

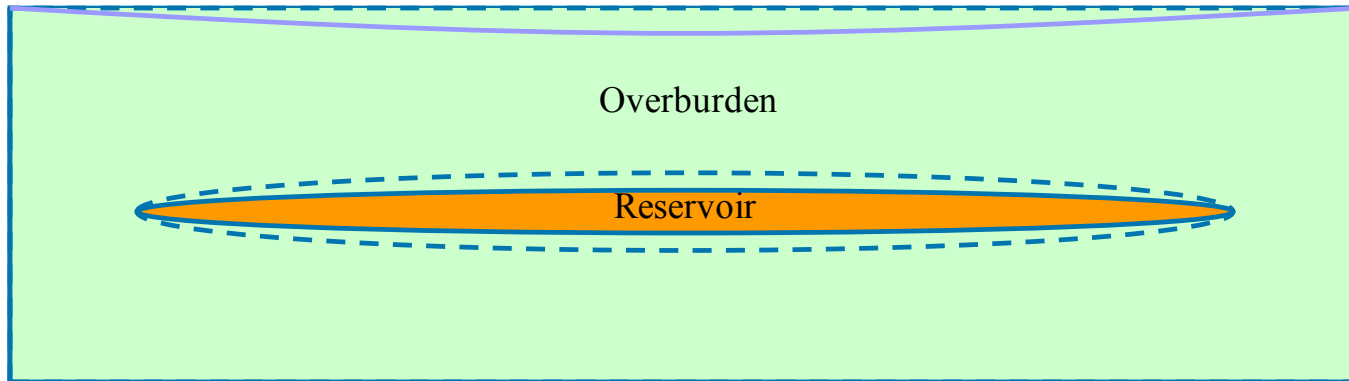
Dilation factor as function of geological time

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Introduction

4D or time-lapse travelttime analysis



$$\frac{\Delta t_0(x_0)}{t_0(x_0)} \approx \frac{\Delta z(x_0)}{z(x_0)} - \frac{\Delta v_{p0}(x_0)}{v_{p0}(x_0)}$$

(Landrø and Stammeijer, 2004)

$$\frac{\Delta v_{p0}(x_0)}{v_{p0}(x_0)} = \alpha \frac{\Delta z(x_0)}{z(x_0)}$$

(Røste et al., 2005)

$$R = -\alpha$$

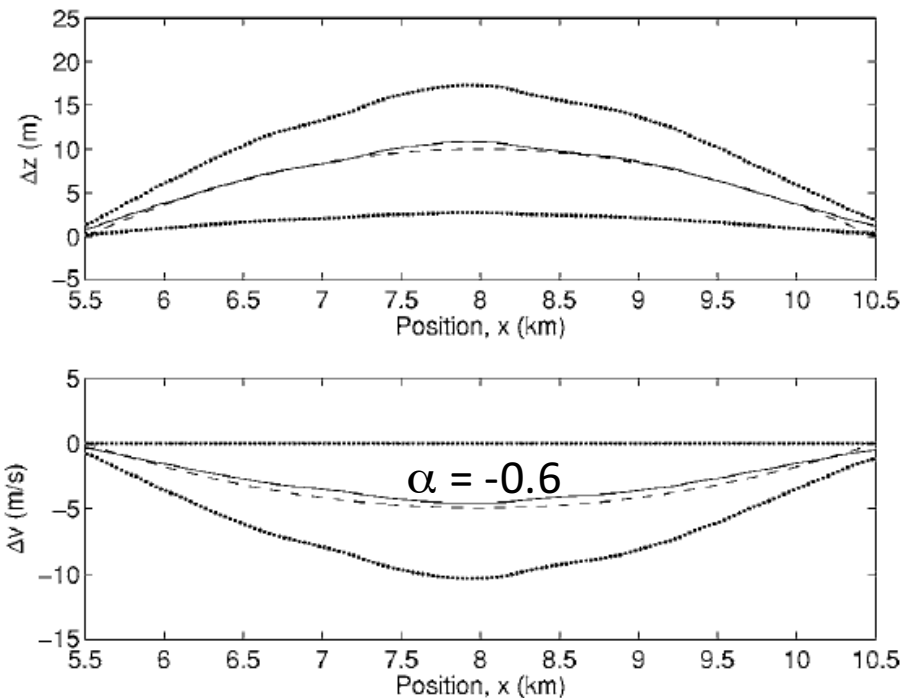
(Hatchell et al., 2005)

Relative thickness change and velocity change

$$\frac{\Delta z(x_0)}{z(x_0)} = \frac{1}{(1 - \alpha)} \frac{\Delta t_0(x_0)}{t_0(x_0)}$$

$$\frac{\Delta v_{p0}(x_0)}{v_{p0}(x_0)} = \frac{\alpha}{(1 - \alpha)} \frac{\Delta t_0(x_0)}{t_0(x_0)}$$

Optimal $\alpha = -0.6$ for all lateral positions

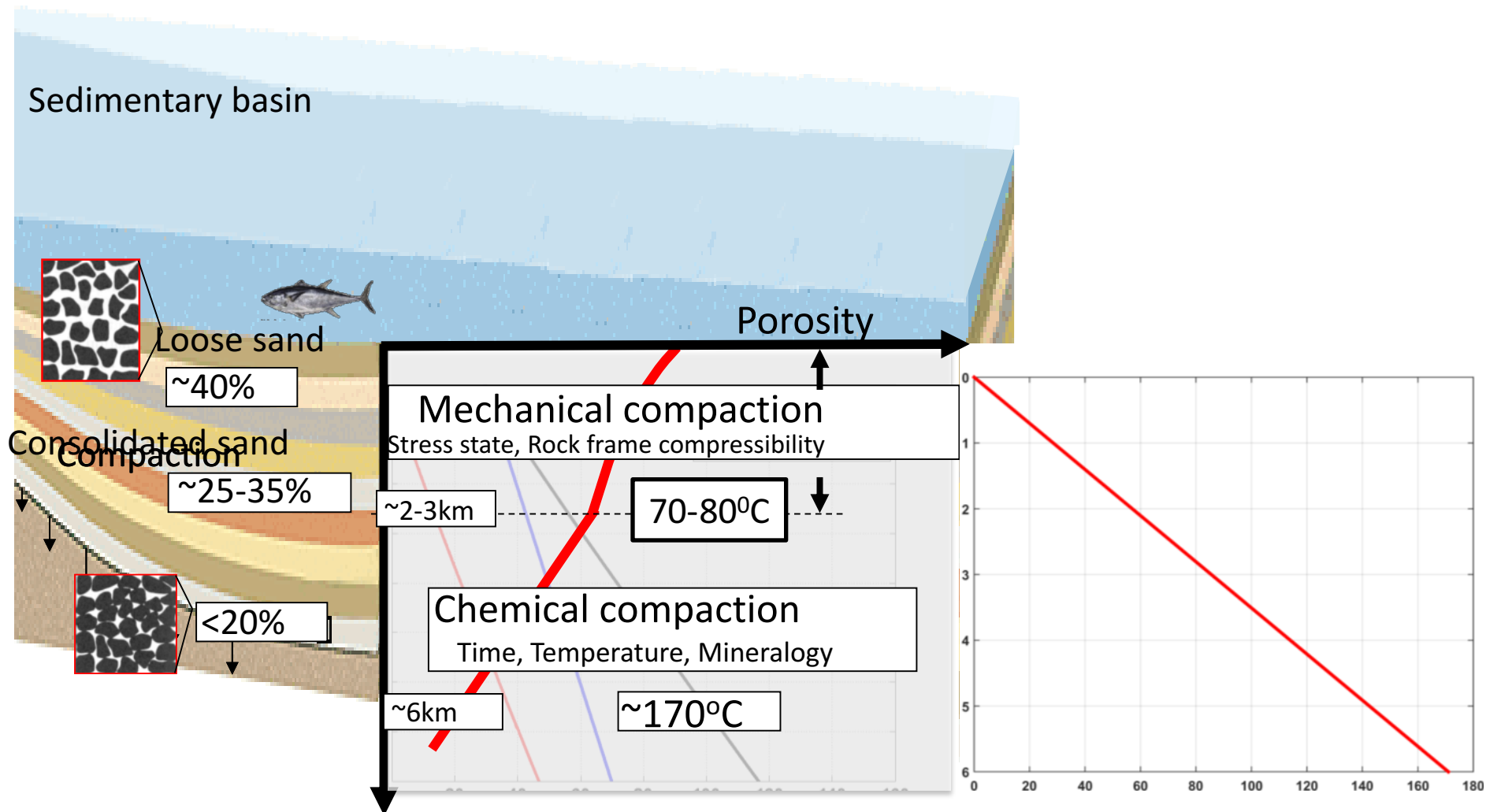


(Figure courtesy: Røste et al., 2006)

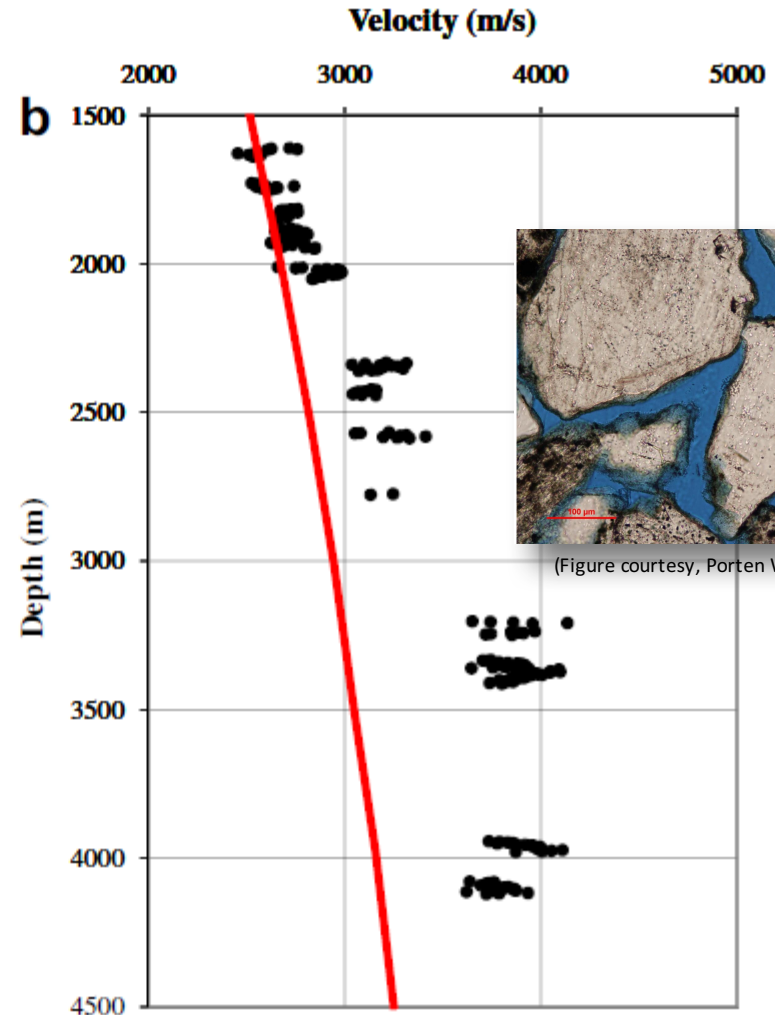
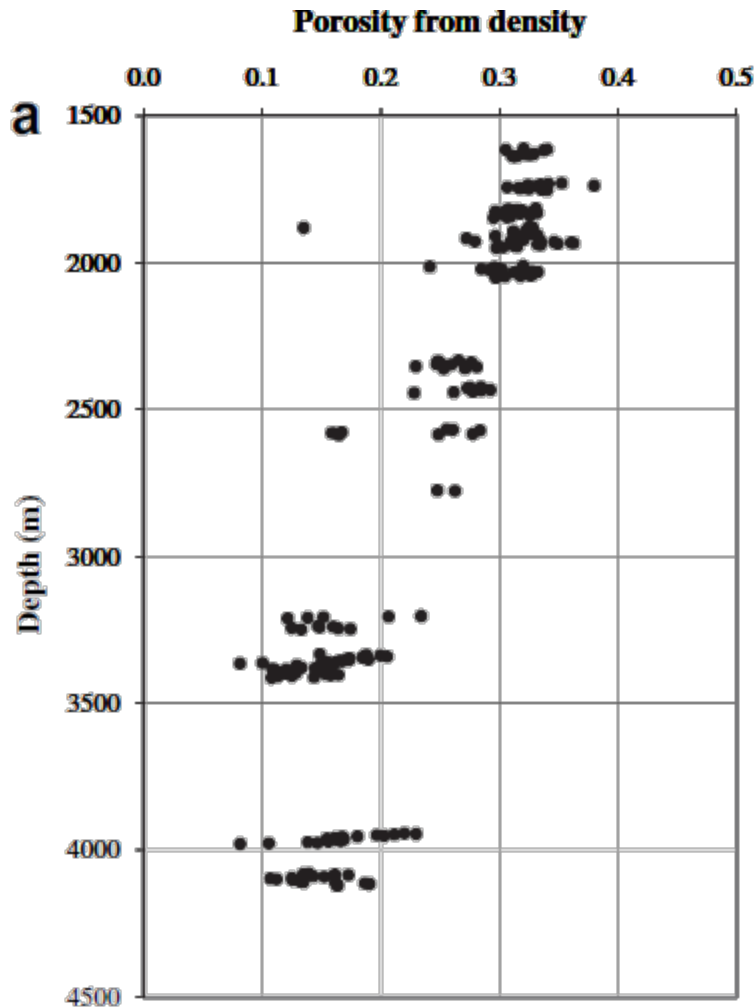
Outline

- Introduction
- The model
- Modelling results
- Conclusions

Sediment Compaction



Porosity-Velocity depth trends – North Sea



(Figure courtesy, Marcussen et al., 2010)

The model

Vertical thinning due to quartz cementation in sandstones

Volume of precipitated quartz (Walderhaug, 1996)

$$\frac{dV_q}{dt} = \frac{MrA}{\rho}$$

Total volume change

$$dV = dV_q + dV \frac{V_\phi}{V}$$

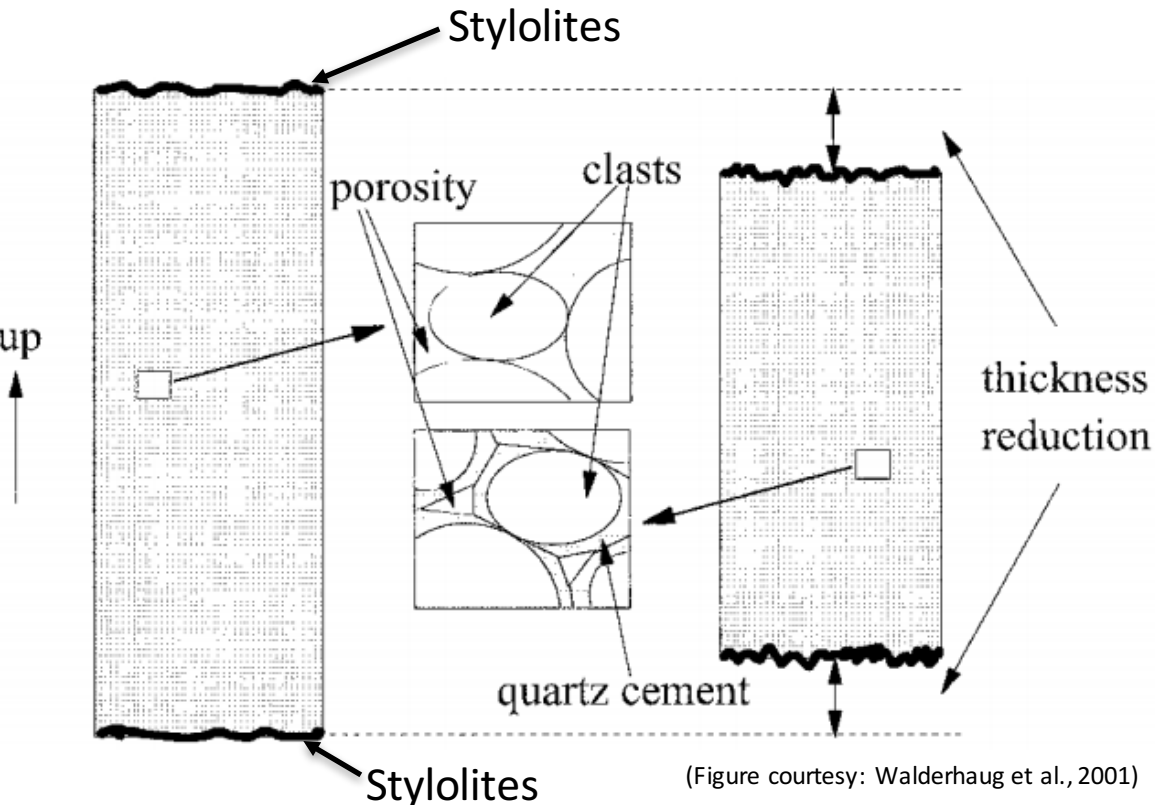
Rearrange and combine

$$\frac{dV}{dt} = -\frac{MrAV}{\rho(V - V_\phi)}$$

Surface area

$$A = A_0 \frac{V_\phi}{V_{\phi_0}}$$

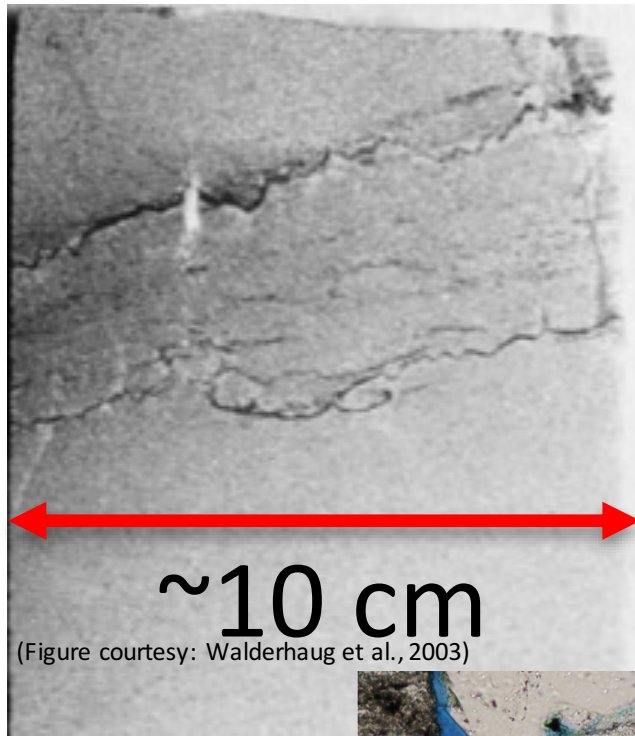
$$\frac{dV}{dt} = -\frac{MrA_0}{\rho V_{\phi_0} V_s} V(V - V_s)$$



(Figure courtesy: Walderhaug et al., 2001)

(Walderhaug et al., 2001)

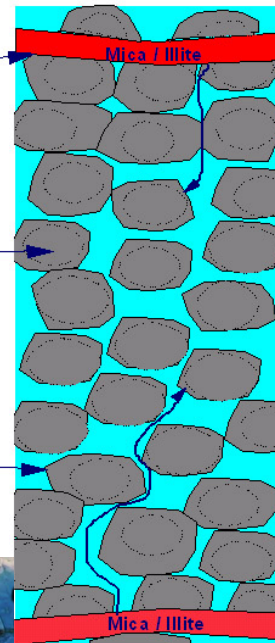
Examples of stylolites and quartz cement



Mica / Quartz
dissolution
surfaces
(stylolites)

Quartz
grain

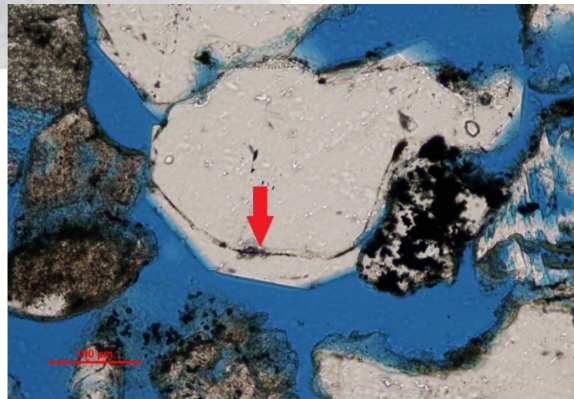
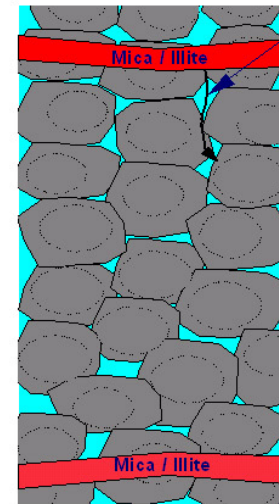
Quartz
cement



Volume reduction
due to
cementation
process

Sequentially
coupled silica
dissolution,
transport and
precipitation

Precipitation rate-
controlled
&
pressure-
insensitive



(Figure courtesy: Porten Walderhaug, 2012)

(Figure courtesy: Bjørkum, 2008)

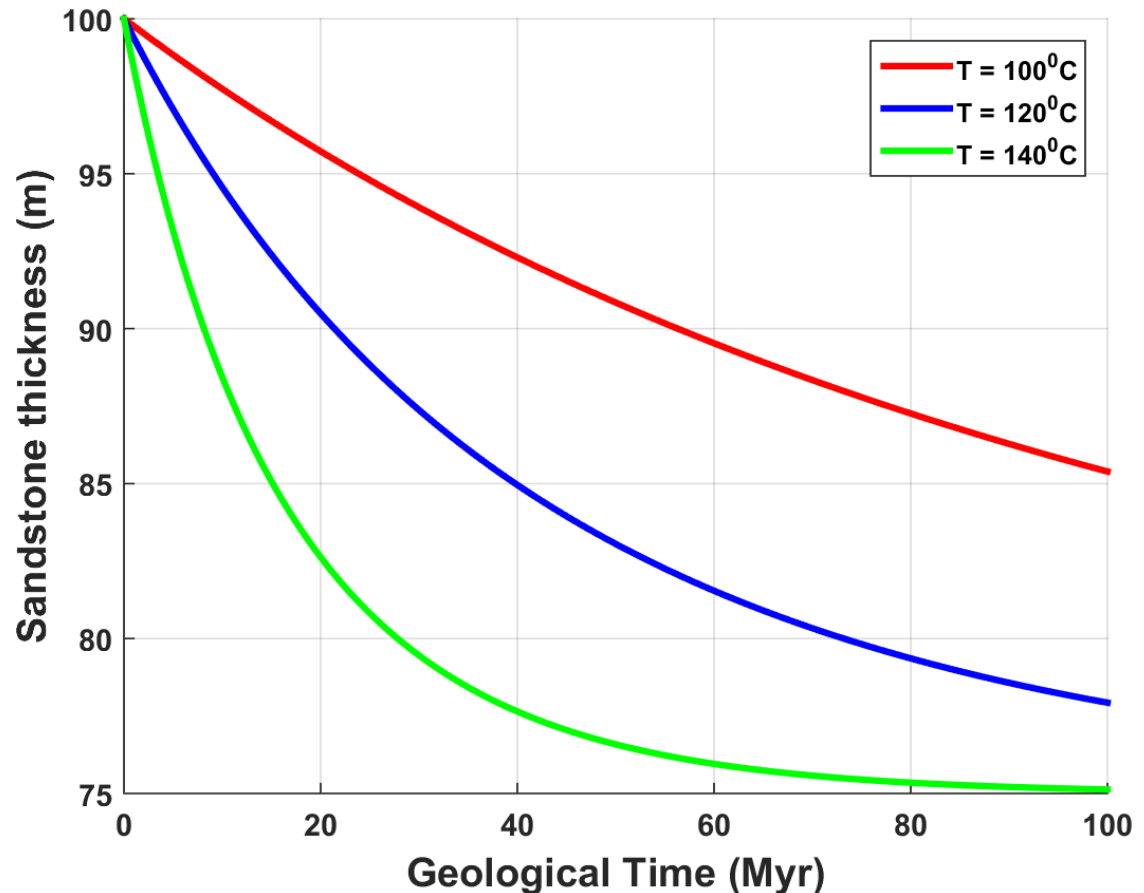
Model assumptions

- Local redistribution, no import or export of silica
- Horizontal stylolites – volume reduction is a vertical thinning
- No mechanical compaction takes place
- No other diagenetic precipitation or dissolution reactions are active

Volume change at constant temperature

$$\frac{dV}{dt} = -\frac{MrA_0}{\rho V_{\phi_0} V_s} V(V - V_s)$$

100 meter thick sandstone layer with 25% porosity @ t=0



Quartz precipitation rate

$$r \approx a10^{bT}$$

$$\int_{V_0}^V \frac{1}{v(v - V_s)} dv = -\frac{MrA_0}{\rho V_{\phi_0} V_s} \int_{t_0=0}^t d\tau$$

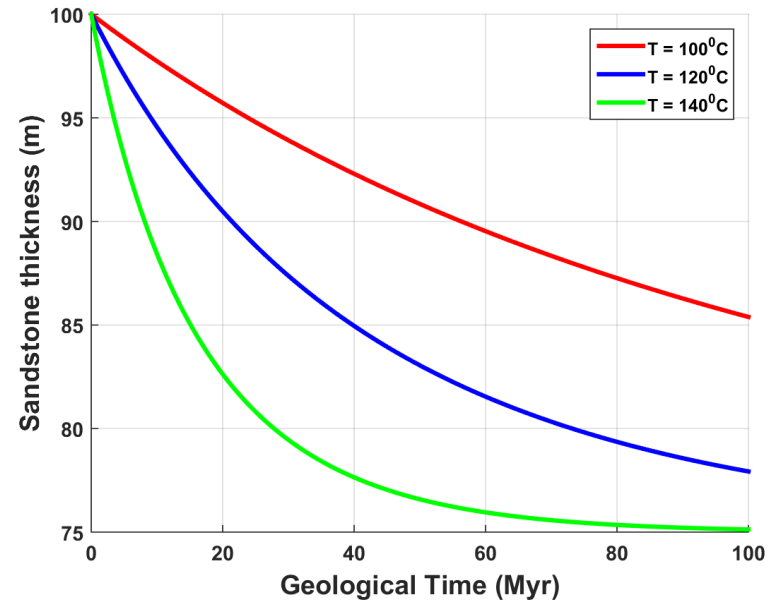
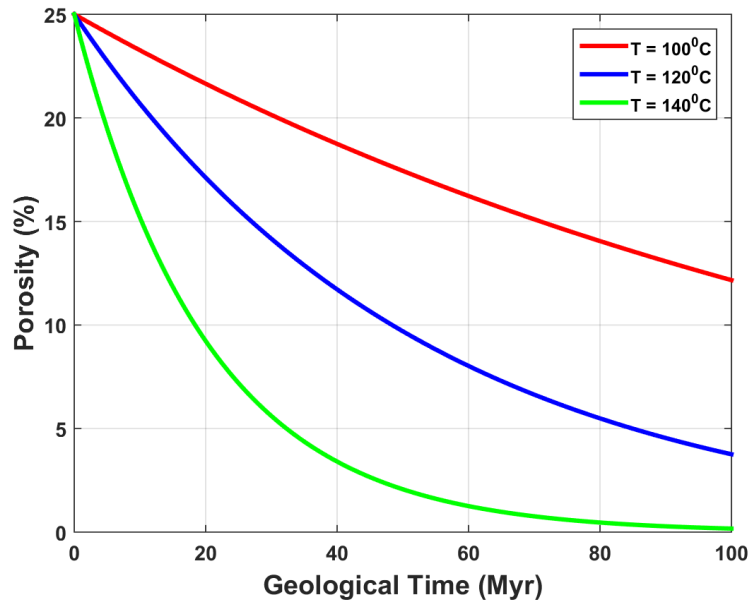
$$V = \frac{V_0 V_s e^{\left(\frac{MrA_0 t}{\rho V_{\phi_0}}\right)}}{\left(V_0 e^{\left(\frac{MrA_0 t}{\rho V_{\phi_0}}\right)} - V_{\phi_0}\right)}$$

(Walderhaug et al., 2001)

Porosity-volume change

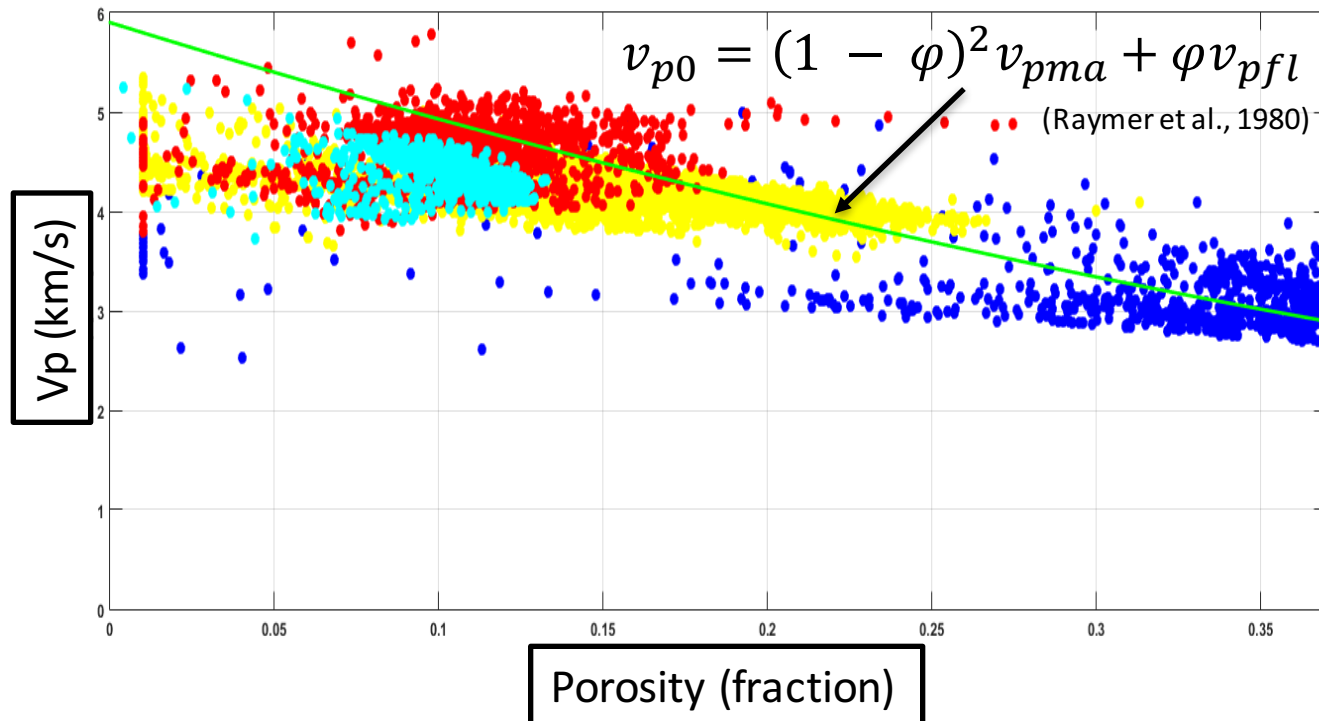
Constant temperature

100 meter thick sandstone layer with 25% porosity @ t=0



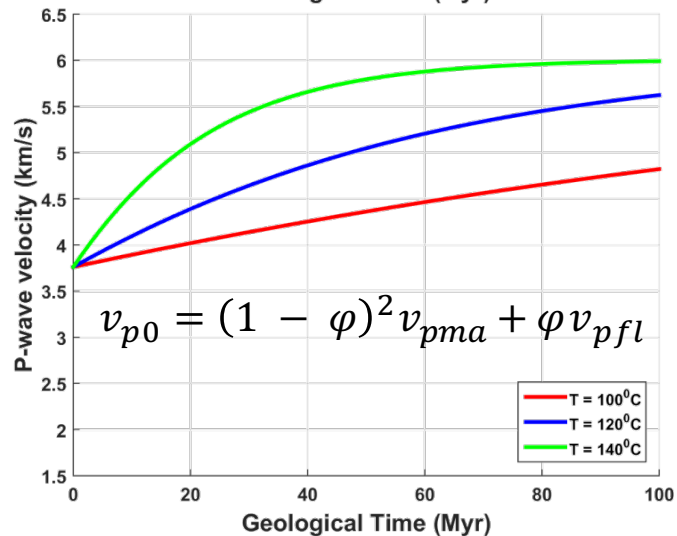
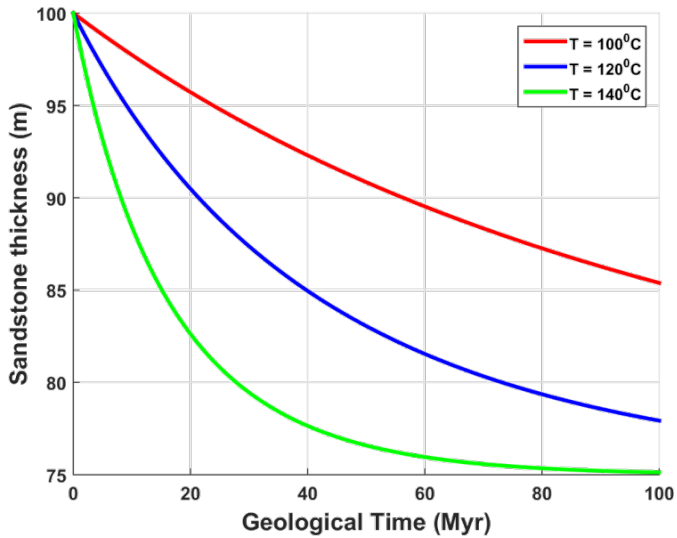
Velocity-porosity model

Consolidated sandstones from NCS

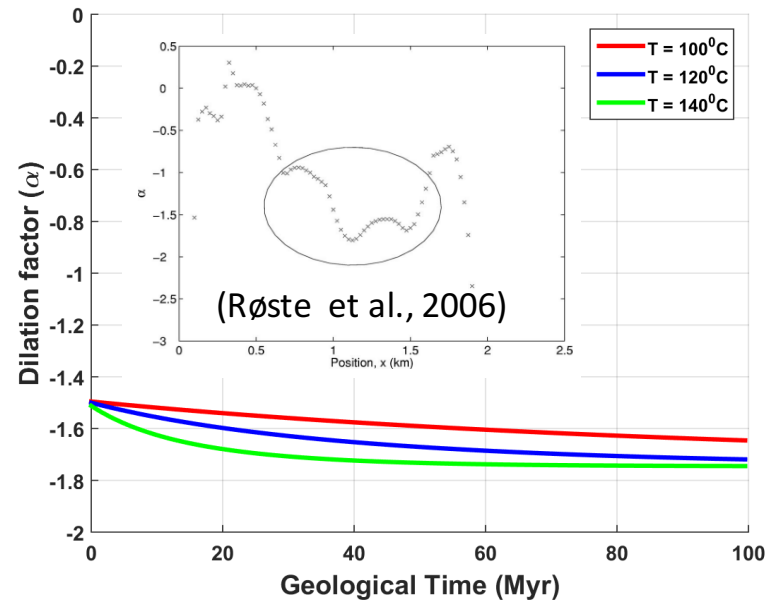


Dilation factor (α) as function of time

Constant temperature



$$\alpha = \frac{\Delta v_{p0}(x_0, t)}{v_{p0}(x_0, t)} \left(\frac{\Delta z(x_0, t)}{z(x_0, t)} \right)^{-1}$$



Volume change during a linear temperature change a.f.o. time

100 meter thick sandstone layer with 25% porosity @ t=0

$$\frac{dV}{dt} = -\frac{MrA_0}{\rho V \varphi_0 V_s} V(V - V_s)$$

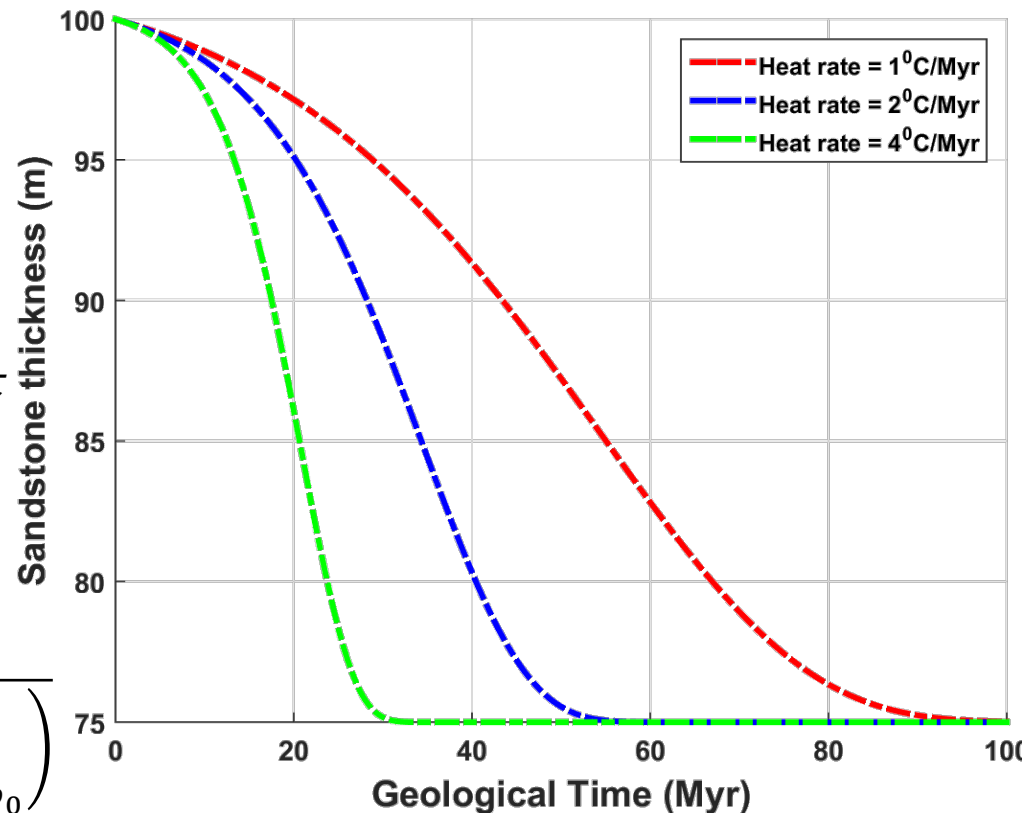
Quartz precipitation rate

$$r \approx a10^{b(ct+d)}$$

$$\int_{V_0}^V \frac{1}{v(v - V_s)} dv = -\frac{MA_0 a}{\rho V \varphi_0 V_s} \int_{t_0=0}^t 10^{b(ct+d)} d\tau$$

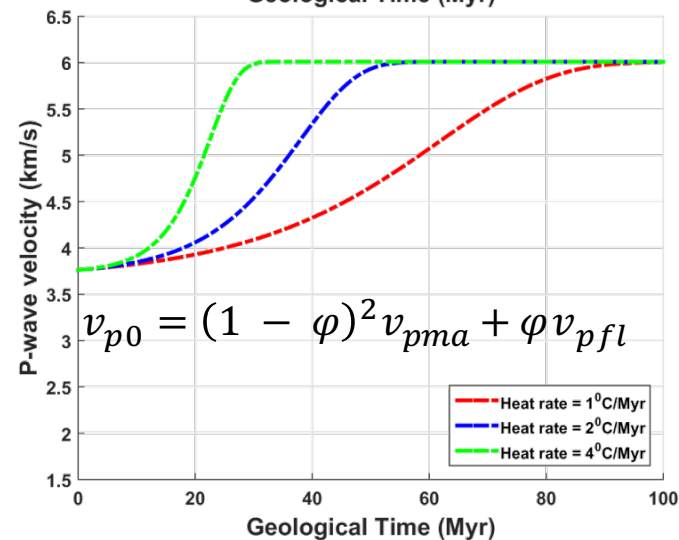
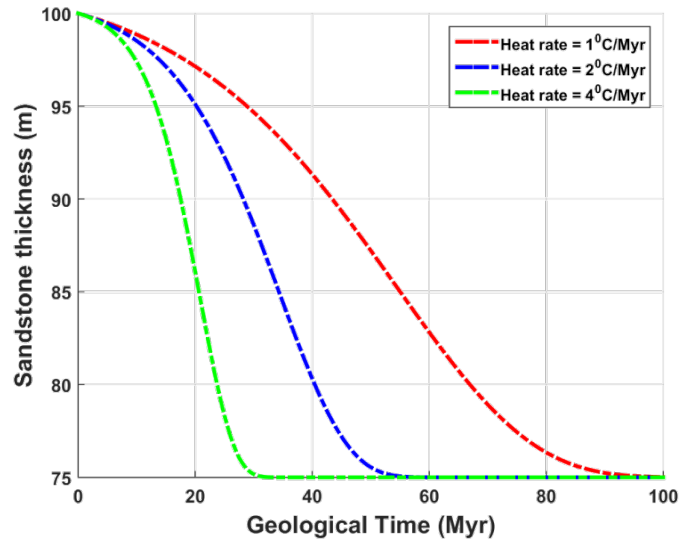
$$V = \frac{V_0 V_s e^{\left(\frac{MA_0 a}{\rho V \varphi_0 b c \ln(10)} (10^{b(ct+d)} - 10^{bd})\right)}}{\left(V_0 e^{\left(\frac{MA_0 a}{\rho V \varphi_0 b c \ln(10)} (10^{b(ct+d)} - 10^{bd})\right)} - V_{\varphi_0}\right)}$$

(Walderhaug et al., 2001)

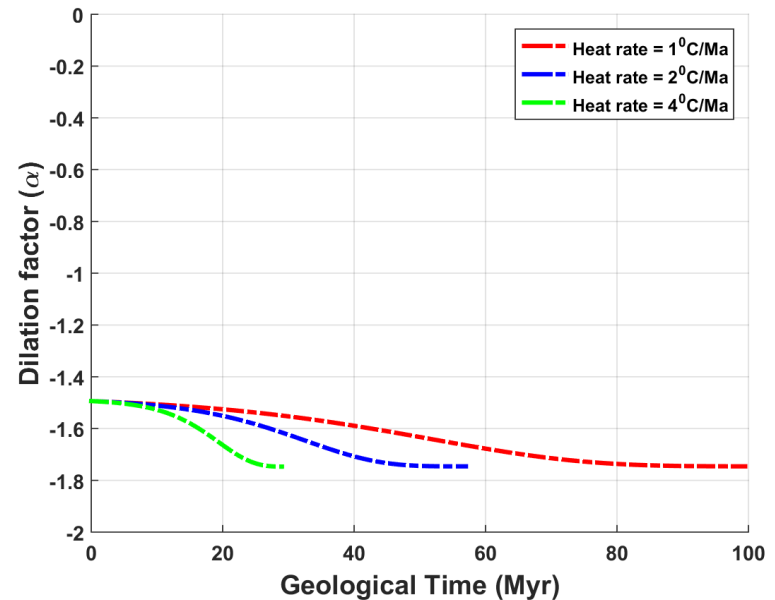


Dilation factor as function of time

Linear temperature change

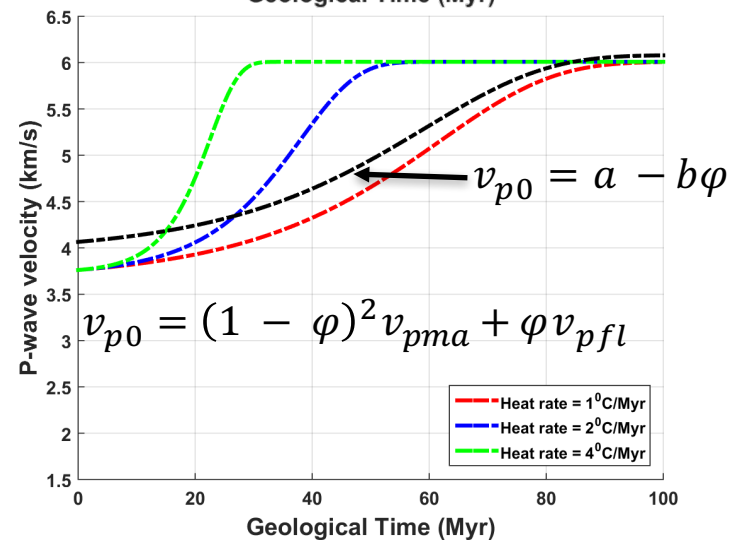
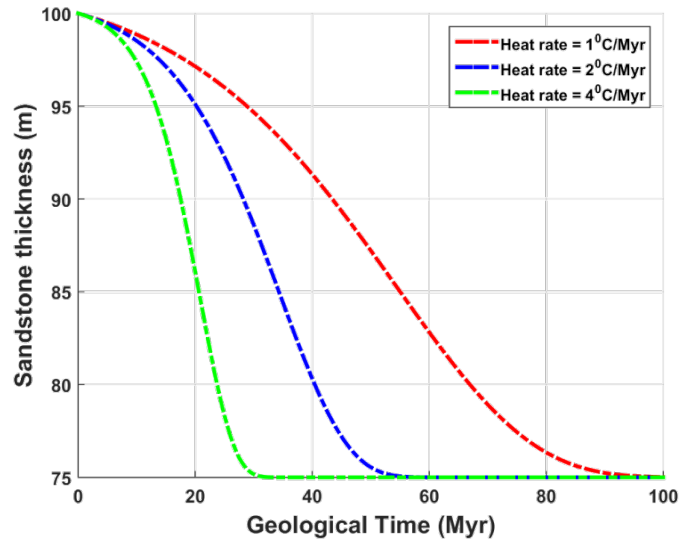


$$\alpha = \frac{\Delta v_{p0}(x_0, t)}{v_{p0}(x_0, t)} \left(\frac{\Delta Z(x_0, t)}{z(x_0, t)} \right)^{-1}$$

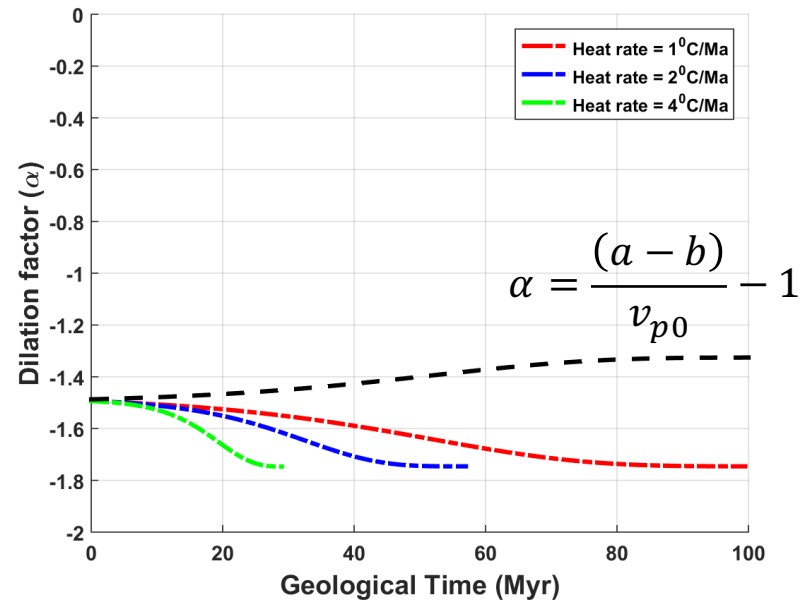


Dilation factor as function of time

Assuming a linear temperature change



$$\alpha = \frac{\Delta v_{p0}(x_0, t)}{v_{p0}(x_0, t)} \left(\frac{\Delta Z(x_0, t)}{Z(x_0, t)} \right)^{-1}$$



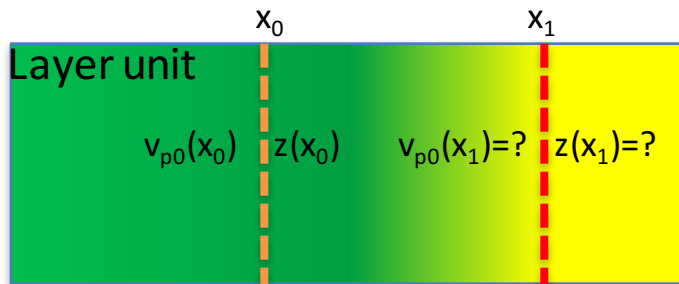
Conclusions

- The dilation factor (α) of the reservoir is described as function of the rate of vertical thinning of sandstones due to quartz cementation.
- Two cases are tested
 - Constant temperature
 - Linear temperature increase
- Given the model assumptions
 - The dilation factor range between -1.5 to -1.75

Spatial zero offset travelttime analysis

RoSe presentation 2015

t_0 = two-way vertical time thickness of unit at x_0
 x_0 = coordinate reference position along a line
 x_1 = a new coordinate position along the line
 z = thickness of formation unit
 v_{p0} = vertical P-wave velocity of unit
 Δ = spatial difference in physical parameters
 α = Dilation factor



$$t_0(x_0) = \frac{2z(x_0)}{v_{p0}(x_0)}$$

$$\frac{\Delta t_0(x_1, x_0)}{t_0(x_0)} \approx \frac{\Delta z(x_1, x_0)}{z(x_0)} - \frac{\Delta v_{p0}(x_1, x_0)}{v_{p0}(x_0)}$$

$$\frac{\Delta v_{p0}(x_1, x_0)}{v_{p0}(x_0)} = \alpha(x_0) \frac{\Delta z(x_1, x_0)}{z(x_0)}$$

$$\frac{\Delta z(x_1, x_0)}{z(x_0)} = \frac{1}{(1 - \alpha(x_0))} \frac{\Delta t_0(x_1, x_0)}{t_0(x_1, x_0)}$$

$$\frac{\Delta v_{p0}(x_1, x_0)}{v_{p0}(x_0)} = \frac{\alpha(x_0)}{(1 - \alpha(x_0))} \frac{\Delta t_0(x_1, x_0)}{t_0(x_1, x_0)}$$

Acknowledgement

- RoSe for giving me the opportunity to present

Thank you for your attention