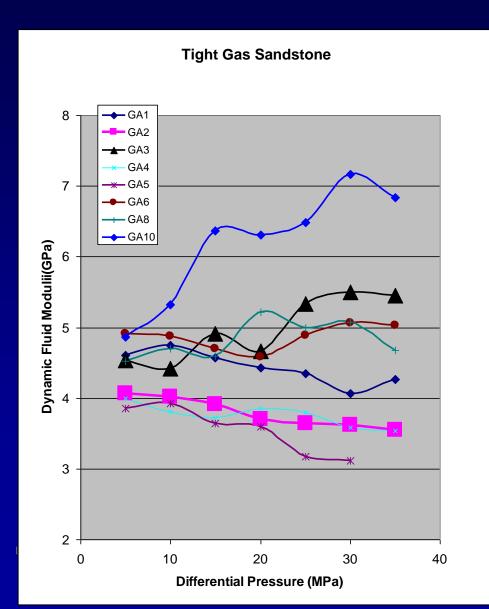


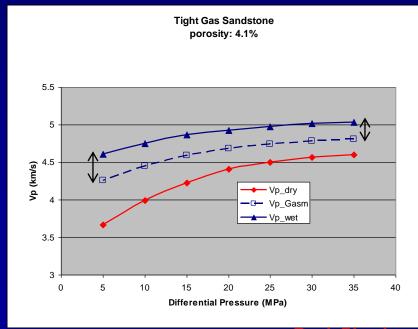


$$q'=-(q_{max}-q)$$

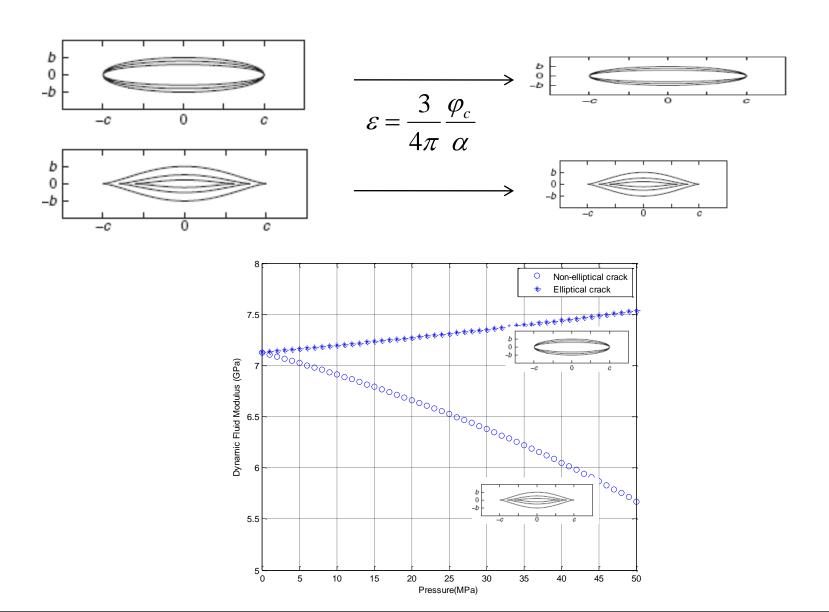


## Pressure effect on heterogeneities

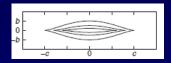


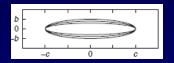


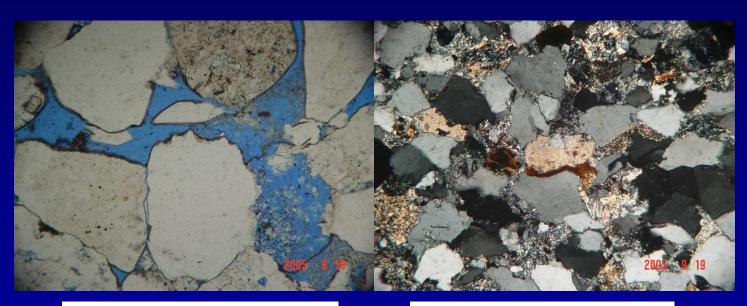
# Elliptical vs. non-elliptical cracks



# Elliptical vs. non-elliptical cracks Tight Gas Sandstone







Sample GA2: ×100

Depth :2080.78

Sample GA3:  $\times$ 100  $\perp$ 

Depth: 2072.53m



#### **Summaries**

- 1. Using "dynamic fluid modulus", Gassmann equation can be extended into heterogeneous rocks at non-zero frequency.
- 2. Explicitly link heterogeneities to dispersion and attenuation, by a fluid term.
- 3. More intuitive physical meaning.
- 4. Modeling: more powerful and flexible.
- 5. Inverting: new insight on rock microstructure



# Thank You!

