

# Spatial reservoir characterization using the dilation factor $\alpha$

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# Outline

- Introduction
- Spatial traveltime analysis - a new application?
- Geological scenarios and numerical examples
- Conclusions

# Introduction

## 4D or time-lapse traveltimes analysis

$$t_0(x_0) = \frac{2z(x_0)}{v_{p0}(x_0)}$$

$t_0$  = two-way vertical time thickness of unit

$x_0$  = coordinate position along a line

$z$  = thickness of formation unit

$v_{p0}$  = vertical P-wave velocity of unit

$\Delta$  = changes in physical parameters

and  $R$  = ratio between relative velocity

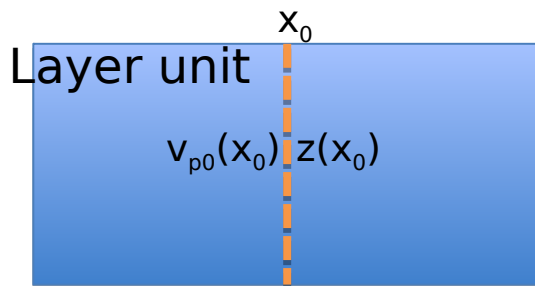
and thickness changes

$$\frac{\Delta t_0(x_0)}{t_0(x_0)} \approx \frac{\Delta z(x_0)}{z(x_0)} - \frac{\Delta v_{p0}(x_0)}{v_{p0}(x_0)}$$

(Landrø and Stammeijer 2004)

$$\frac{\Delta v_{p0}(x_0)}{v_{p0}(x_0)} = \alpha \frac{\Delta z(x_0)}{z(x_0)}$$

(Røste et al., 2005)



$$\frac{\Delta v_{p0}(x_0)}{v_{p0}(x_0)} = -R \frac{\Delta z(x_0)}{z(x_0)}$$

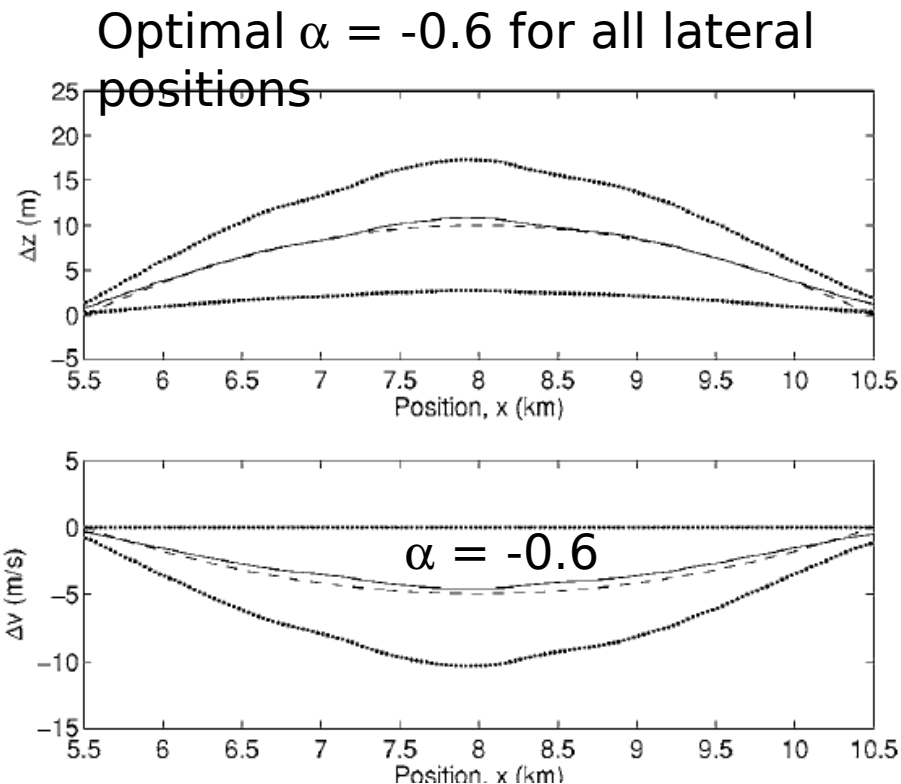
(Hatchell et al., 2005)

# Introduction

## Relative changes in layer thickness and velocity

$$\frac{\Delta z(x_0)}{z(x_0)} = \frac{1}{(1 - \alpha)} \frac{\Delta t_0(x_0)}{t_0(x_0)}$$

$$\frac{\Delta v_{p0}(x_0)}{v_{p0}(x_0)} = \frac{\alpha}{(1 - \alpha)} \frac{\Delta t_0(x_0)}{t_0(x_0)}$$



(Figure courtesy: Røste et al., 2006)

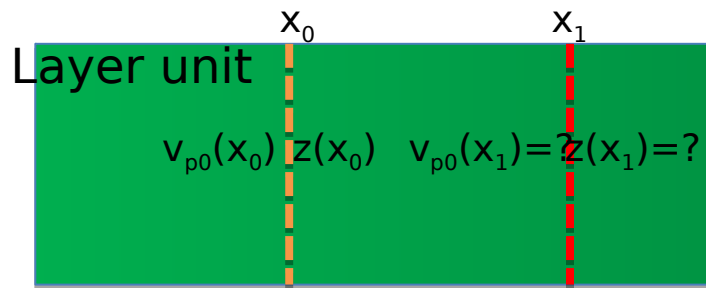
# Spatial zero offset traveltimes analysis

- = two-way vertical time thickness of unit at  $x_0$
- = coordinate reference position along a line
- = a new coordinate position along the line
- = thickness of formation unit
- $v_{p0}$  = vertical P-wave velocity of unit
- = spatial difference in physical parameters
- = Dilation factor

$$t_0(x_0) = \frac{2z(x_0)}{v_{p0}(x_0)}$$

$$\frac{\Delta t_0(x_1, x_0)}{t_0(x_0)} \approx \frac{\Delta z(x_1, x_0)}{z(x_0)} - \frac{\Delta v_{p0}(x_1, x_0)}{v_{p0}(x_0)}$$

$$\frac{\Delta v_{p0}(x_1, x_0)}{v_{p0}(x_0)} = \alpha(x_0) \frac{\Delta z(x_1, x_0)}{z(x_0)}$$



# Porosity-strain relation

Assume only changes in pore volume

Porosity  $\varphi = \frac{V_b - V_s}{V_b} = 1 - \frac{V_s}{V_b}$  **Solid volume**  
**Bulk volume**

Rewriting we get

$$\frac{V_s}{V_b} = 1 - \varphi$$

Differentiation gives

$$\frac{d\varphi}{dV_b} = \frac{V_s}{V_b^2} \Rightarrow d\varphi = \frac{V_s}{V_b} \frac{dV_b}{V_b} \quad \text{where} \quad \frac{dV_b}{V_b} = \epsilon_{vol}$$

We get porosity-strain relation

$$\frac{d\varphi}{(1 - \varphi)} = \epsilon_{vol}$$

# Spatial varying layer thickness and velocity

Relative changes in layer thickness and velocity

$$\frac{\Delta z(x_1, x_0)}{z(x_0)} = \frac{1}{(1 - \alpha(x_0))} \frac{\Delta t_0(x_1, x_0)}{t_0(x_1, x_0)}$$

$$\frac{\Delta v_{p0}(x_1, x_0)}{v_{p0}(x_0)} = \frac{\alpha(x_0)}{(1 - \alpha(x_0))} \frac{\Delta t_0(x_1, x_0)}{t_0(x_1, x_0)}$$

$$\frac{\Delta v_{p0}(x_1, x_0)}{v_{p0}(x_0)} = \alpha(x_0) \frac{\Delta \varphi(x_1, x_0)}{(1 - \varphi(x_0))}$$

# Dilation factor in clean and shaly sandstone Uniaxial deformation

Clean sandstone

$$v_{p0} = a - b\phi$$

$$\frac{d\phi}{(1-\phi)} = \frac{dz}{z}$$

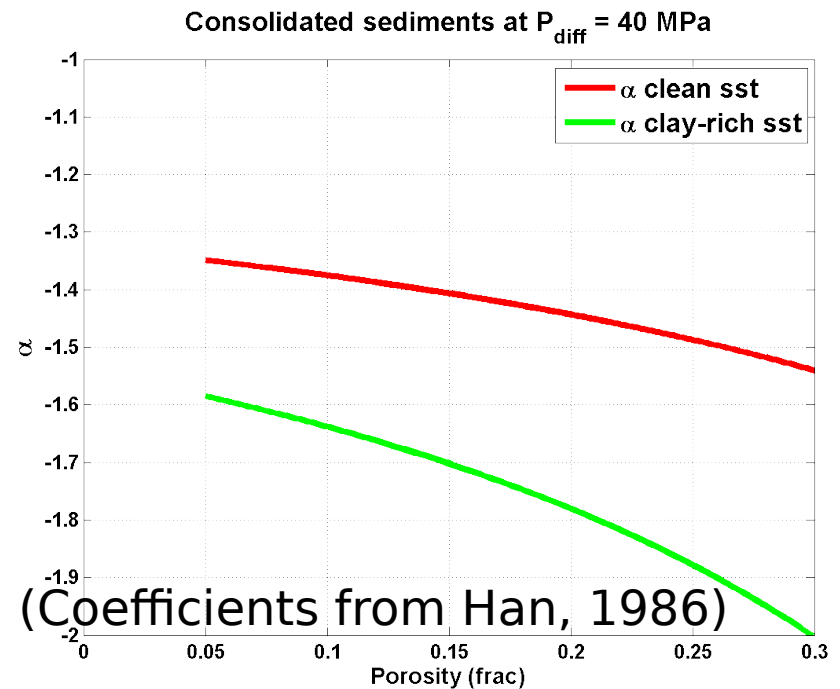
$$\frac{\Delta\phi}{(1-\phi)} = \frac{1}{\alpha} \frac{\Delta v_{p0}}{v_{p0}}$$

$$\alpha = \frac{(a-b)}{v_{p0}} - 1 \quad (\text{R\o{ost}e et al., 2006})$$

Clay-rich sandstone

$$v_{p0} = a - b\phi - cv_{cl}$$

$$\alpha = \frac{(\phi-1)b}{v_{p0}} \quad (\text{Carcione et al., 2007})$$

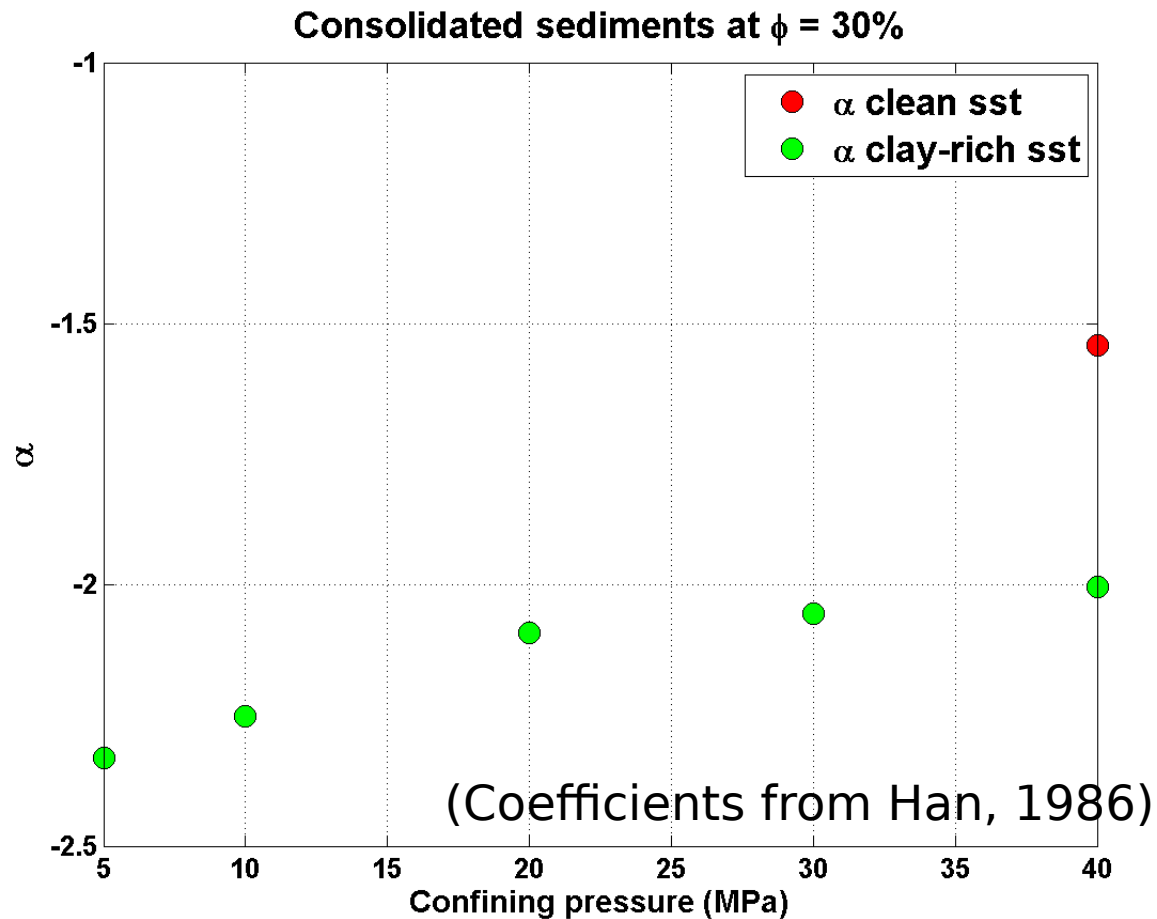


$\alpha$  decreases with decreasing porosity 



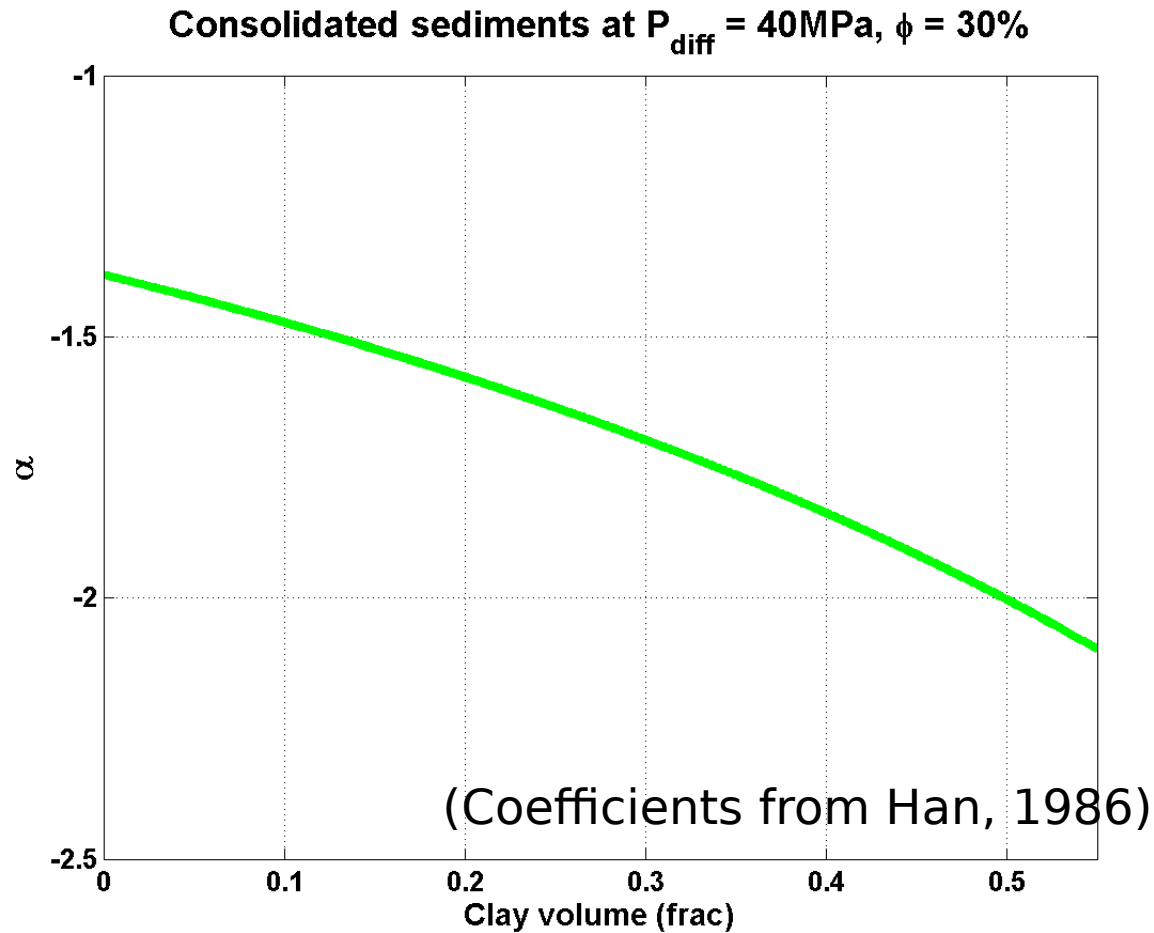
# Dilation factor vs. stress

## Clean and clay-rich sandstone



$\alpha$  decreases with increasing net stress 

# Dilation factor vs. volume of clay



$\alpha$  decreases with decreasing clay content 

# Geological scenarios - numerical examples

- Plausible geological reasons causing spatially variation in layer thickness and velocity in a unit
  - Equal depositional layer thickness followed by “differential compaction” or diagenetic effects within a formation unit
  - Erosion
  - Lithology changes
    - clean vs. shaly sandstone

# Spatial porosity and thickness change within unit

## Differential compaction of clean sandstone

$x_0$  location with a well

$$\varphi = 24\%$$

$$Z = 33\text{m}$$

$x_1$  location without a well

$$\varphi = 20\%$$

$$Z = 31\text{m}$$

$$v_{p0}(x_0) = 6.08 - 8.06 * 0.24 \approx 4146\text{m/s}$$

$$v_{p0}(x_1) = 6.08 - 8.06 * 0.20 = 4468\text{m/s}$$

$$t_0(x_0) = \frac{2z(x_0)}{V_{p0}(x_0)}$$

$$t_0(x_0) = \frac{66\text{m}}{4146\text{m/s}} \approx 0.01592\text{s}$$

$$t_0(x_1) = \frac{62\text{m}}{4468\text{m/s}} \approx 0.01388\text{s}$$

$$\alpha_0 = \frac{(a-b)}{v_{p0}} - 1 = \frac{(6.08-8.06)}{4.146} - 1 \approx -1.48$$

# Estimates of layer thickness and velocity at new location

## Case: Differential compaction

**Layer thickness estimate at  $x_1$  location:**

$$\frac{\Delta z(x_1, x_0)}{z(x_0)} = \frac{1}{2.48} \frac{(0.01388 - 0.01592)}{(0.01592)} \approx -0.0517$$

$$\tilde{z}(x_1) = z(x_0) + \Delta z(x_1, x_0) = 33 \text{ m} - 1.71 \text{ m} \approx 31.3 \text{ m}$$

Correct  $z(x_1) = 31 \text{ m}$

**Velocity estimate at  $x_1$  location:**

$$\frac{\Delta v_{p0}(x_1, x_0)}{v_{p0}(x_0)} = \frac{-1.48}{2.48} \frac{(0.01388 - 0.01592)}{(0.01592)} \approx 0.07647$$

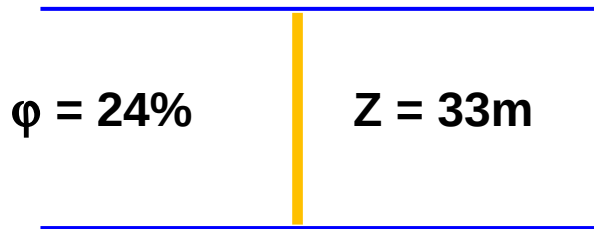
$$\tilde{v}_{p0}(x_1) = v_{p0}(x_0) + \Delta v_{p0}(x_1, x_0) = 4146 + 317 = 4463 \text{ m/s}$$

Correct  $v_{p0}(x_1) = 4468 \text{ m/s}$

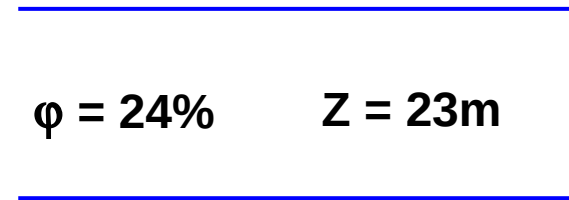
# Spatially varying layer thickness

## Erosion of a clean sandstone unit

$x_0$  location with a well



$x_1$  location without a well



$$v_{p0}(x_0) = 6.08 - 8.06 * 0.24 \approx 4146 \text{ m/s}$$

$$v_{p0}(x_1) = 6.08 - 8.06 * 0.24 = 4146 \text{ m/s}$$

$$t_0(x_0) = \frac{2z(x_0)}{V_{p0}(x_0)}$$

$$t_0(x_0) = \frac{66 \text{ m}}{4146 \text{ m/s}} \approx 0.01592 \text{ s}$$

$$t_0(x_1) = \frac{46 \text{ m}}{4146 \text{ m/s}} \approx 0.01110 \text{ s}$$

$$\alpha_0 = \frac{(a-b)}{v_{p0}} - 1 = \frac{(6.08-8.06)}{4.146} - 1 \approx -1.48$$

# Estimates of layer thickness and velocity at new location

## Case: Erosion

**Layer thickness and porosity estimate at  $x_1$  location:**

$$\frac{\Delta z(x_1, x_0)}{z(x_0)} = \frac{1}{2.48} \frac{(0.01110 - 0.01592)}{(0.01592)} \approx -0.12221$$

$$\tilde{z}(x_1) = z(x_0) + \Delta z(x_1, x_0) = 33\text{m} - 4\text{m} = 29\text{m}$$

Correct  $v_{p0}$  = 23m

**Velocity estimate at  $x_1$  location:**

$$\frac{\Delta v_{p0}(x_1, x_0)}{v_{p0}(x_0)} = \frac{-1.48}{2.48} \frac{(0.01110 - 0.01592)}{(0.01592)} \approx 0.1809$$

$$\tilde{v}_{p0}(x_1) = v_{p0}(x_0) + \Delta v_{p0}(x_1, x_0) = 4146\text{m/s} + 750\text{m/s} = 3396\text{m/s}$$

Correct  $v_{p0}$  = 4146m/s

# Spatial varying lithology clean vs. shaly sandstone unit

$x_0$  location with a well

$$\varphi = 24\%$$

$$Z = 33\text{m}$$

$x_1$  location without a well

$$\varphi = 12.5\% \quad Z = 28\text{m}$$

$$V_{cl} = 30\%$$

$$v_{p0}(x_0) = 6.08 - 8.06 * 0.24 \sim 4146\text{m/s} \quad v_{p0}(x_1) = 5.59 - 6.93 * 0.125 - 2.18 * 0.30 = 4070\text{m/s}$$

$$t_0(x_0) = \frac{2z(x_0)}{V_{p0}(x_0)}$$

$$t_0(x_0) = \frac{66\text{m}}{4146\text{m/s}} \approx 0.01592\text{s}$$

$$t_0(x_1) = \frac{56\text{m}}{4070\text{m/s}} \approx 0.01376\text{s}$$

$$\alpha_0 = \frac{(a-b)}{v_{p0}} - 1 = \frac{(6.08-8.06)}{4.146} - 1 \approx -1.48$$



# Estimates of layer thickness and velocity at new location

## Case: Lithological change

**Layer thickness and porosity estimate at  $x_1$  location:**

$$\frac{\Delta z(x_1, x_0)}{z(x_0)} = \frac{1}{2.48} \frac{(0.01376 - 0.01592)}{(0.01592)} \approx -0.0547$$

$$\tilde{z}(x_1) = z(x_0) + \Delta z(x_1, x_0) = 33\text{m} - 1.8\text{m} \approx 31.2\text{m}$$

Correct  $v_{p0}$  = 28m

**Velocity estimate at  $x_1$  location:**

$$\frac{\Delta v_{p0}(x_1, x_0)}{v_{p0}(x_0)} = \frac{-1.48}{2.48} \frac{(0.01376 - 0.01592)}{(0.01592)} \approx 0.081$$

$$\tilde{v}_{p0}(x_1) = v_{p0}(x_0) + \Delta v_{p0}(x_1, x_0) = 4146\text{m/s} + 336\text{m/s} \approx 4482\text{m/s}$$

Correct  $v_{p0}$  = 4069m/s

# Conclusions

- New approach tested by estimating relative changes in layer thickness and velocity using the dilation factor ( $\alpha$ ) and spatially traveltimes differences of a unit
  - A few numerical examples are shown.
    - Only the “differential compaction” case gave good estimates of laterally variable layer thickness (porosity) and velocity
- Method depends on
  - reference location
  - porosity-strain relation of the unit
- $\alpha$  is not constant when assuming Han’s model. It decreases as
  - Porosity decreases (clean and shaly sandstones)
  - Clay content decreases
  - Net stress increases (shaly sandstone)

# Acknowledge

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Thank you for your attention