

Quantitative analysis of CO₂ saturation at the Snøhvit field from analytical solutions and time lapse seismic data

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Objective

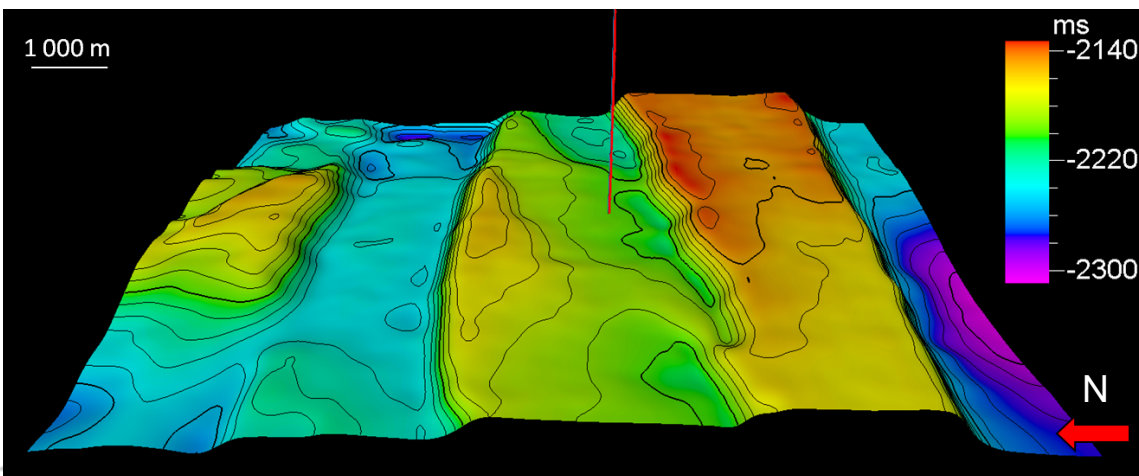
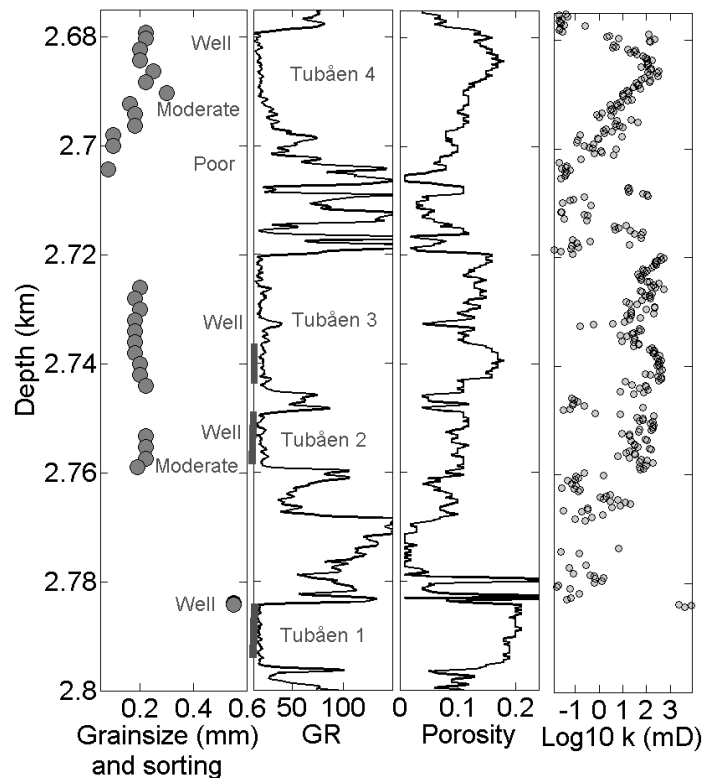
Quantify initial 400 ktons of CO₂ injected in the Tubåen 1 sand layer based on analytical solution and two time lapse seismic methods

Outline

- Reservoir fluid flow
- Analytical estimate of CO₂ plume
- CO₂ plume quantification from time lapse seismic data
 - Inverted from amplitude versus offset (AVO)
 - Spectral decomposition
- Conclusions

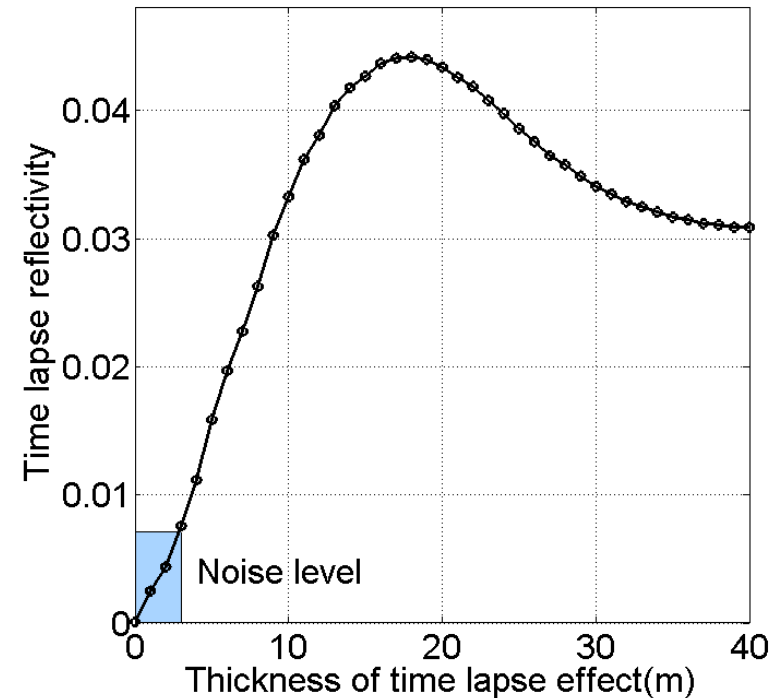
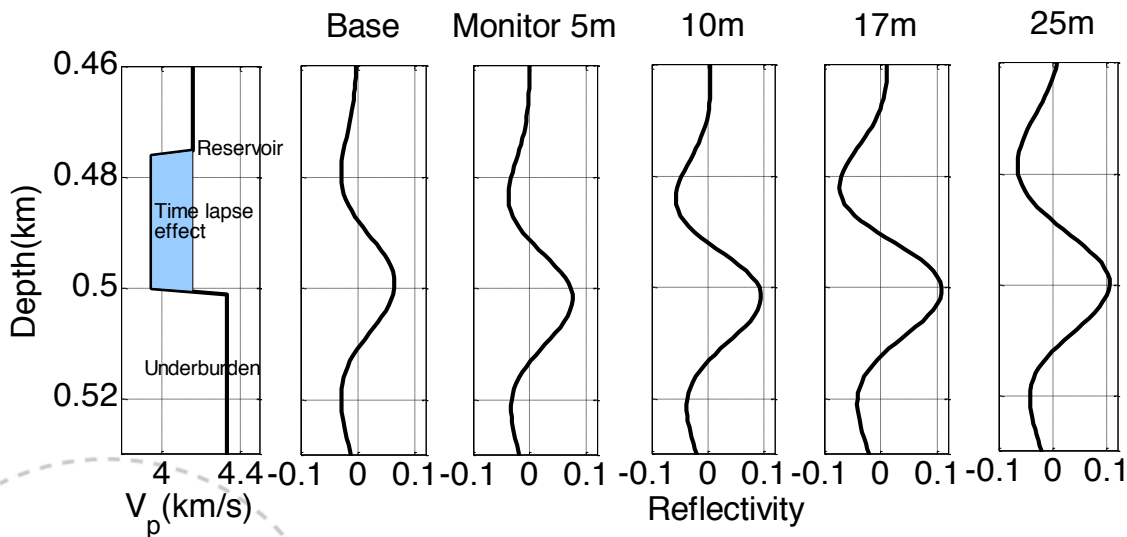
Tubåen formation, Snøhvit field

- Faults East-West, injection zone ~2 500 m wide
- 2.67 km depth, ~110m thick
- 500 ktons injected from April 2008 to Sept 2009
- Tubåen 1-4 sand layers, separated by shale
- CO₂ migrated into Tubåen 1 sand (80%), and Tubåen 2 and 3 sand layer (20%)
- Will focus on the Tubåen 1 sand layer



Time lapse reflectivity and noise level

- 1D Synthetic seismograms for increasing thickness of time lapse effect
- Time lapse seismic reflectivity noise level approximately 0.007
- CO₂ layer above ~3m detectable



Multiphase CO₂ flow on a short time scale

Nordbotten & Celia (2006)

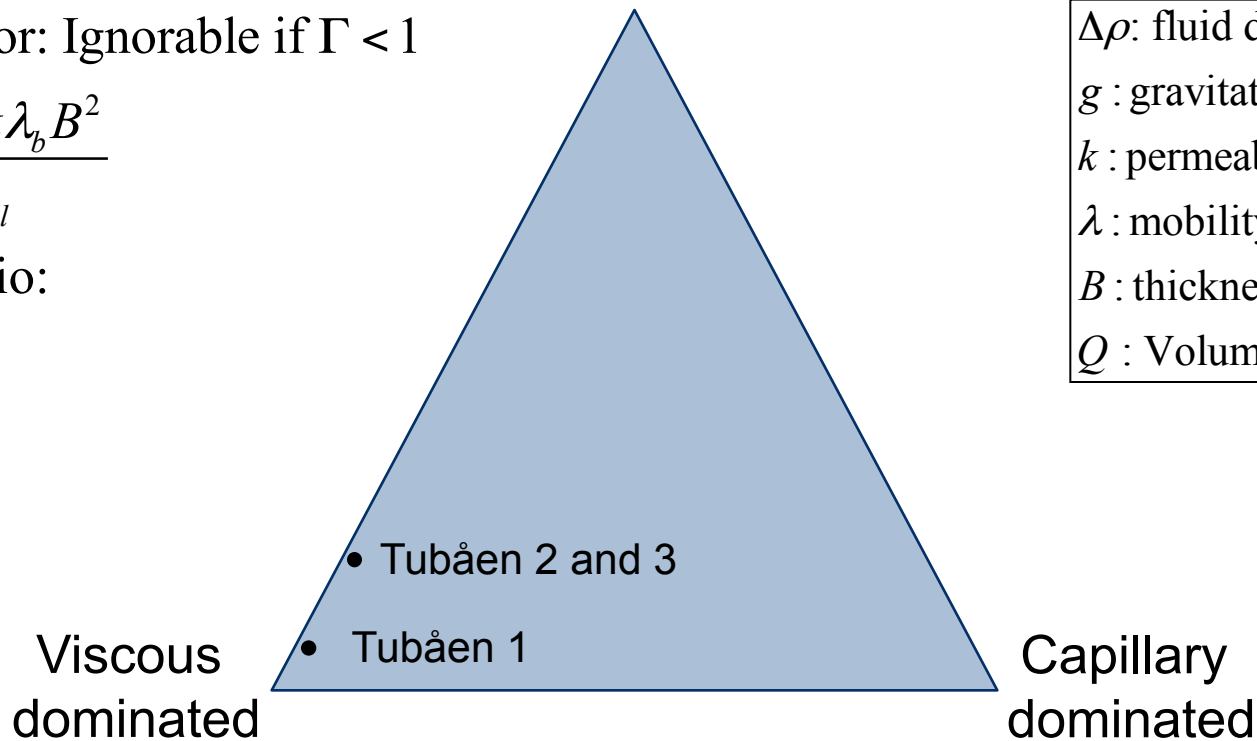
Gravity factor: Ignorable if $\Gamma < 1$

$$\Gamma = \frac{2\pi\Delta\rho g k \lambda_b B^2}{Q_{well}}$$

Mobility ratio:

$$\lambda = \frac{\lambda_{CO_2}}{\lambda_{Brine}}$$

Gravity
dominated

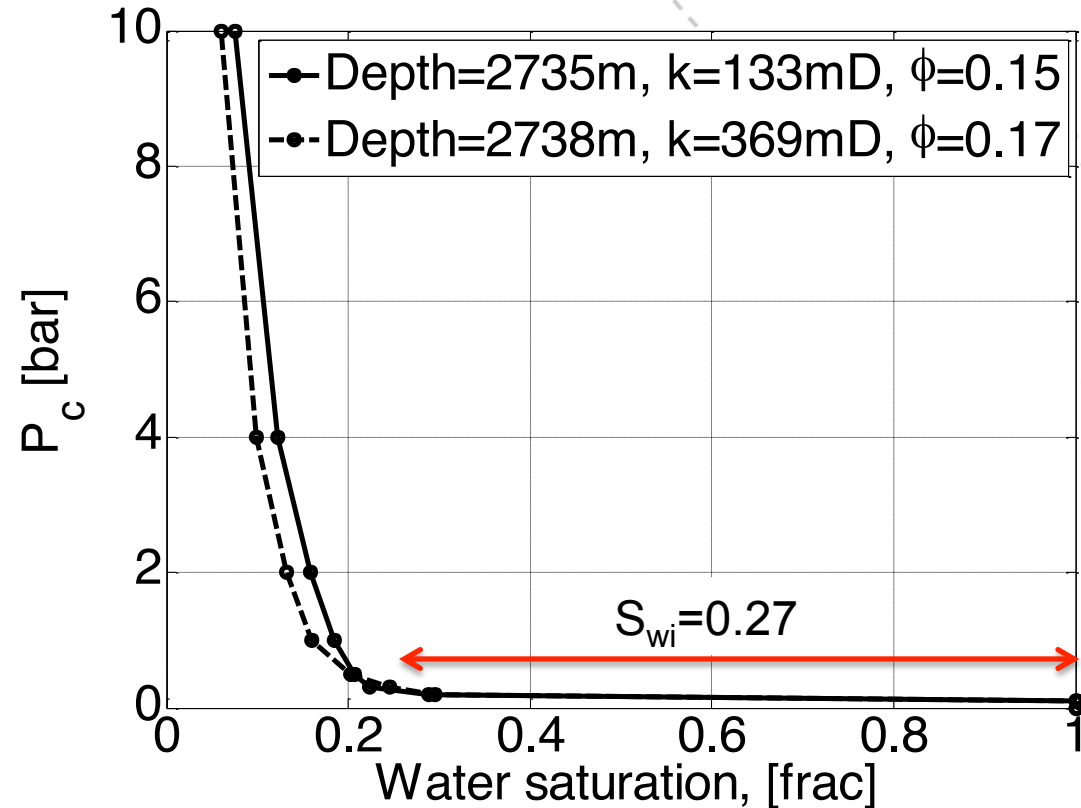


$\Delta\rho$: fluid density difference
 g : gravitational acceleration
 k : permeability
 λ : mobility
 B : thickness of formation
 Q : Volumetric flow rate

Modified from Philip Ringrose's presentation at the CLIMIT & BICCCS PhD Seminar

Capillary pressure

- Laboratory measurements on core plugs from the Tubåen 3 sand layer (courtesy Statoil)
- Low capillary entry pressure
- Low capillary pressure at residual water saturation



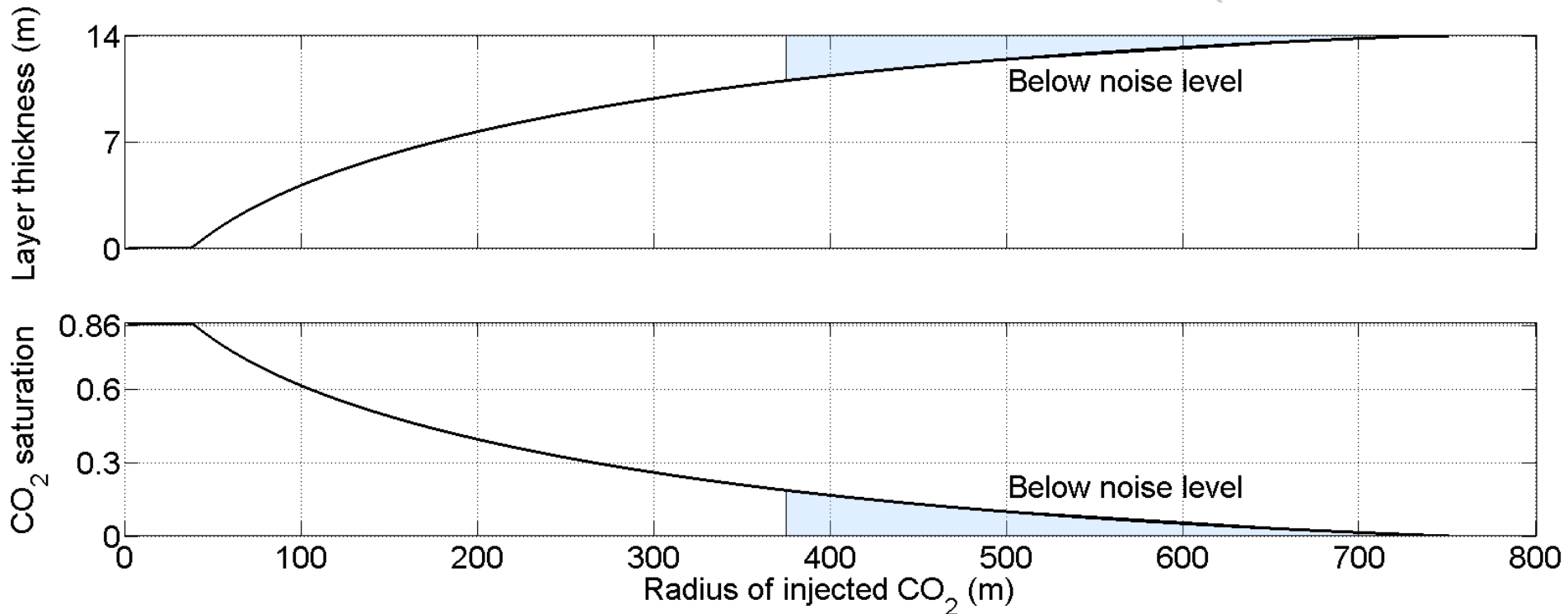
Analytical estimate of the CO₂ plume

		<i>Tubåen 1</i>	
<i>Thickness</i>	h	14	<i>m</i>
<i>Porosity</i>	Φ	0.19	<i>fraction</i>
<i>Permability</i>	k	750	<i>mD</i>
<i>Relative permeability of CO₂</i>	k_{r,CO_2}	0.7 ¹	<i>fraction</i>
<i>Residual water saturation</i>	S_{wi}	0.13 ²	<i>fraction</i>
<i>Injected volume</i>	$Q \times t$	400	<i>ktons</i>
<i>Mobility ratio</i>	λ	6	
<i>Gravity factor</i>	Γ	0.6	

¹ No laboratory measurement available from Tubåen 1. Based on comparison with laboratory measurements of the relative permeability on three composite cores from 2735m measured depth (mD) and published literature ([Bennion and Bachu, 2008](#))

² Estimated from ([Sengupta, 2000](#))

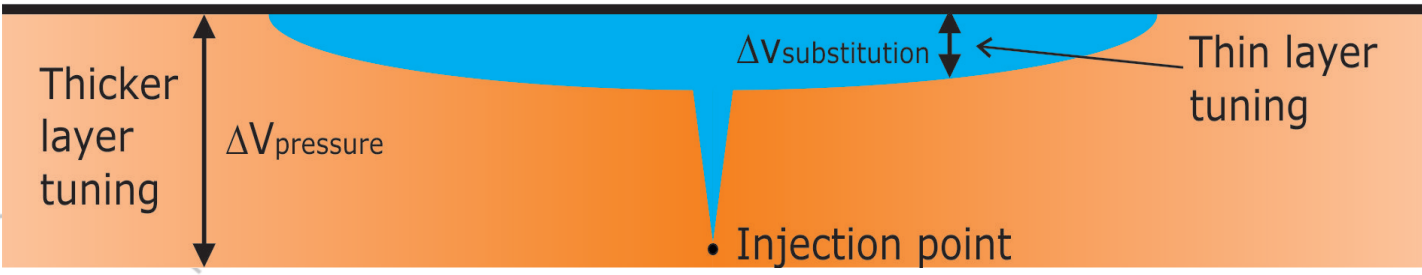
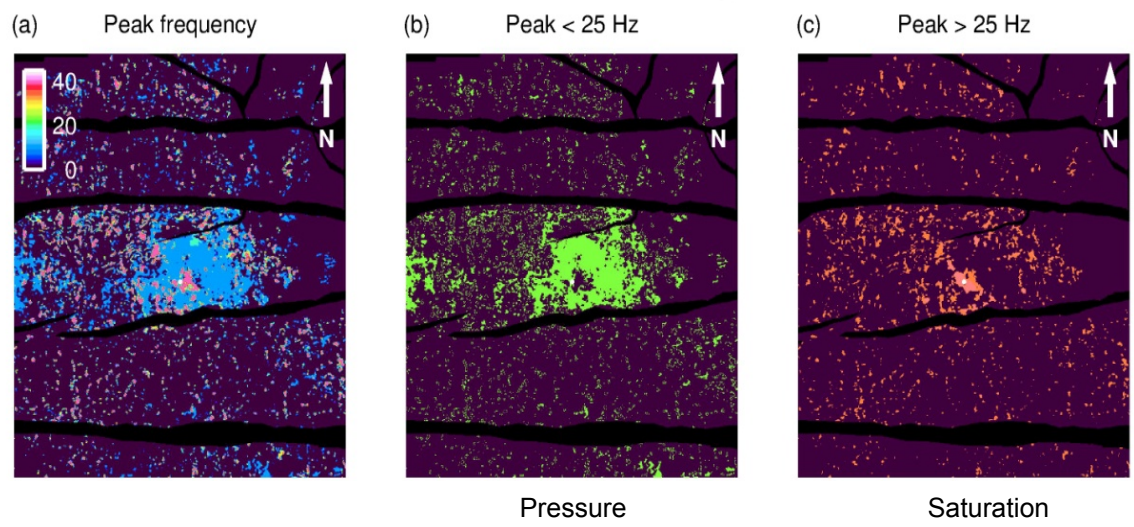
Analytical estimate of the CO₂ plume



Link saturation vs thickness: CO₂ Layer thickness = $16 \cdot S_{CO_2, effective}$

Pressure-saturation discrimination from spectral decomposition

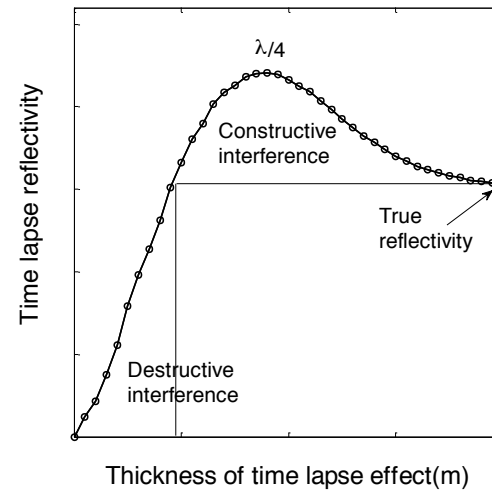
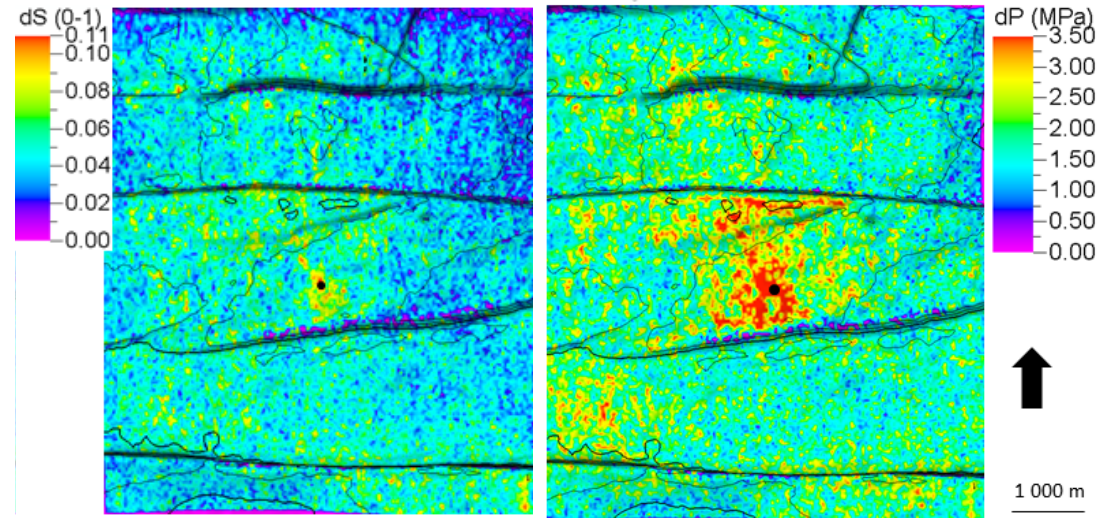
- Assume CO₂ thin layers, pressure more extensively
- Different vertical extents tune different frequencies in seismic wavelet
- 25 Hz cut-off separate the pressure and saturation changes



White et al., 2014

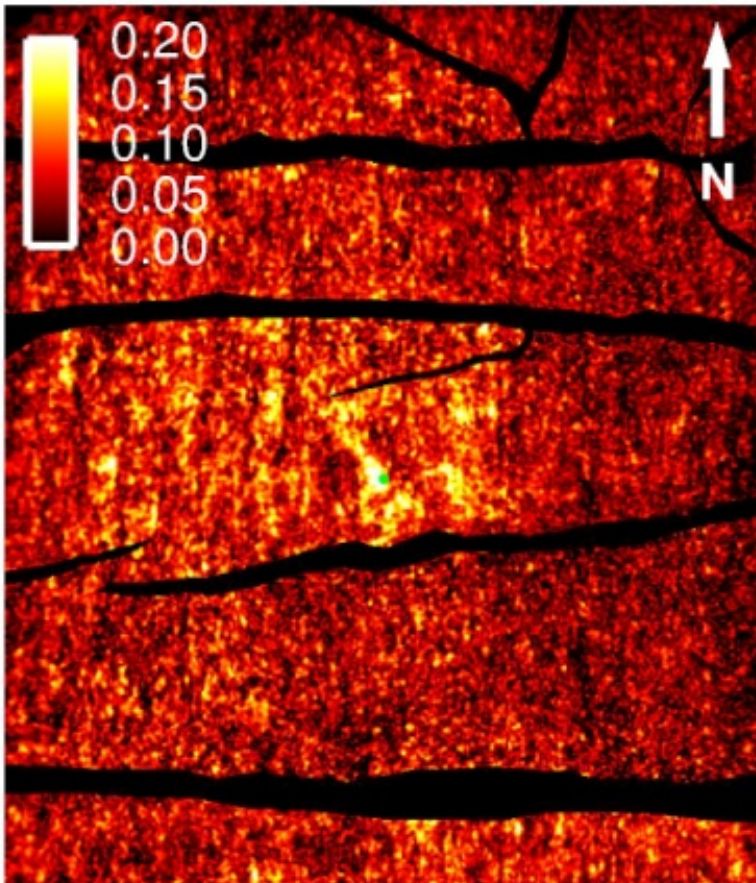
Pressure-saturation discrimination from time lapse AVO seismic

- Introduced by Landrø (2001)
- Exploits time-lapse near- and far offset seismic stacks as independent measurements
- Invert for fluid- and pressure related changes in seismic amplitude
- Thin layer tuning correction of saturation effect

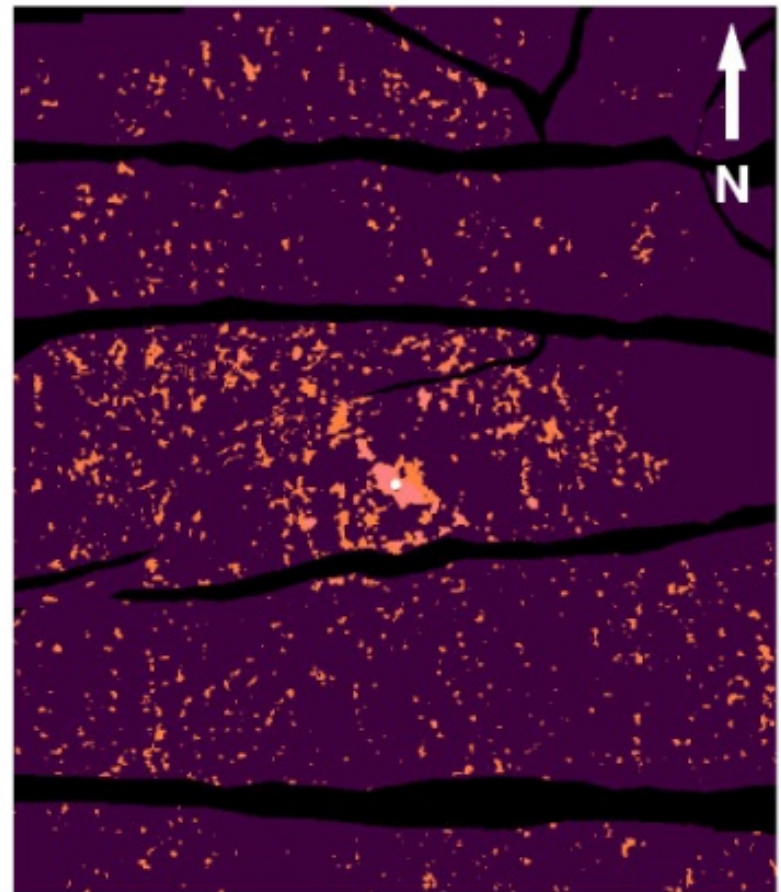


CO₂ saturation from time lapse seismic

AVO

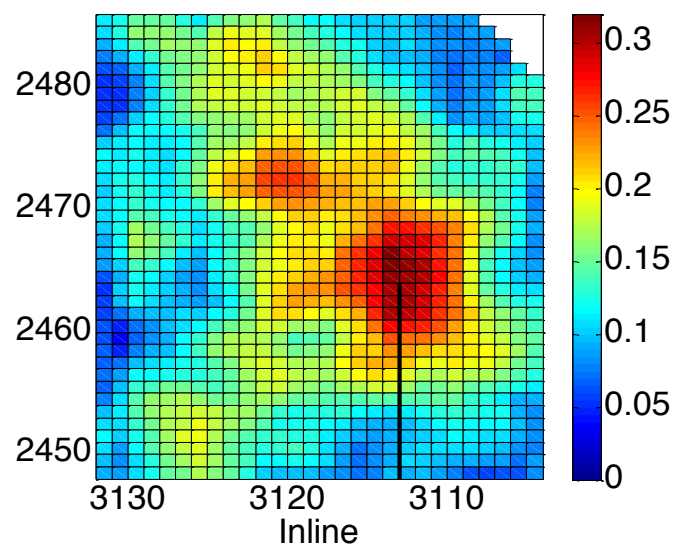


Spectral decomposition

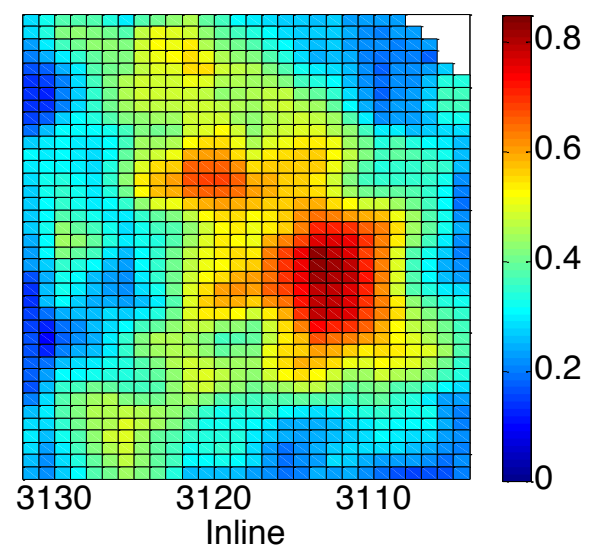


CO₂ saturation inverted from time lapse AVO seismic

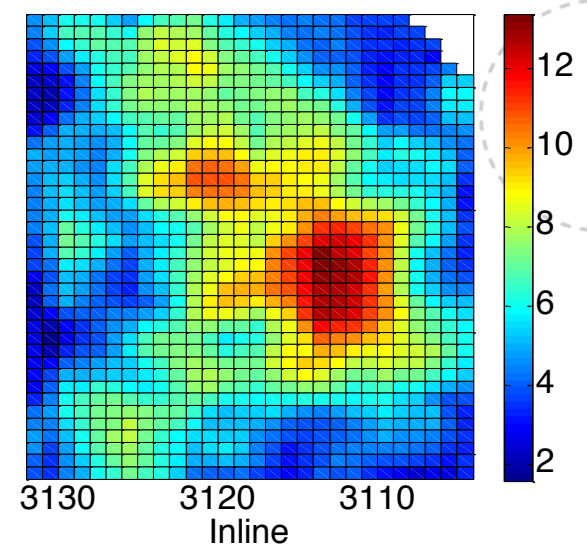
CO₂ saturation



Tuning corrected saturation
(Lin and Phair, 1993)

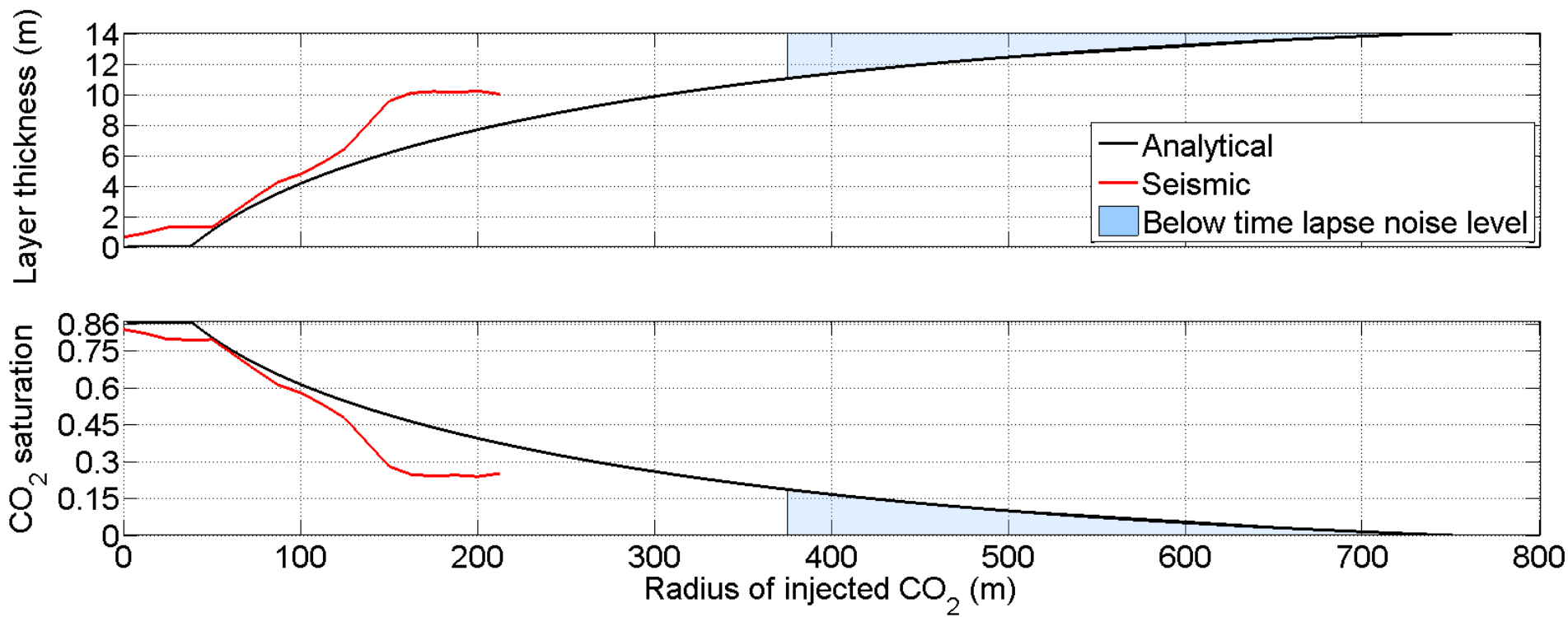


CO₂ plume thickness, (m)



12.5*12.5m grid cells

Comparisons with analytical estimates



11% of injected CO₂ below noise level

Summary CO₂ saturation and radius of plume

	<i>Maximum observable radius, m</i>	<i>Maximum CO₂ saturation, [frac]</i>
<i>Analytical solution</i>	375	0.87
<i>Time lapse AVO</i>	363 ¹	0.83 ²
<i>Time lapse Spectral decomposition</i>	470 ¹	

¹In channel direction (SE-NW)

²After thin layer tuning correction

Discussion

- Early injection times, stratigraphic and structural trapping
- Later: Dissolution of CO₂ in brine, evaporation of brine into the dry CO₂ and mineral trapping dominant trapping mechanisms
- Assume constant injection-rate into confined, homogeneous, isotropic and saline aquifer, wrong for fluvial environment?
- Uncertainty in capillary pressure (core plug measurements representative for large scale?)

Conclusions

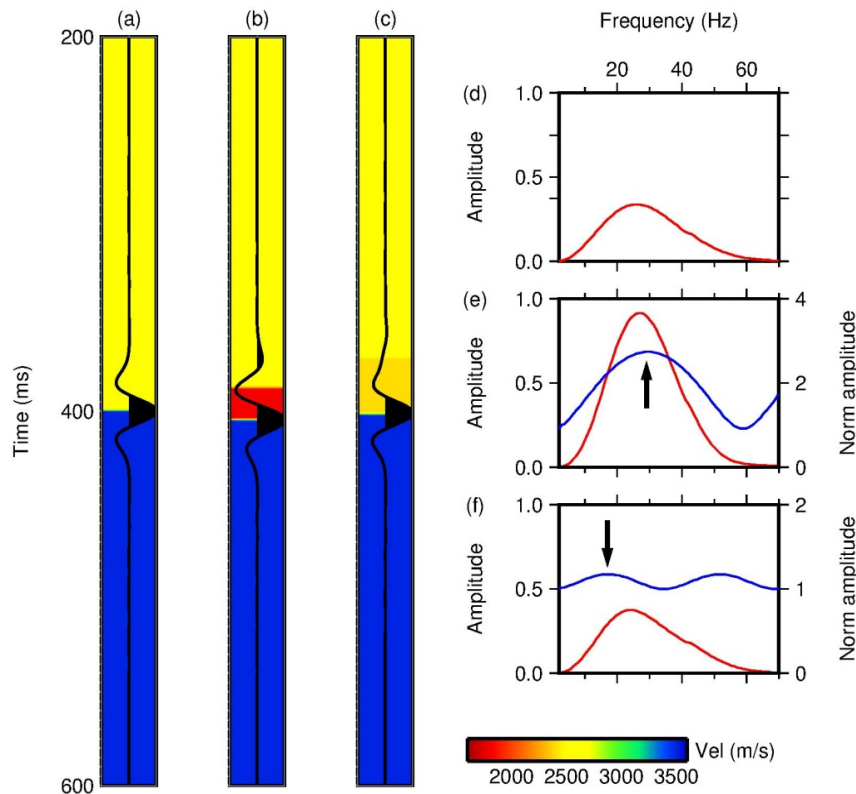
- Viscous dominated fluid flow
- Wedge shaped CO₂ plume
- Thinning of wedge below time lapse noise level (<3m CO₂)
- Spectral decomposition indicate tuning of time lapse saturation effect, and pressure communication with the surroundings
- Plume radius from analytical solution comparable to average radius from time lapse seismic
- CO₂ saturation from time lapse seismic in good agreement with analytical solution after thin layer scaling, with a partially patchy fluid distribution in the pore space

Acknowledgments

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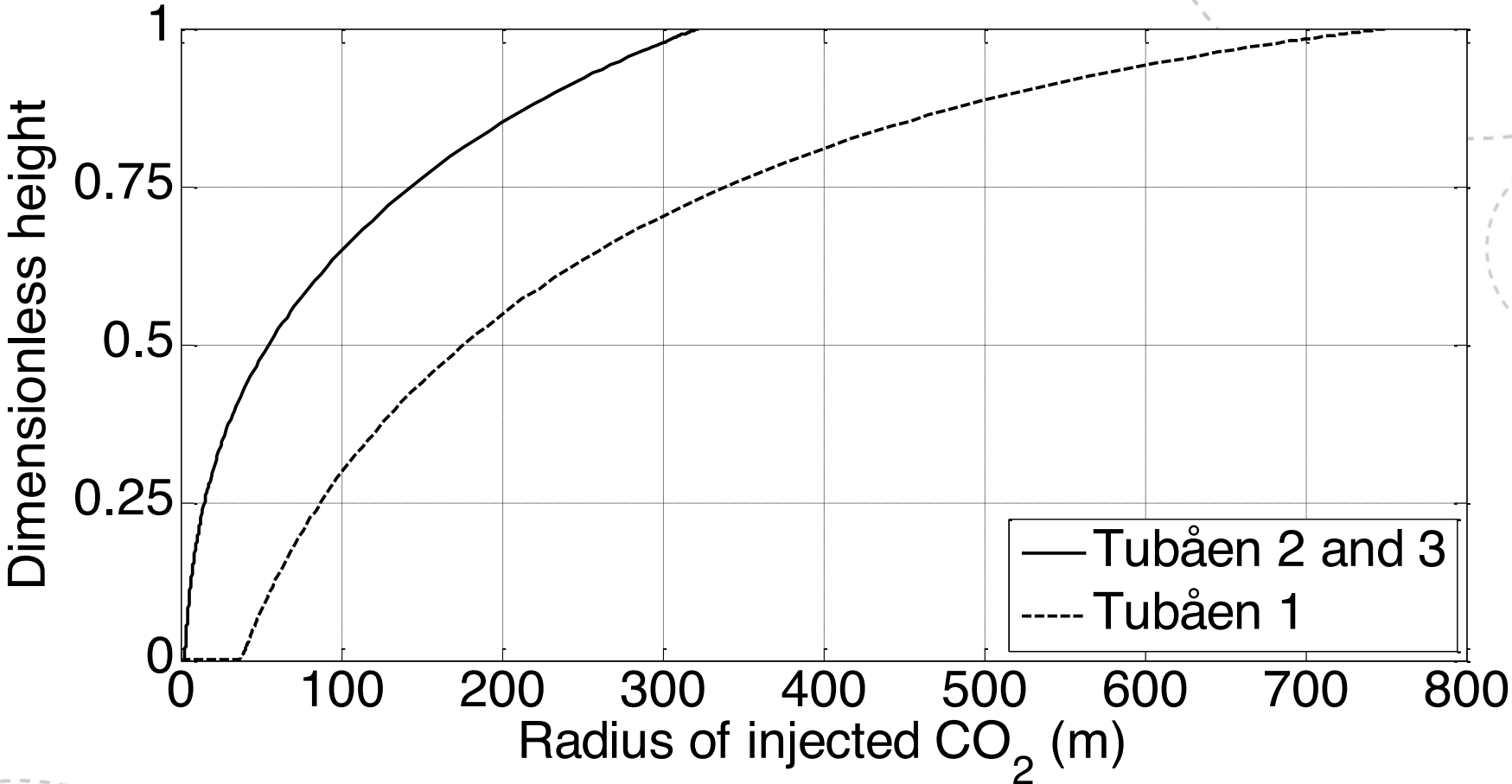
Backup-slides

Spectral decomposition: Pore pressure and fluid substitution changes generated during CO₂ injection



- Baseline wavelet
- 15 m layer where CO₂ has replaced brine
- 36 m thick pore pressure anomaly
- Baseline spectral response
- Red line- spectral content of the reflected energy following fluid substitution changes. Blue line- the same spectra normalised by the baseline response. The peaks define the tuning frequency for the top and base reflection. Peak frequency of ~29 Hz. Corresponds to a temporal thickness of 17 ms.
- Red line- spectral content of the reflected energy following pore pressure changes. Blue line- the same spectra normalised by the baseline response. The peaks define the tuning frequency for the top and base reflection. Peak frequency ~17 Hz, equivalent to temporal thickness of 30 ms. The second peak frequency at 51 Hz corresponds to the 2nd tuning harmonic.

Analytical estimate of the CO₂ plume



Comparisons with analytical estimates

Opposite direction in blue

