

# Thermally induced compaction of shales

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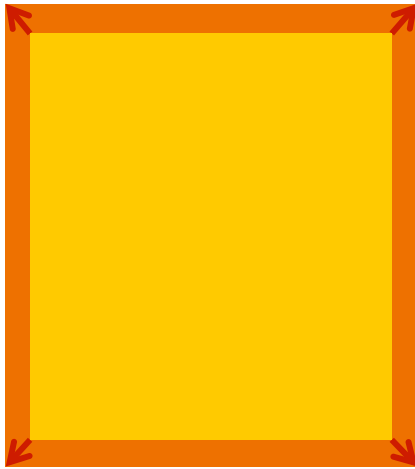
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## What we learn in physics:

Solids expand upon heating and retract upon cooling

Heat expansion



- Heat expansion and thermal stresses have a large impact on borehole stability and caprock integrity

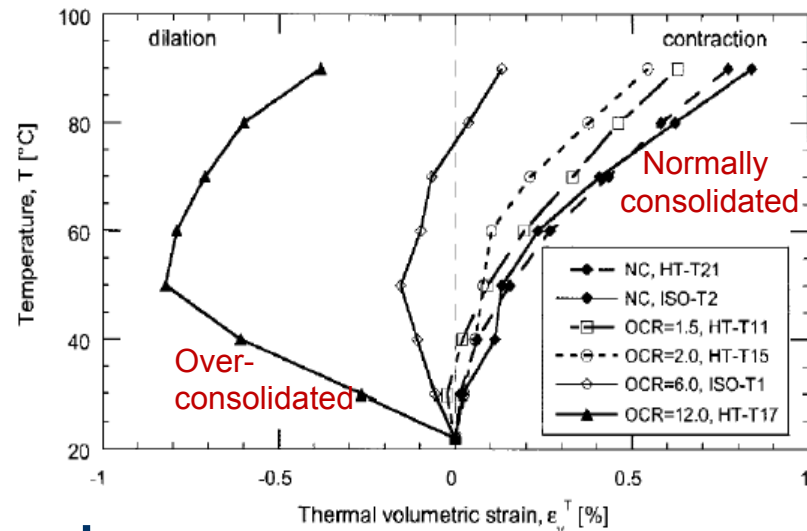
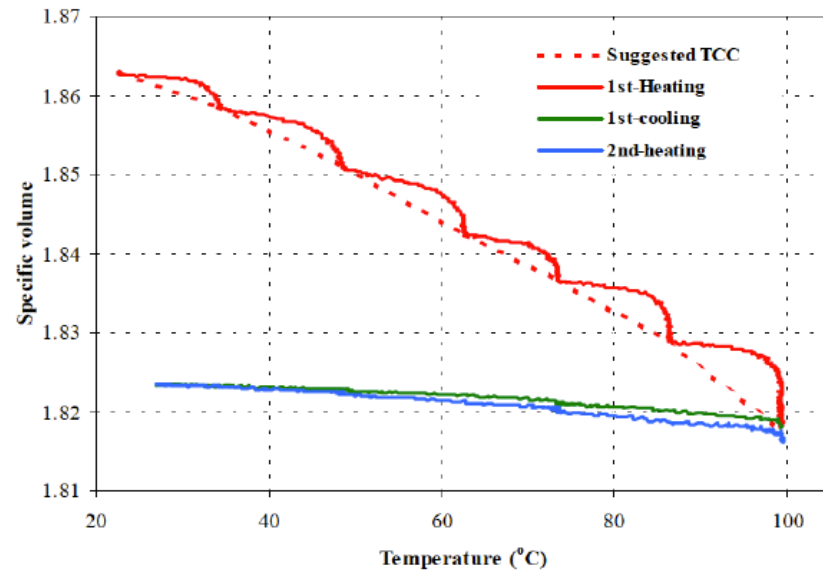
## However ...

- Heating may result in phase transitions, or non-reversible processes that involve an activation energy (chemical reactions, mechanical processes on the atomic scale)
- Rock compaction is visco-plastic (creep behavior), and creep processes may be accelerated at increased temperature

# Thermally induced compaction in clay

□ Thermal compaction of reconstituted clay (Ghahremannejad, 2003): Largely plastic behaviour during initial heating, elastic behaviour during cooling and reheating

□ Stress history dependent thermal behaviour of kaolin clay during heating; initial consolidation at 0.6 MPa (Cekerevac and Laloui, 2004)

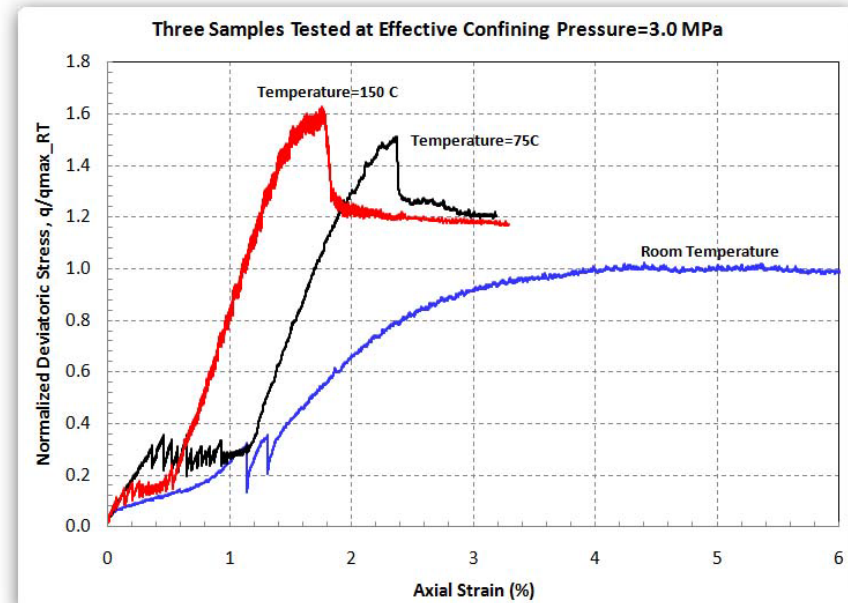
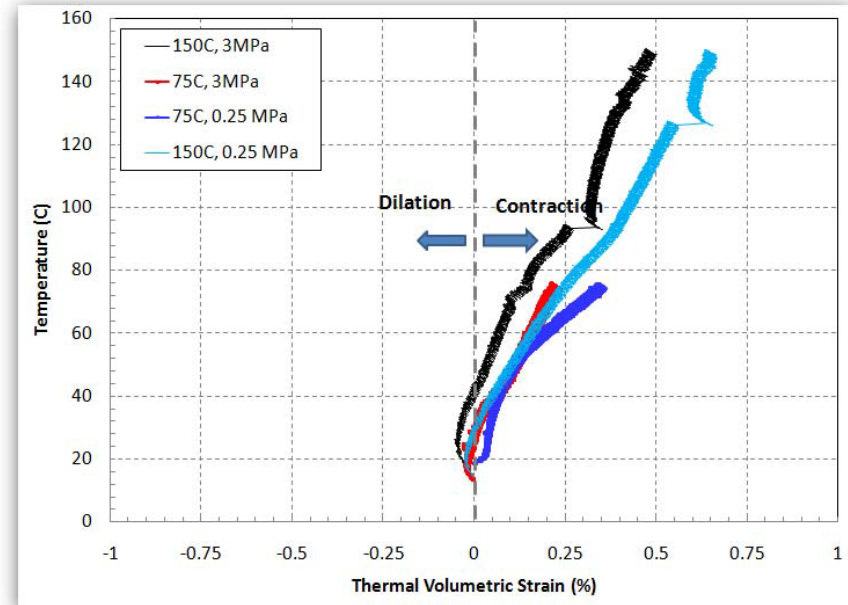


↪ Degree of compaction depends on overconsolidation ratio

# Thermally induced compaction in caprock shale

B. Xu, Y.G. Yuan, and Z.C. Wang (2011)

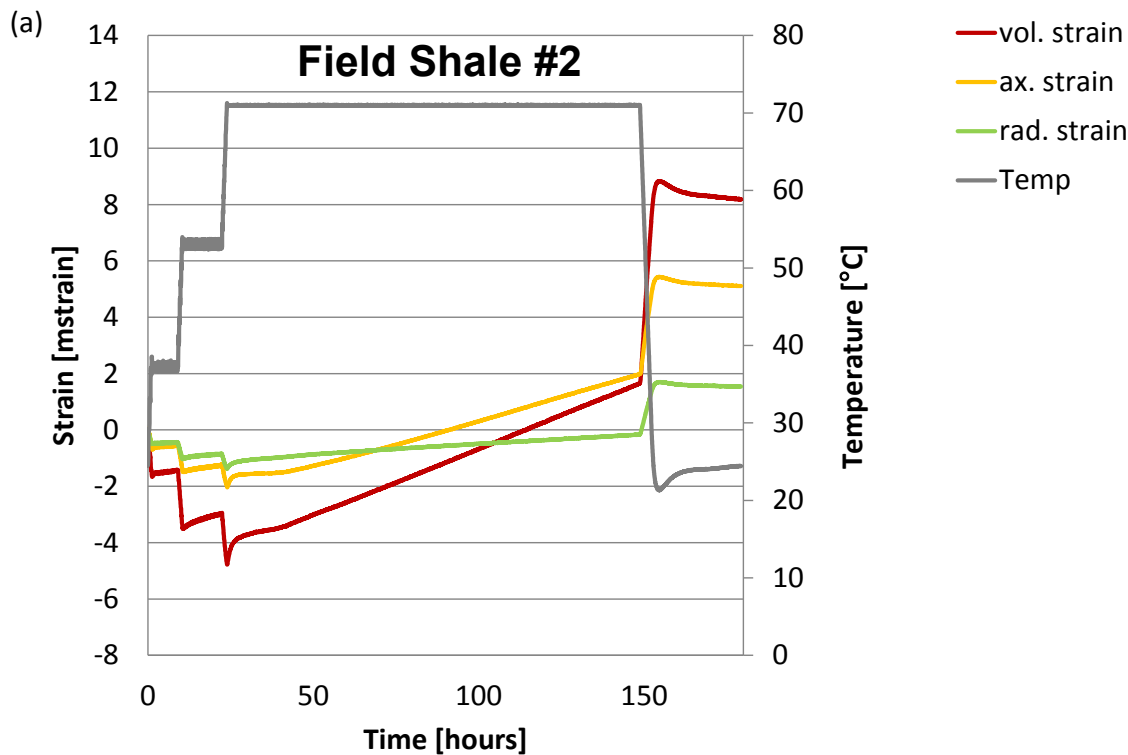
- ❑ Thermal contraction of water-saturated Clearwater Shale (10–30% clay content) during drained heating to 75°C and 150°C.
- ❑ Sample stiffness and strength increase (for slow heating rates, allowing for pore pressure equilibration)



At SINTEF we perform extensive rock-mechanical testing of field shales within the Shale Rock Physics JIP and the Dynamic Borehole Stability JIP

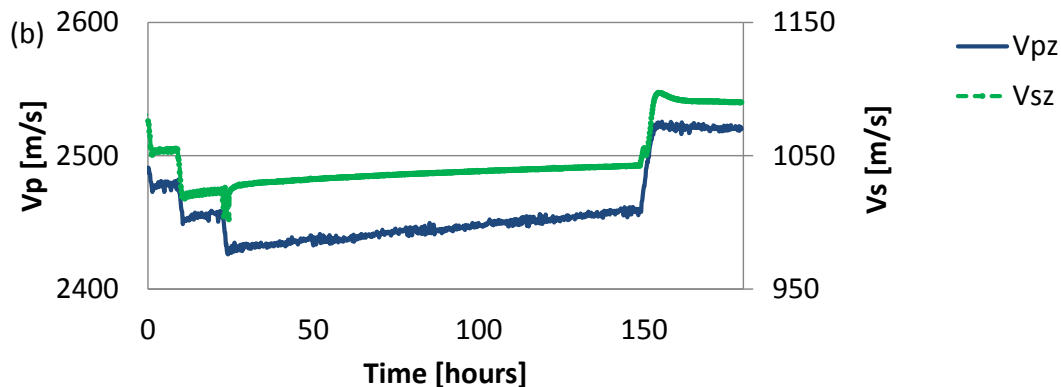
**For competent, (over-)consolidated caprock shales, we expected negligible or small thermally-induced compaction for temperatures around in-situ temperature**

# Test results reveal large compaction upon heating at temperature at or below in-situ temperature



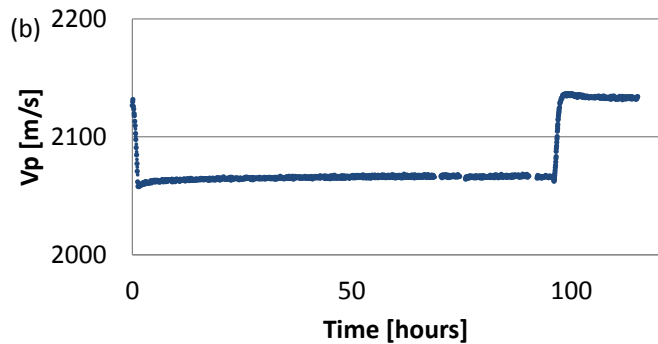
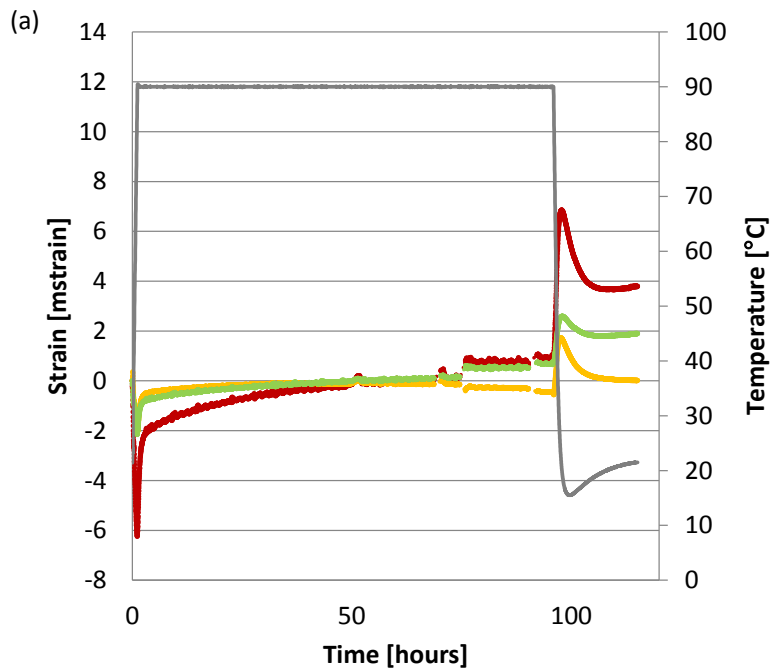
## Experimental procedure:

- Shale core plugs fully saturated and brought to in-situ stress condition
- Stepwise heating
- Strains and ultrasonic velocities recorded as a function of time
- Cooling to room temperature

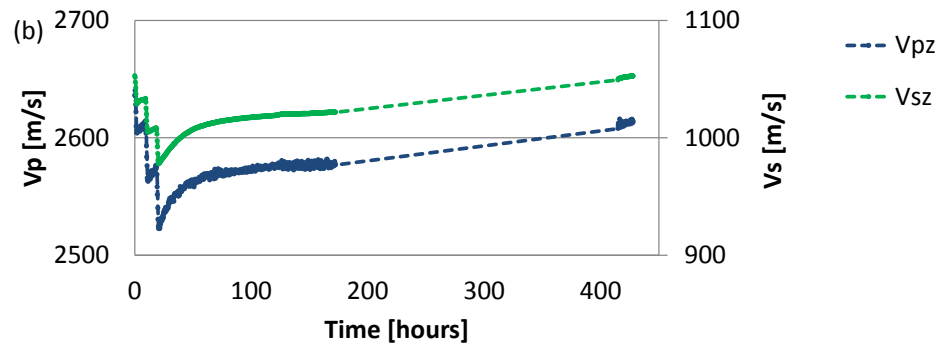
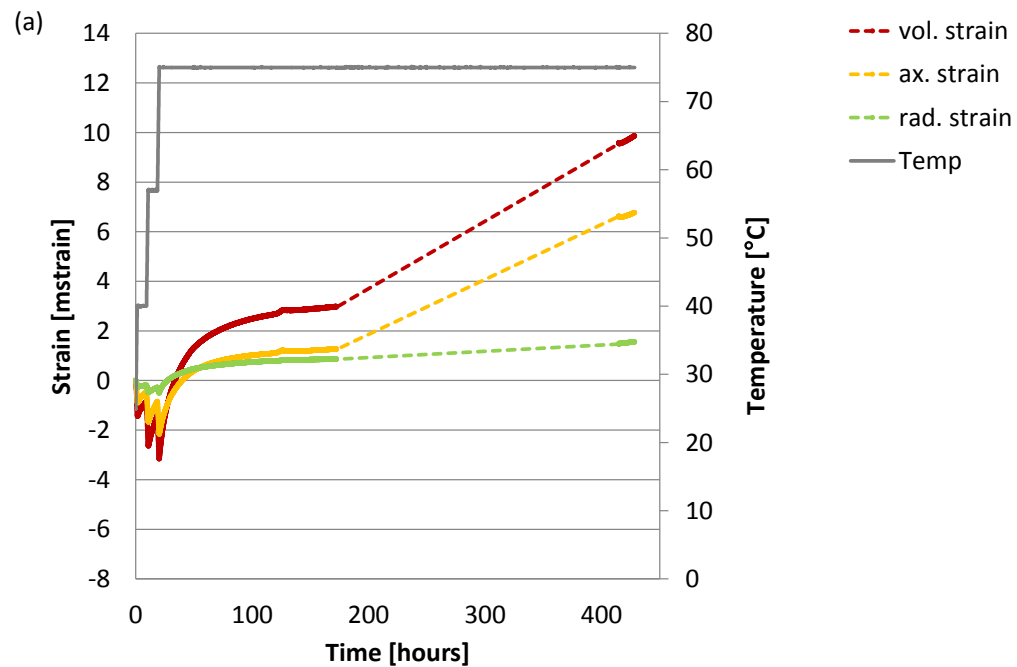


- Heating/cooling cycle results in 0.8% volumetric compaction
- Creep behavior (primary, secondary, and tertiary creep)
- Compaction is also reflected by increase in ultrasonic velocities  $v_p$  and  $v_s$

## Field Shale #1

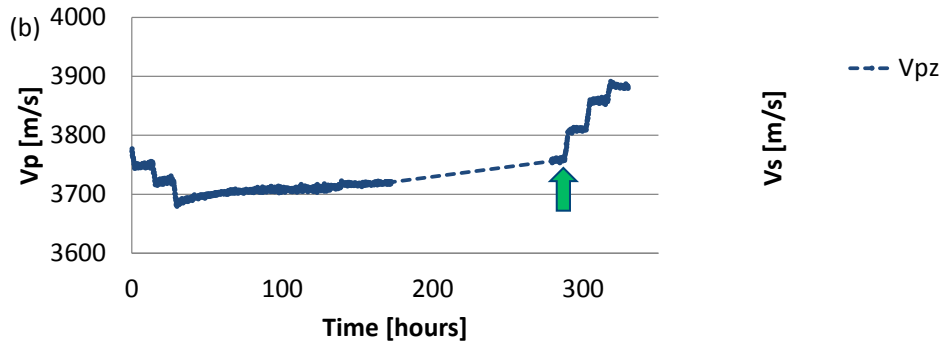
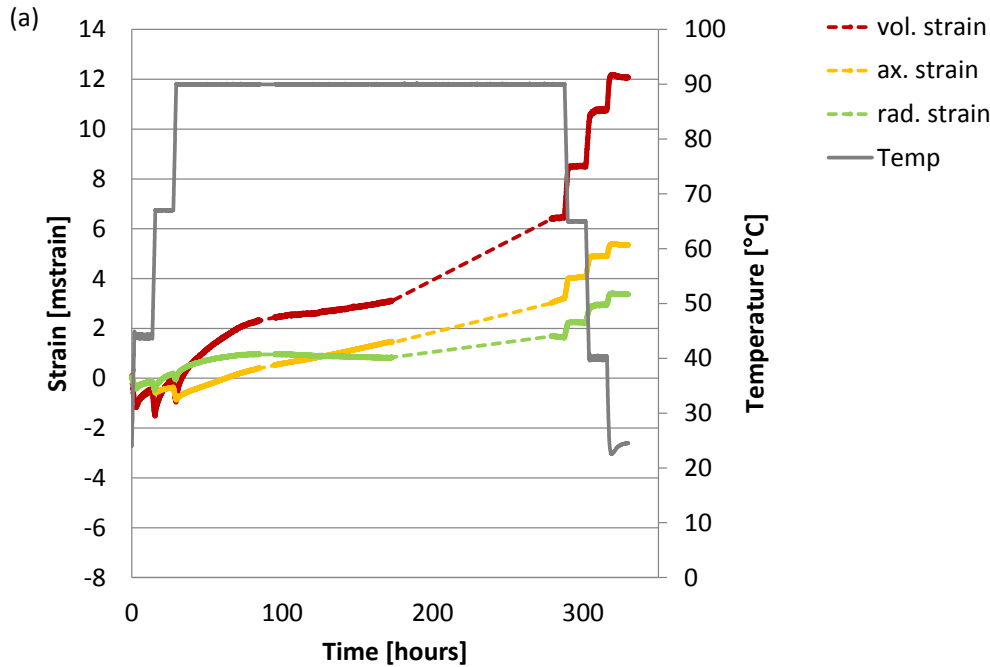


## Field Shale #3





# Field Shale #4



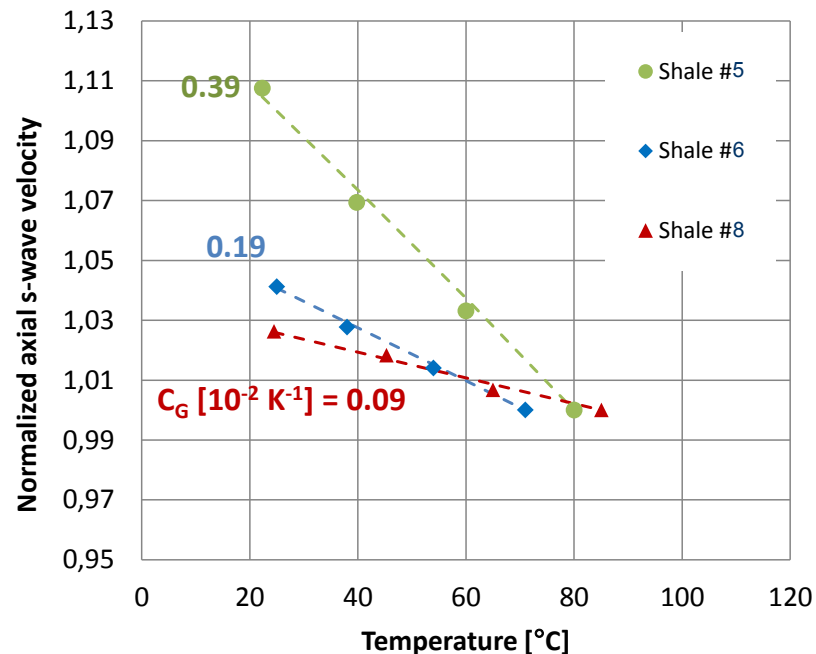
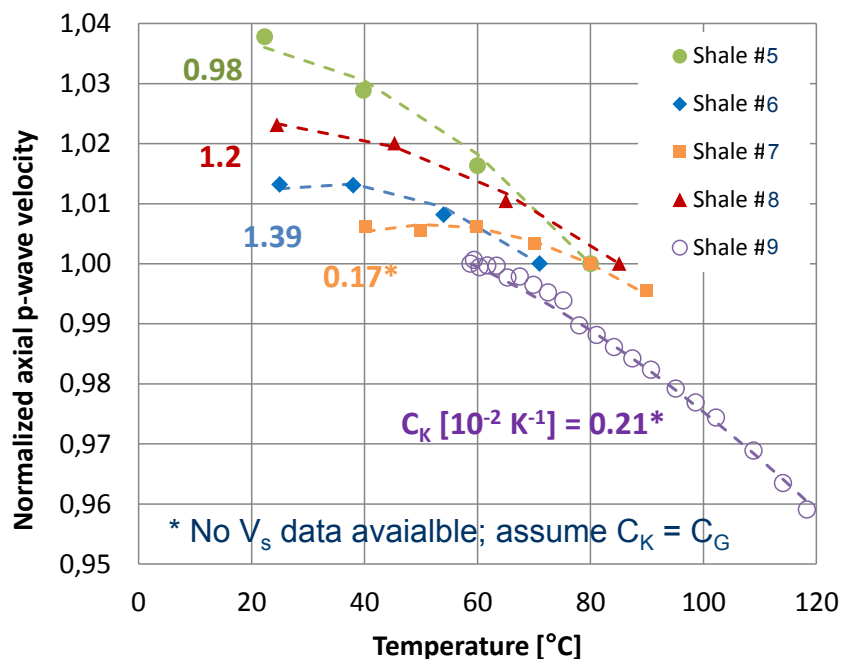
## General features are the same for all shales:

- Heat expansion is followed by compaction (primary creep), with the compaction rate increasing with temperature (compaction may partly be explained by drainage effects)
- Reduction of compaction rate until a constant creep rate is established (secondary creep)
- Acceleration of compaction rate after a certain time (tertiary creep)
- Asymmetry between axial and radial strain
- Velocity increase (↑) upon cooling is reversible; reason for relatively large, reversible velocity changes (both  $v_p$  and  $v_s$ ) is not understood

# Reversible temperature dependence of ultrasonic velocities of Shales

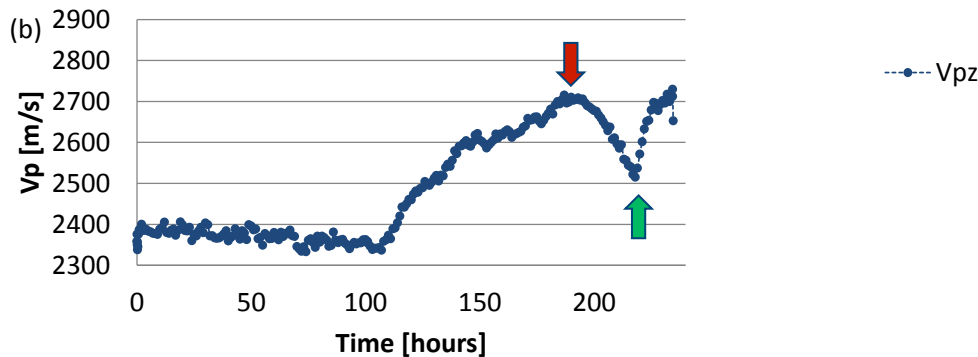
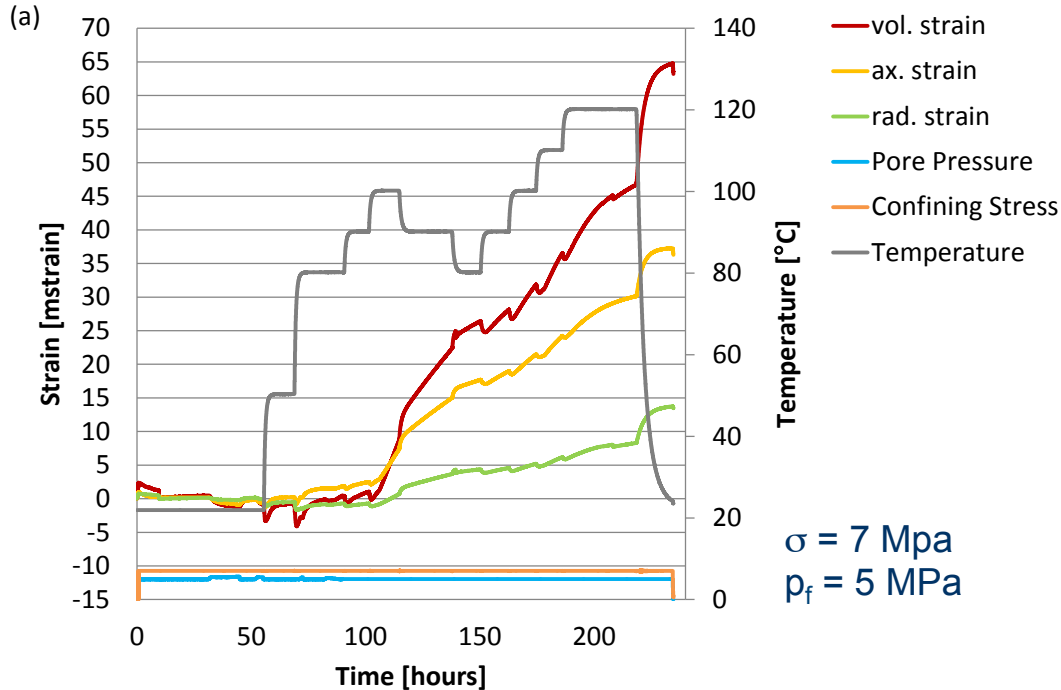
- Apply Gassmann model
- Assume linear temperature dependence of dry-rock stiffness:

$$K_{dry} = K_{dry,0} (1 - C_K \cdot \Delta T) \quad G_{dry} = G_{dry,0} (1 - C_G \cdot \Delta T)$$



Good fit can be obtained by selecting the right temperature-sensitivity coefficients; physics not understood yet

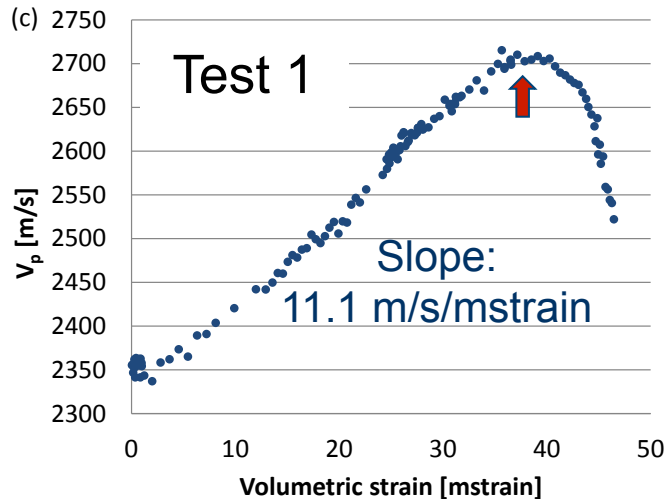
# Thermally-induced compaction in Pierre shale (outcrop)



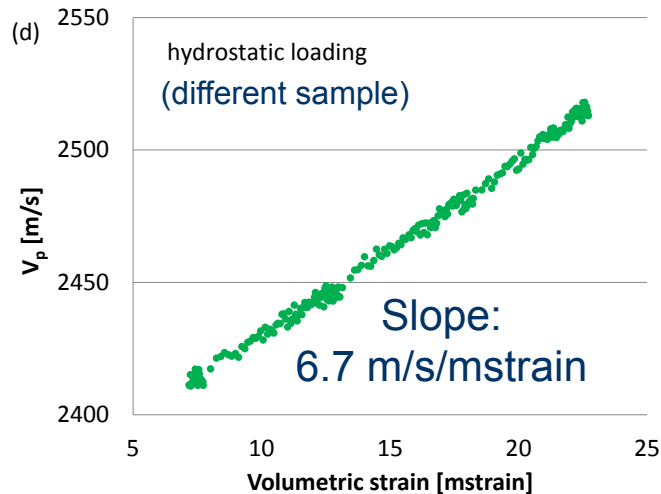
- Volumetric compaction of 6.5% induced by heating to 120°C
- Strong increase in compaction for  $T \geq 80^\circ\text{C}$
- Temperature dependence of creep rates is consistent with Arrhenius law for thermally activated processes ( $\propto e^{-E_a/k_B T}$ )
- Strong increase in  $v_p$
- velocity decrease (↓) during consolidation phase at 120°C possibly caused by rock failure at large strains
- Velocity increase (↑) upon cooling

# Thermally-induced compaction in Pierre shale (outcrop)

## Heat-induced compaction

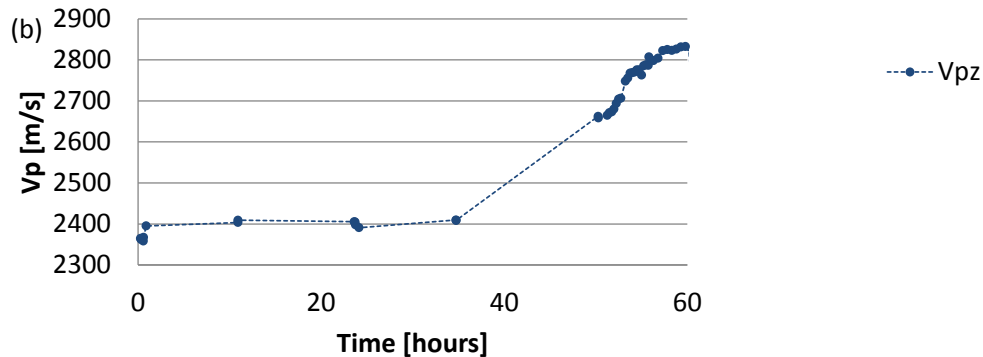
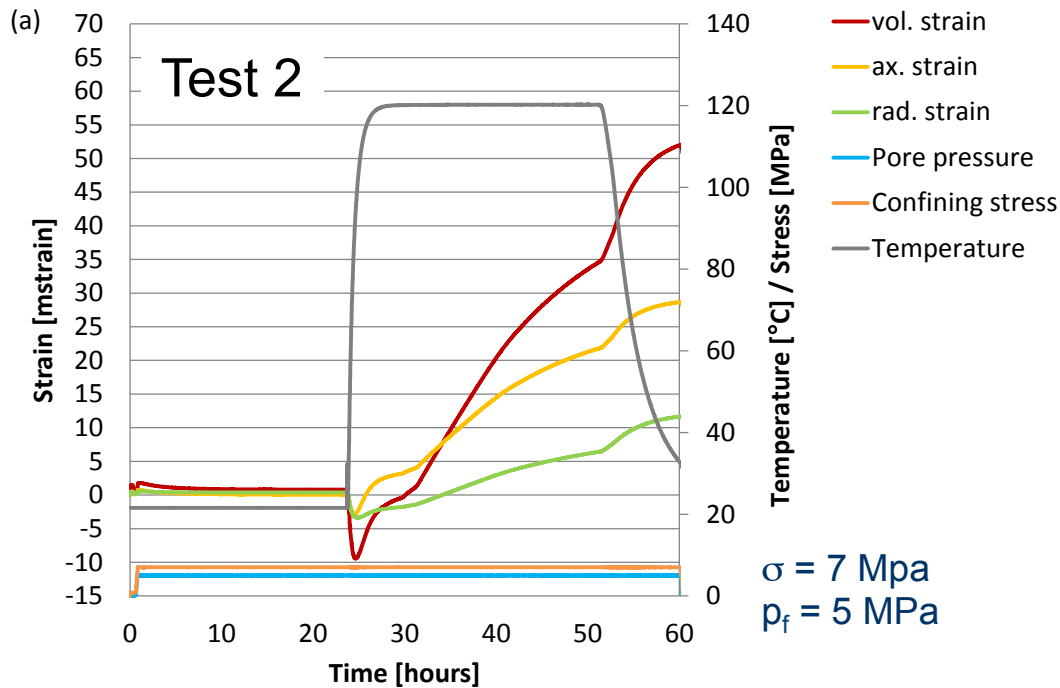


## Stress-induced compaction



- Nearly linear dependence of  $v_p$  on vol. strain
- Volumetric-strain dependence comparable to that measured during hydrostatic loading, suggesting that velocity changes are mainly controlled by volumetric strain (porosity changes).
- Field shales: volumetric-strain dependence of velocities during heat-induced compaction is by a factor 3 – 9 smaller than during hydrostatic loading, suggesting that heat-induced compaction has a different impact on the microscopic rock structure than stress increments

# Thermally-induced compaction of Pierre shale (outcrop)



- Heating to 120°C in one step
- Observation of characteristic creep behavior at constant stress and temperature
- Irreversible volumetric compaction of 5.1%, which corresponds to a porosity reduction from 19.2% to 14.1%
- Irreversible increase of  $v_p$  by 423 m/s (17.6%)

# Summary of test results

	porosity [%]	Clay [wt%]	Smectite [wt%]	$\sigma'_{conf}$ [MPa]	T [°C]	T <sub>in-situ</sub> [°C]	$\Delta\varepsilon_{vol}$ [%]	$\Delta V_p$ [m/s]	$\Delta V_s$ [m/s]	
Pierre shale (outcrop)	test 1	19,2	57,4	31,5	2	120	5,1	423	-	
	test 2	19,2	57,4	31,5	2	120	6,5	330	-	
Field shales	#1	34,2	64	0	4,9	90	74	0,38	6	-
	#2	20	56	29	8,2	71	50	0,8	32	10
	#3	18,6	65	53	10,5	75	75	≈ 1,34	≈ 49	≈ 42
	#4	5	69	0	19,1	90	103	1,2	110	-

- Very large thermally-induced compaction ( $\varepsilon_{vol} > 5\%$ ) for outcrop shale at 120°C
- Large thermally-induced compaction ( $\varepsilon_{vol} > 1\%$ ) for field shales #3 and #4 at temperatures at or below in-situ temperature
- No apparent correlation with initial porosity, smectite content, or applied stress
- Large compaction observed even in a highly compacted, low-porosity shale (shale #4)

# Possible reasons for thermally-induced compaction in shale core plugs at or below in-situ temperature

1. **Core damage:** (micro-)cracks formed during core retrieval might not get closed again upon re-establishing the in-situ stress state. Heat may enhance crack healing.

Relevance for field applications: Damage zone around borehole

2. **Heat-rate effects:** For high heat rates, pore pressure builds up in the rock as a result of fluid expansion and slow drainage, which can result in shear or tensile rock failure. Typical pore-pressure changes:  $\approx 0.2 \text{ MPa}/^\circ\text{C}$ .

Relevance for field applications: In shales, heat diffusion is generally faster than pore-pressure diffusion; rock failure as a result of pore-pressure increase is a likely to happen for large enough heating

# Possible reasons for thermally-induced compaction in shale core plugs at or below in-situ temperature

3. **Brine chemistry:** Samples are exposed to synthetic brine, which might result in rock weakening for not perfectly matched brines

Relevance for field applications: Drilling fluids may alter viscoplastic properties of surrounding shale formations



# Summary

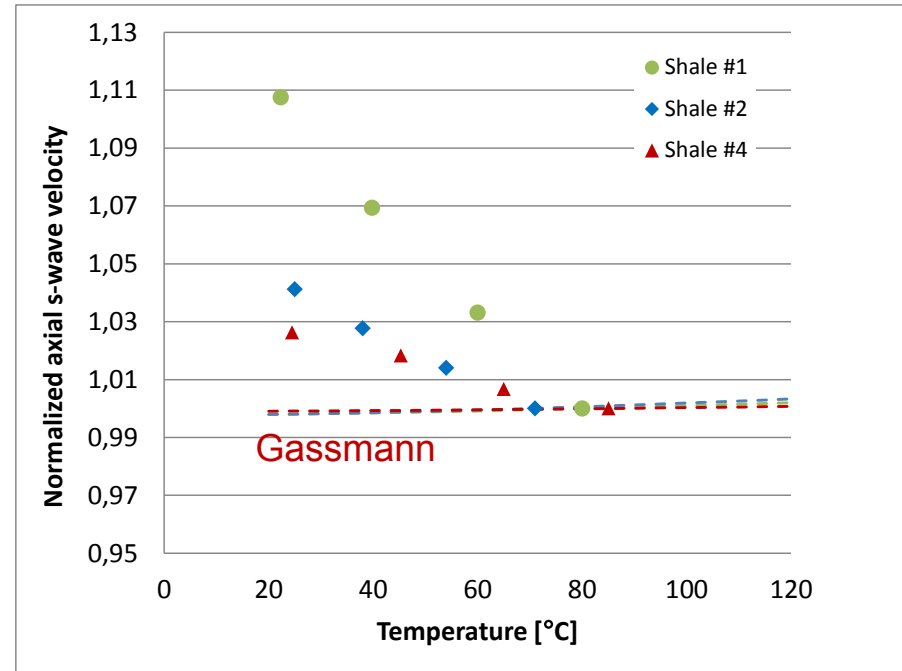
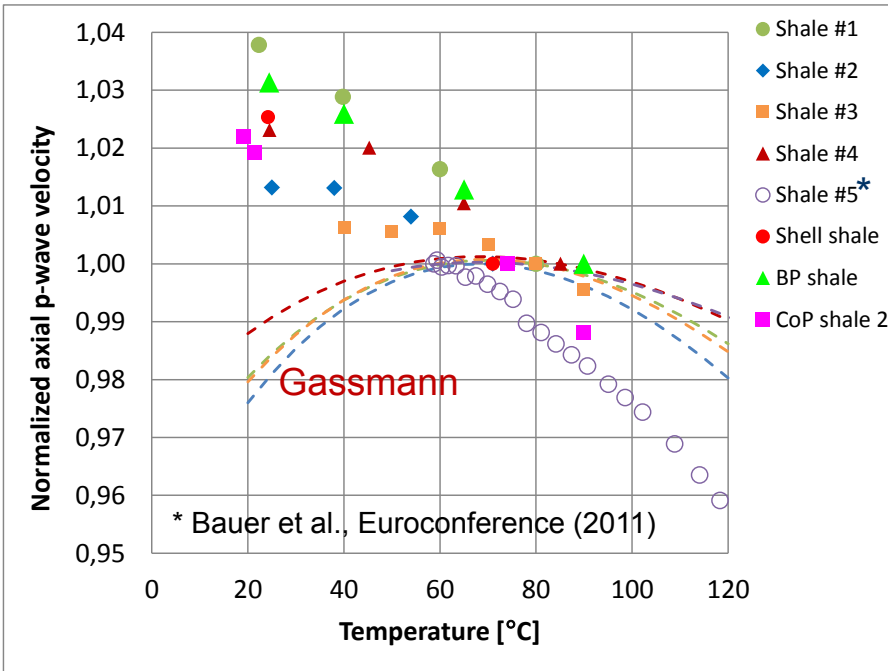
- ❑ Laboratory tests with core plugs of 5 different shales exhibit sizable compaction (up to  $> 1\%$  volumetric strain) upon heating accompanied by increases in ultrasonic velocities
- ❑ Temperature are not high enough to cause chemical changes (e.g. smectite-illite transitions)
- ❑ Possible explanations for observed compaction include core damage, heat-rate effects, and brine-chemistry effects
- ❑ Thermally-induced compaction of shales may have strong impact on field applications such as steam injection


# Backup slides



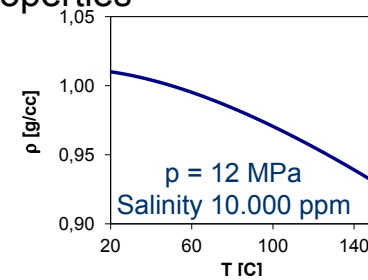
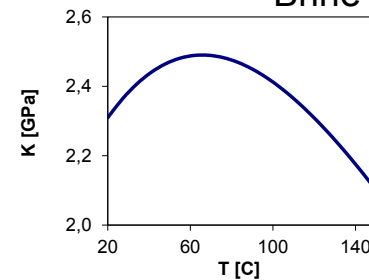
# Temperature dependence of ultrasonic velocities of Shales

- Apply Gassmann model
- Assume temperature-independent dry-rock modulus



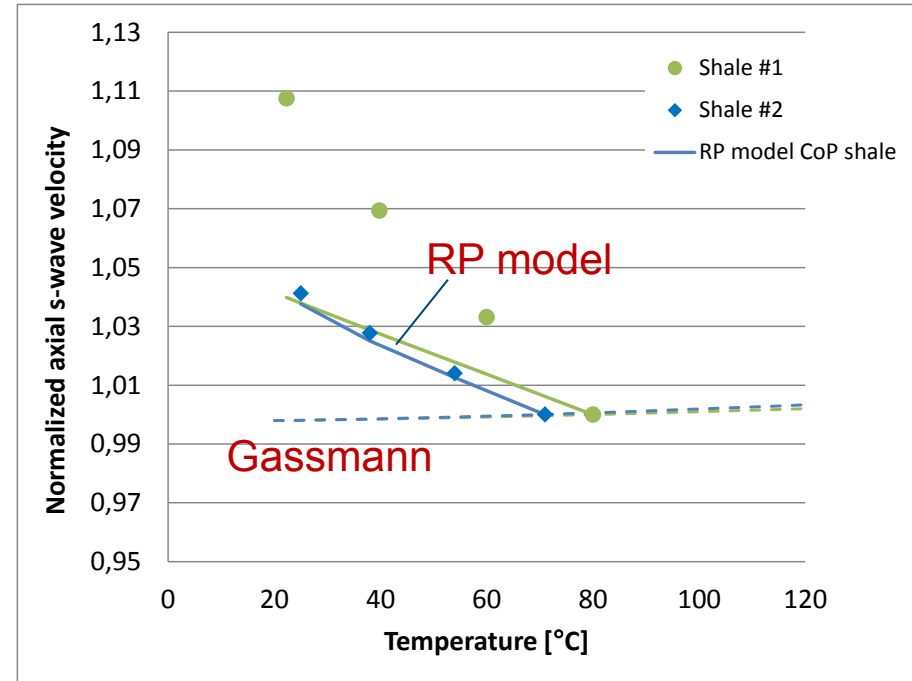
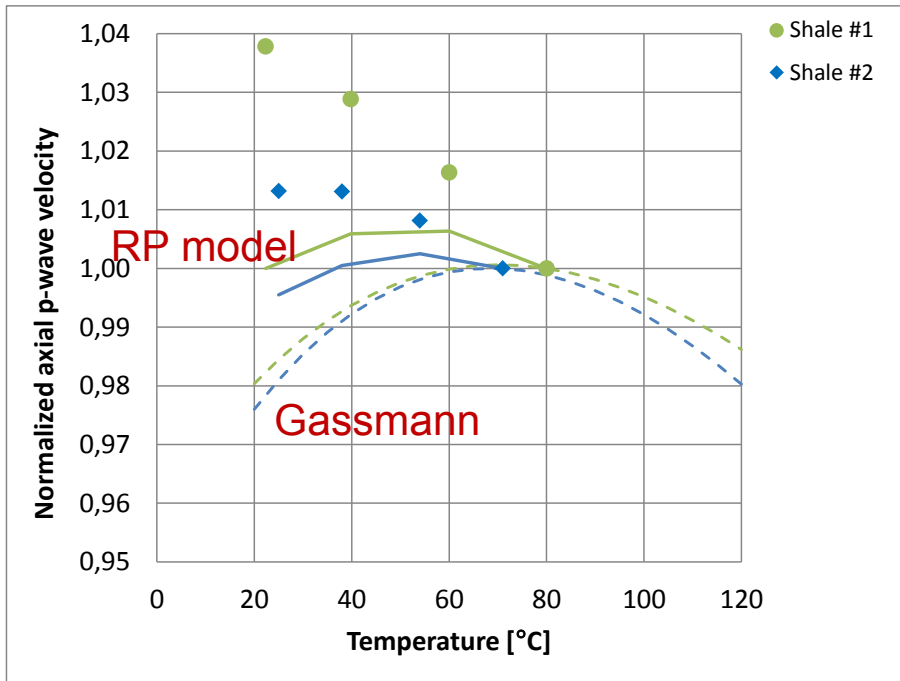
 Gassmann model predicts much weaker temperature dependence and wrong trend at low temperatures


## Brine properties



# Temperature dependence of ultrasonic velocities of Shales

- Apply SINTEF's RP model
- Assume that temperature dependence of bulk and shear stiffness of bound water is the same as that of ice

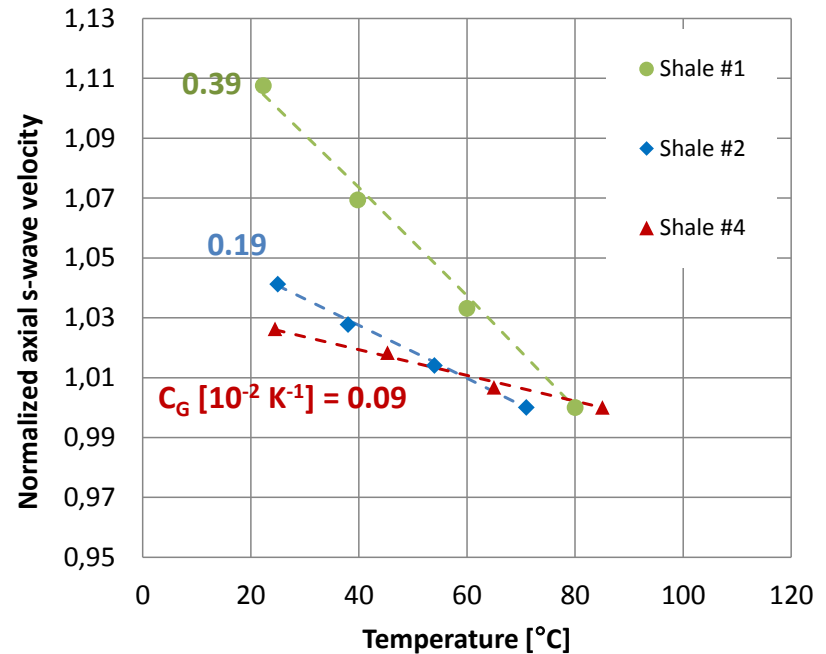
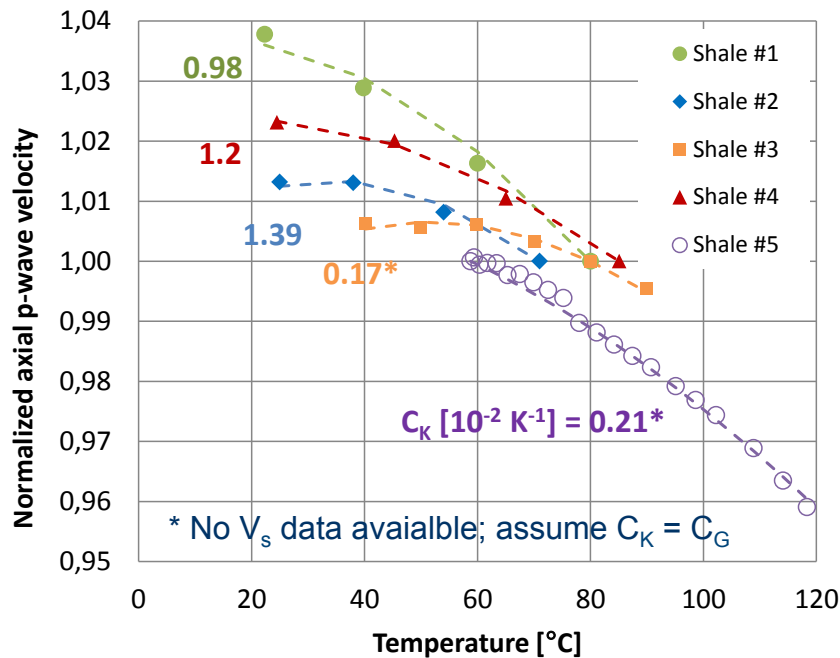


 The Shale RP model provides a better description than the Gassmann model (for temperature independent dry-rock stiffnesses) but still underpredicts the temperature dependence of  $V_p$  and  $V_s$  (except for shale #2 where a good match of the  $V_s$  data is obtained).

# Temperature dependence of ultrasonic velocities of Shales

- Apply Gassmann model
- Assume linear temperature dependence of dry-rock stiffness:

$$K \downarrow_{dry} = K \downarrow_{dry,0} (1 - C \downarrow K \cdot \Delta T) \quad G \downarrow_{dry} = G \downarrow_{dry,0} (1 - C \downarrow G \cdot \Delta T)$$



Good fit can be obtained by selecting the right temperature-sensitivity coefficients; physics not understood yet