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Seismic waveform tomography

Centre for Reservoir Geophysics

Yanghua Wang

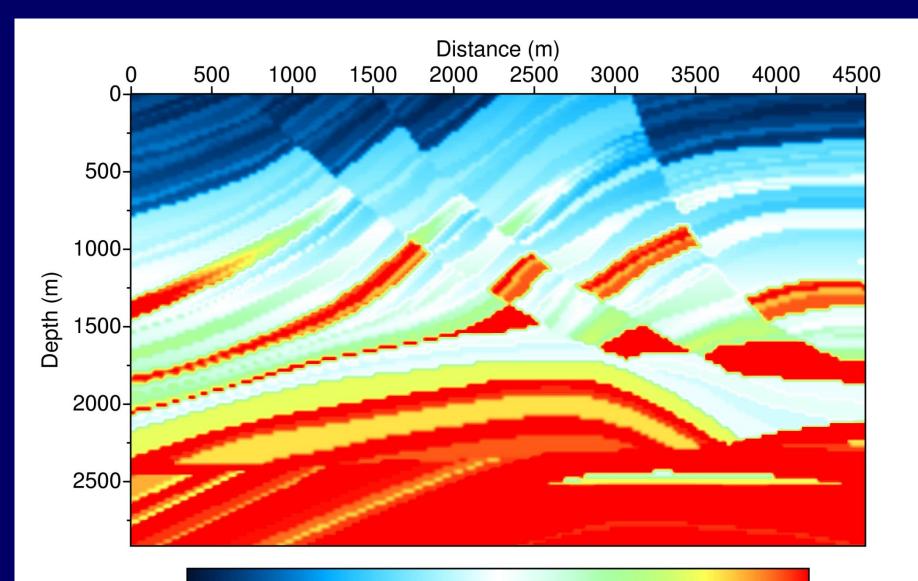
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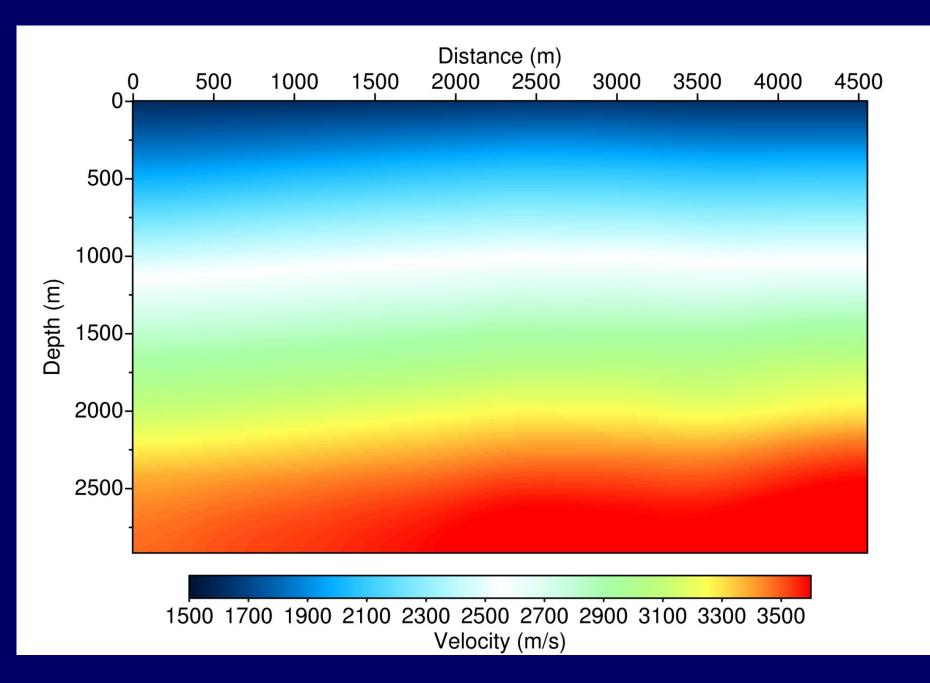
Professor Yanghua Wang

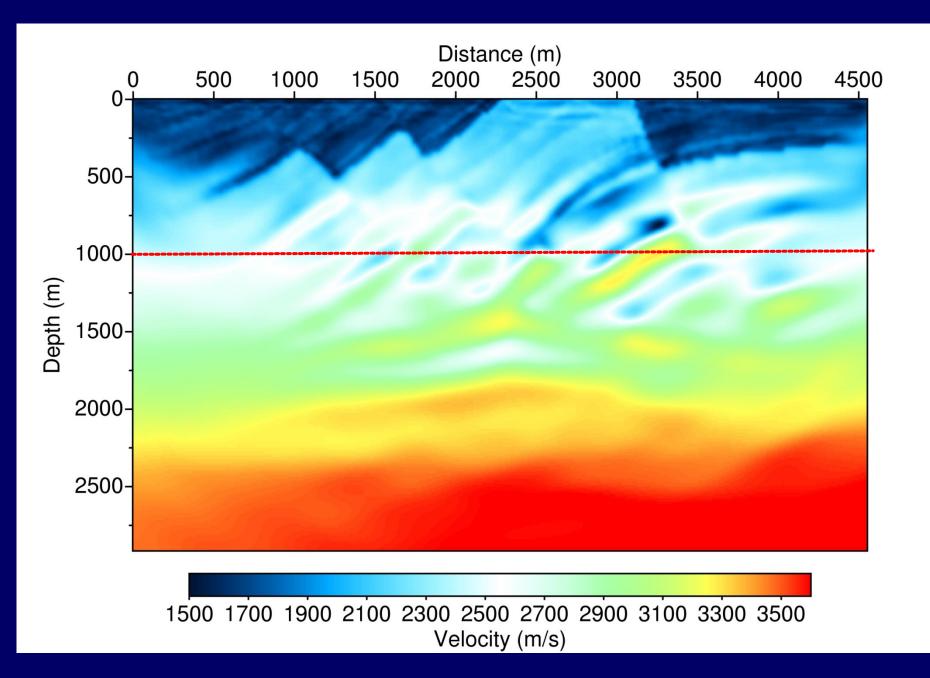
Centre for Reservoir Geophysics

- Chair Professor in Reservoir Geophysics
- Director, Centre for Reservoir Geophysics



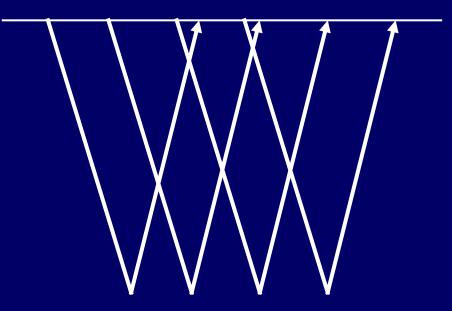
1500 1700 1900 2100 2300 2500 2700 2900 3100 3300 3500 Velocity (m/s)





Reflection waveform tomography

In seismic waveform tomography, if using reflection data with limited source-receiver offsets, it is difficult to reconstruct the deep part of the subsurface velocity model.



Weighted Updating Scheme

By definition, tomography is an integral along the ray path; that is, data residuals are spread evenly along the ray path.

The residual projected to a cell is then linearly related to the model update of the cell, and thus the model update depends upon the total length of ray segments within the cell, i.e. the ray density.

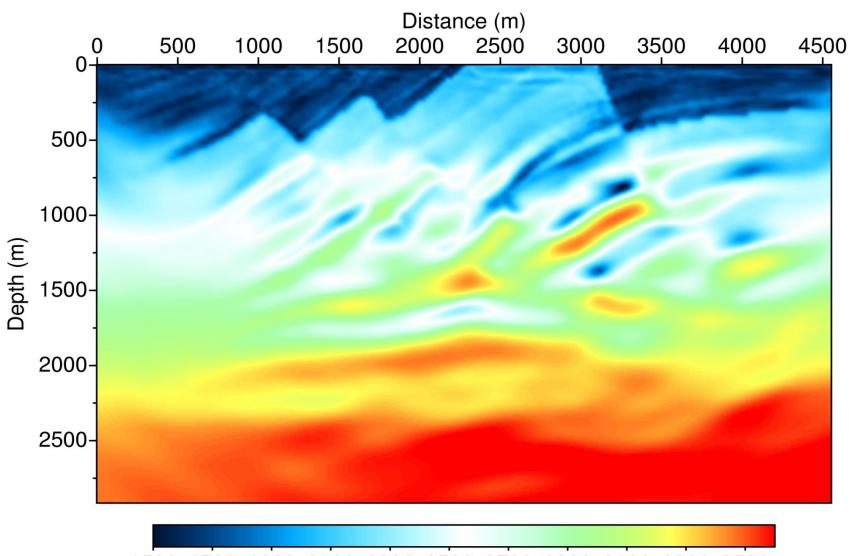
Weighted Updating Scheme

Since the ray density in the shallow depth is much higher than that in deeper parts of the model, the model update in the shallow depth is much higher than that in the deep portion of the model.

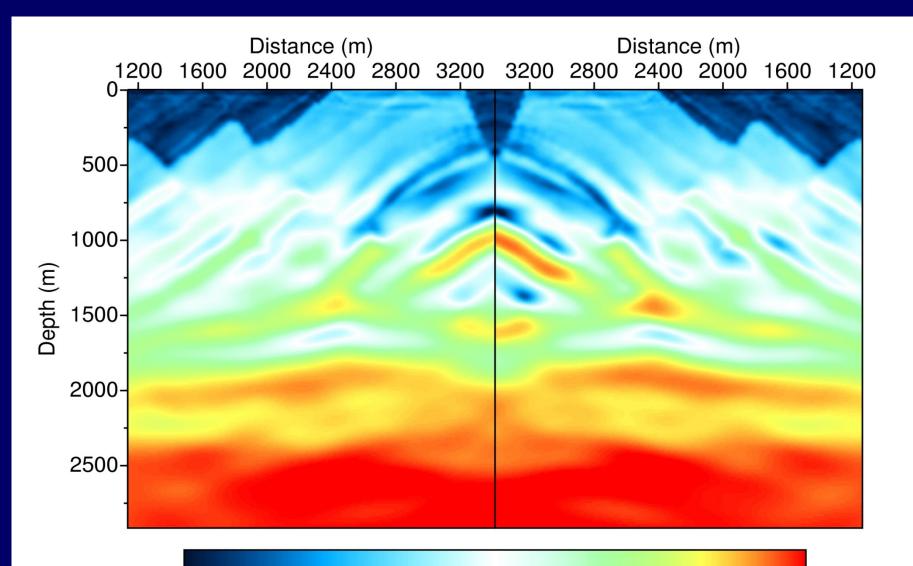
This is a underlying rationale for using depth-dependent weights in model updating.

Weighted Updating Scheme

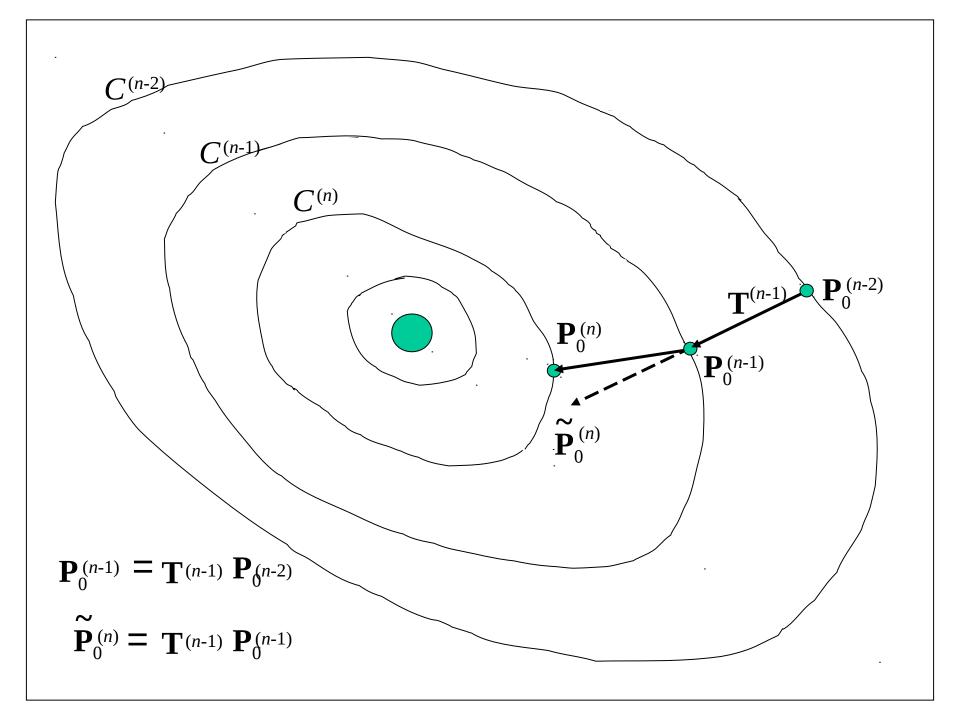
$$\mathbf{m}^{(\ell+1)} = \mathbf{m}^{(\ell)} - \alpha \mathbf{\gamma}$$
$$\mathbf{m}^{(\ell+1)} = \mathbf{m}^{(\ell)} - \alpha \mathbf{H}^{-1} \mathbf{\gamma}$$
$$\mathbf{m}^{(\ell+1)} = \mathbf{m}^{(\ell)} - \alpha \beta(z) \mathbf{\gamma}$$

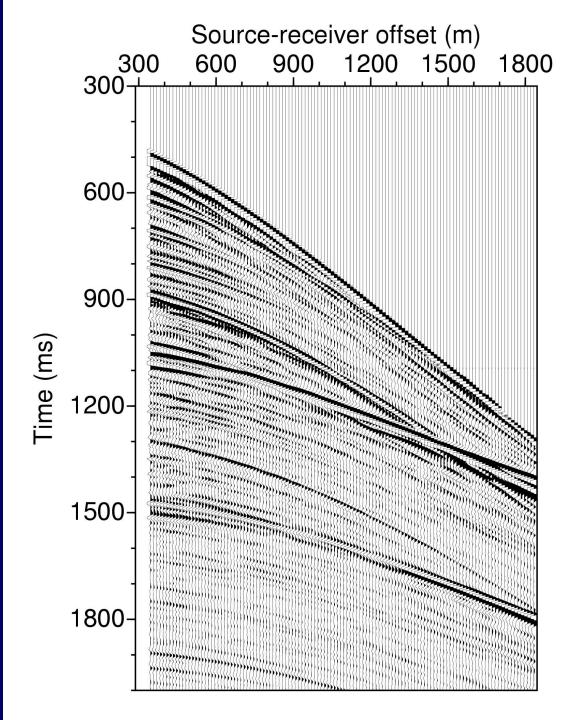


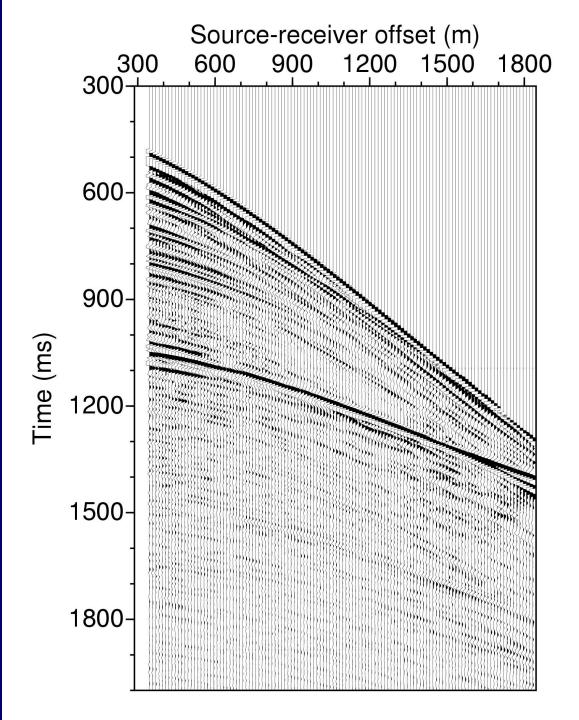
1500 1700 1900 2100 2300 2500 2700 2900 3100 3300 3500 Velocity (m/s)

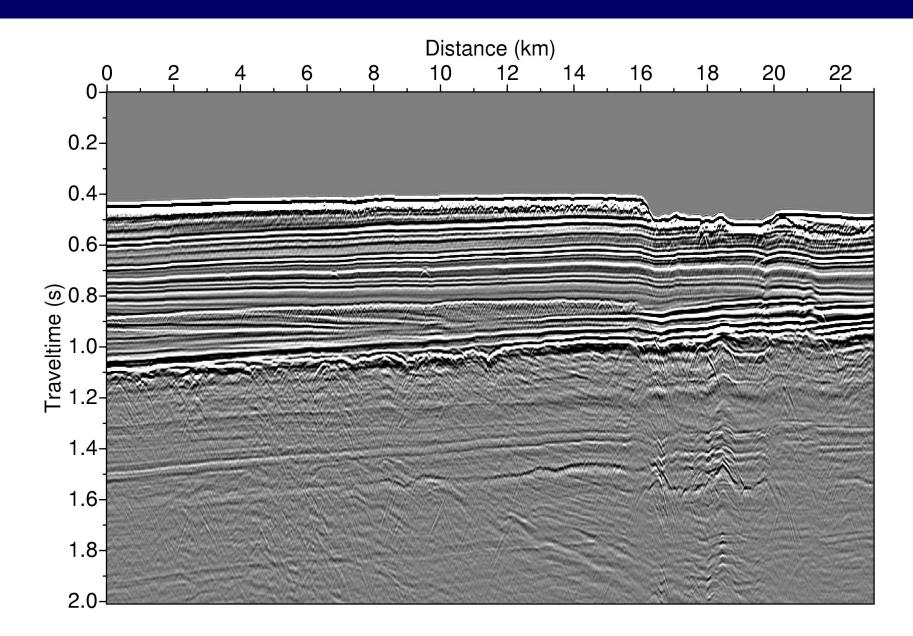


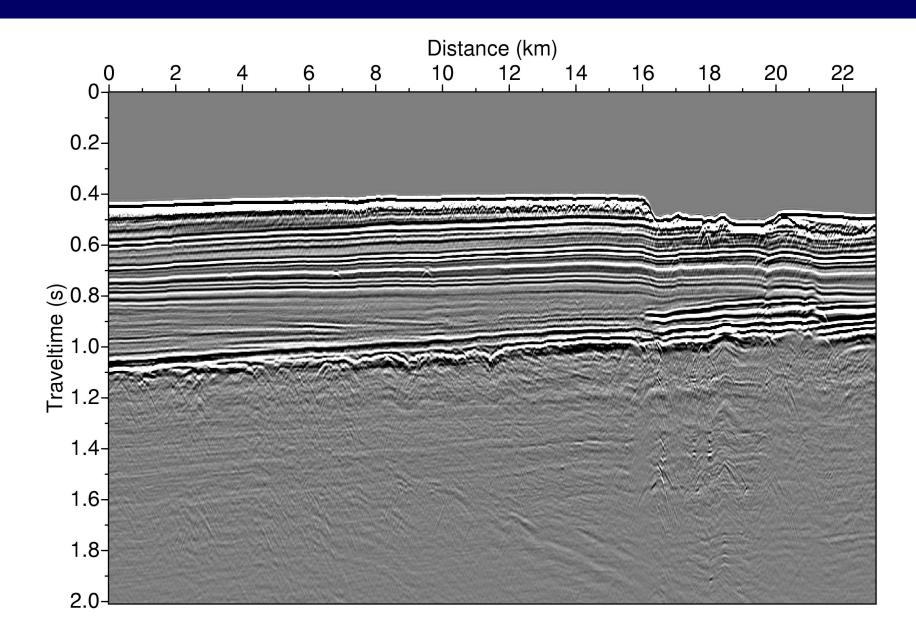
1500 1700 1900 2100 2300 2500 2700 2900 3100 3300 3500 Velocity (m/s)



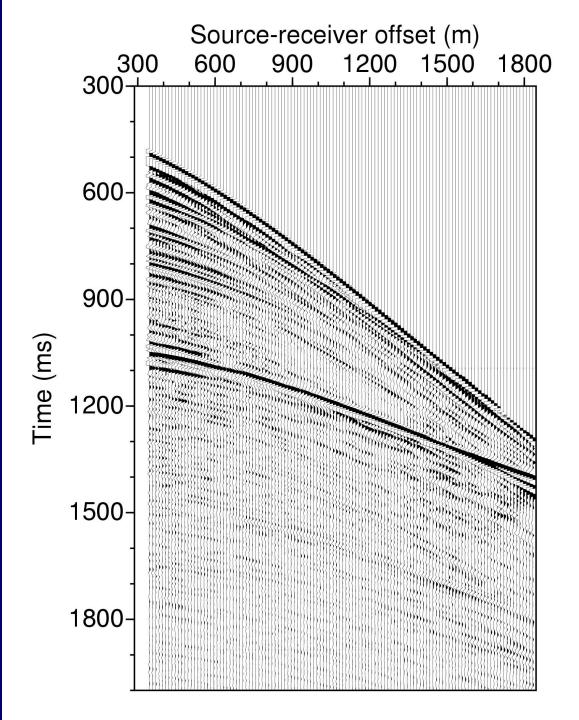


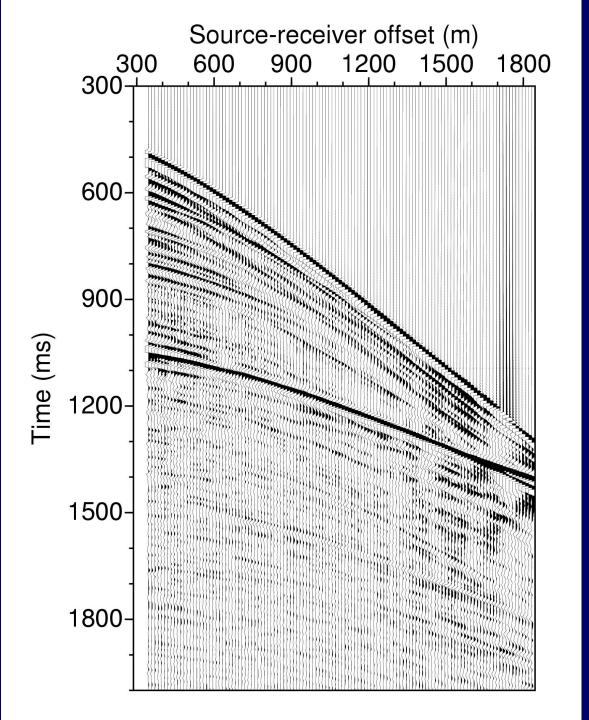






Multiple attenuation





2D

$$G_{2D}(\mathbf{r},\mathbf{r}_{s},\omega) = \left(\frac{-\mathrm{i}c}{8\pi\omega R}\right)^{1/2} \exp\left(-\mathrm{i}\omega\frac{R}{c}\right)$$
3D

$$G_{3D}(\mathbf{r},\mathbf{r}_{s},\omega) = \frac{1}{4\pi R} \exp\left(-\mathrm{i}\omega\frac{R}{c}\right)$$

Partial compensation

$$W(\omega) = \sqrt{\frac{2\pi Rc}{i\omega}}$$

(frequency domain)

$$W(t) = D_{-1/2}(t)\sqrt{2\pi Rc}$$

(time domain)

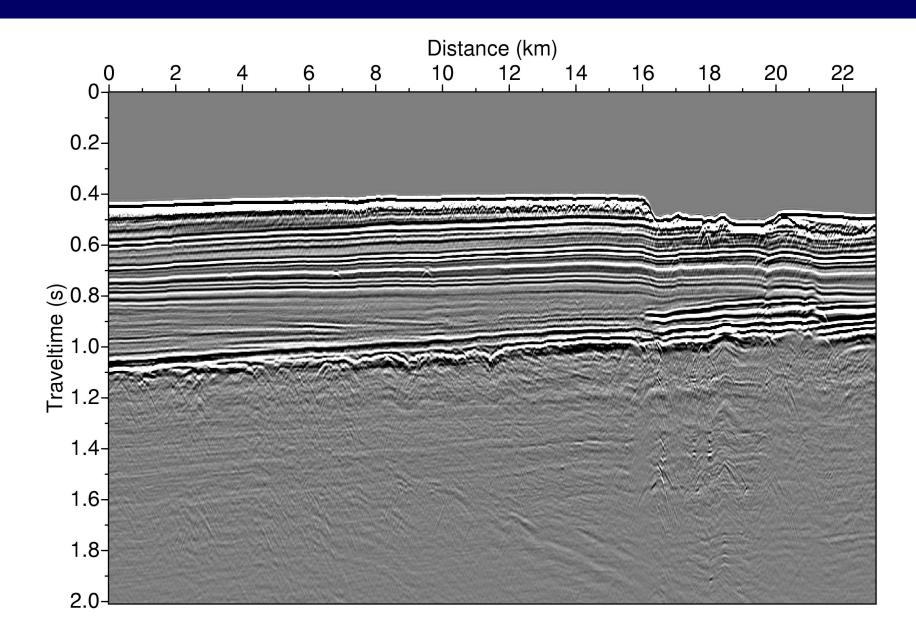
Partial compensation

 $W(t) \propto D_{-1/2}(t) c(t) \sqrt{t}$

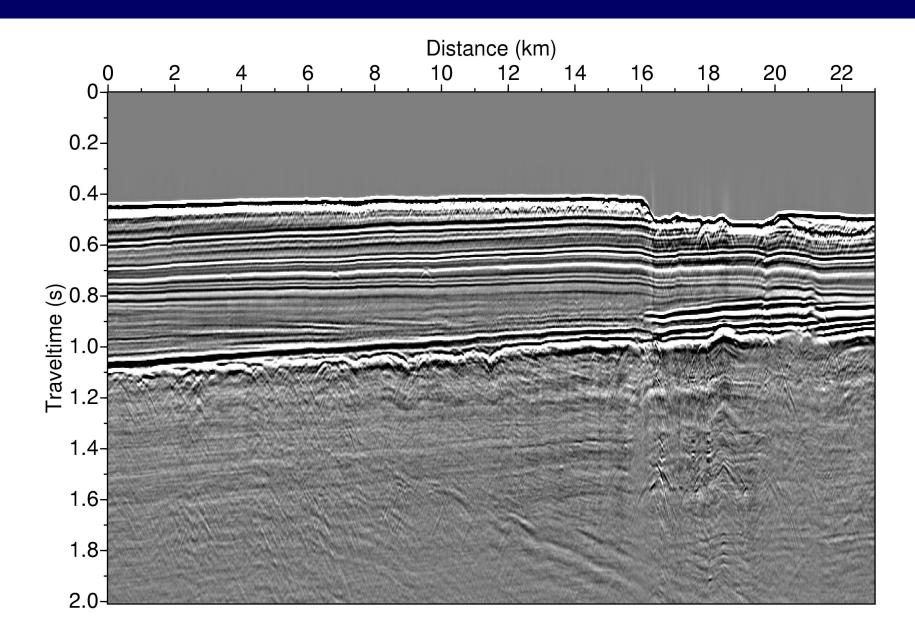
Implementation

Amplitude compensation in the time domain

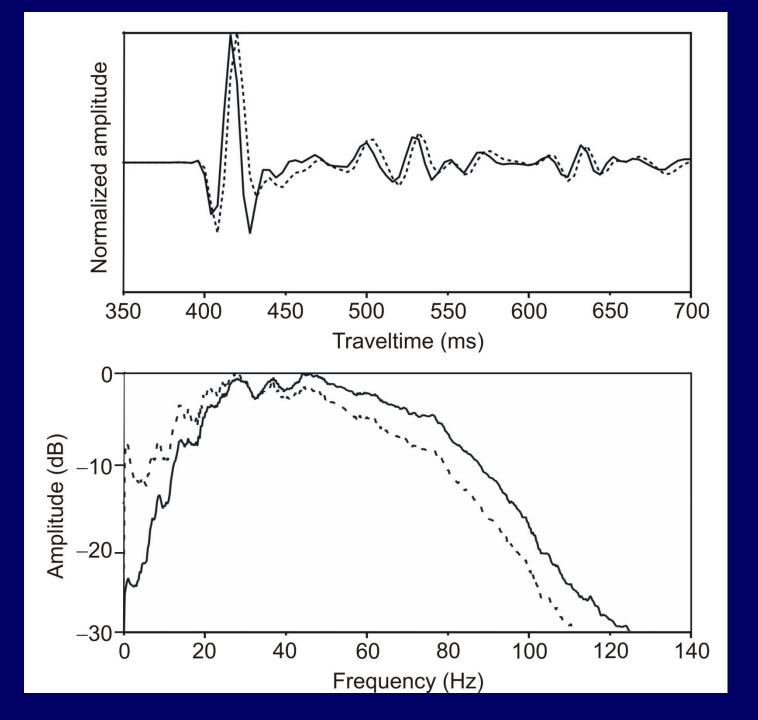
Phase correction in the frequency domain

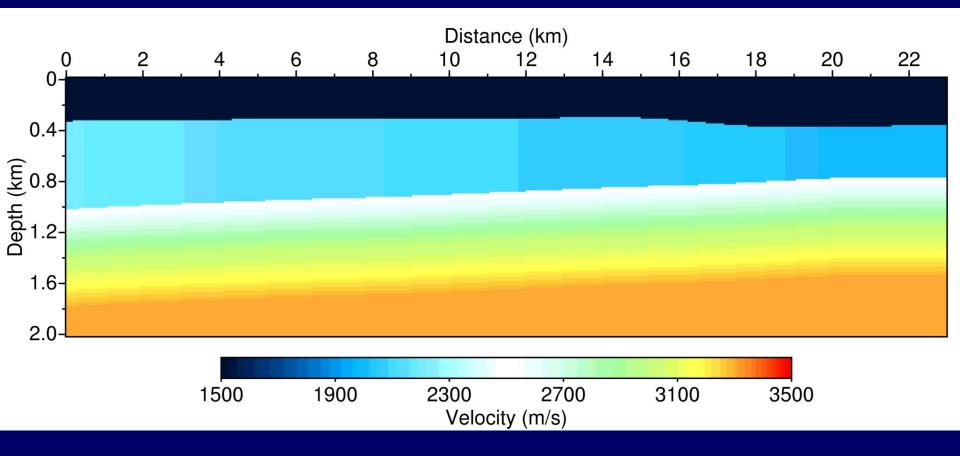


Multiple attenuation

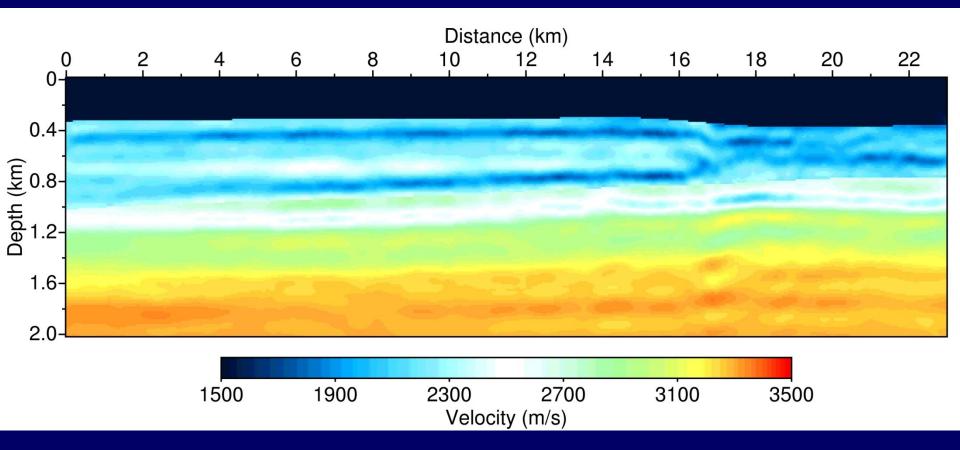


Amplitude and phase compensation

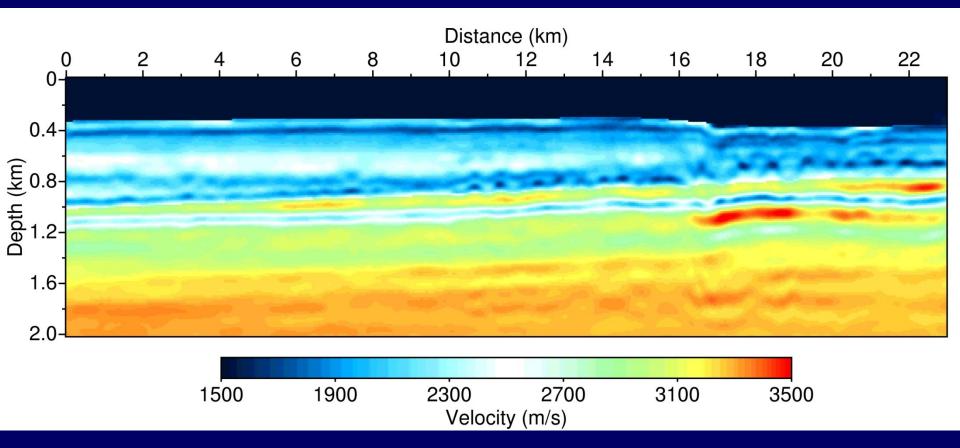




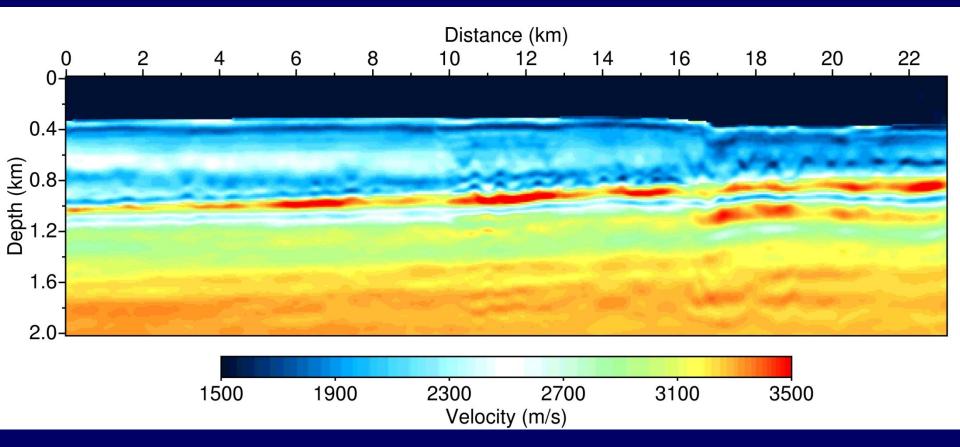
Traveltime tomography



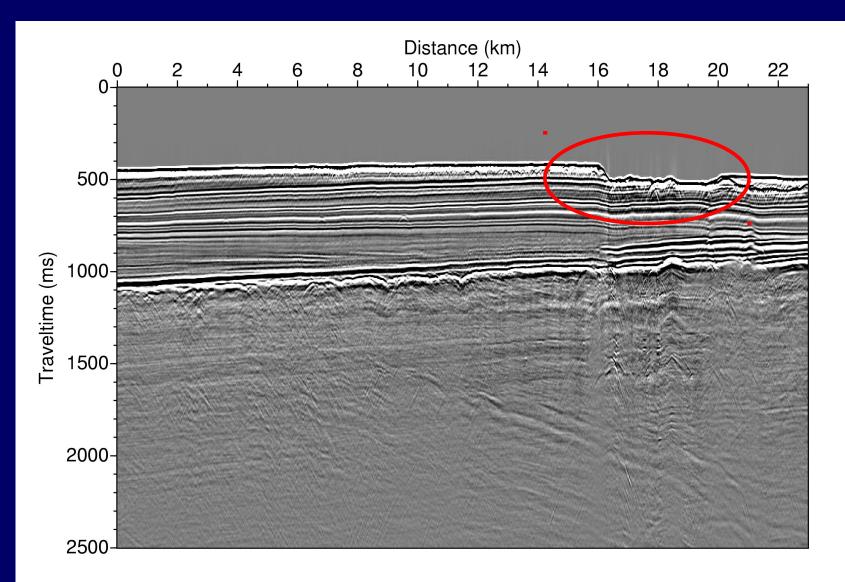
Tomography using frequencies of 6.9-7.5 Hz



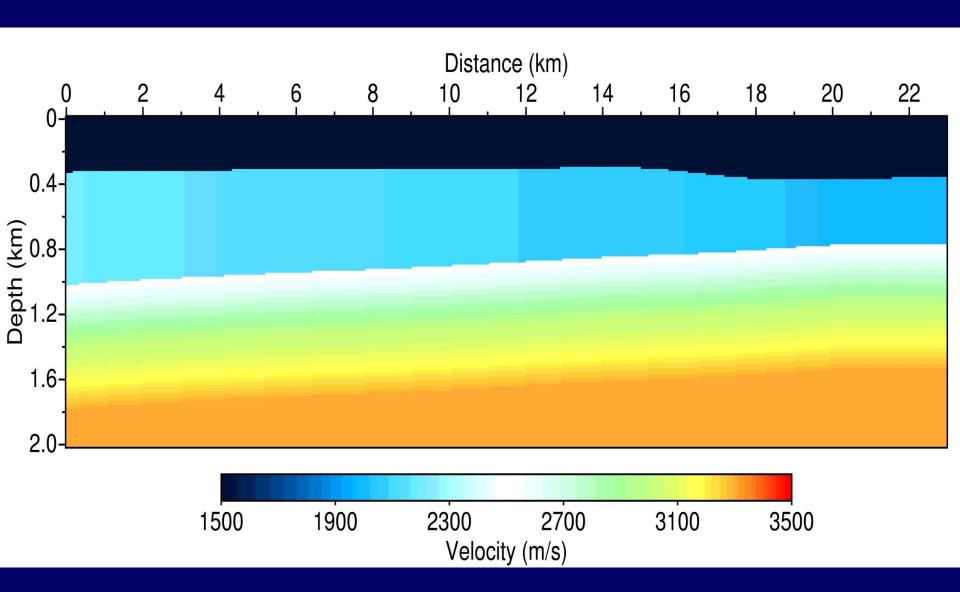
Tomography using frequencies of 6.9-13.8 Hz



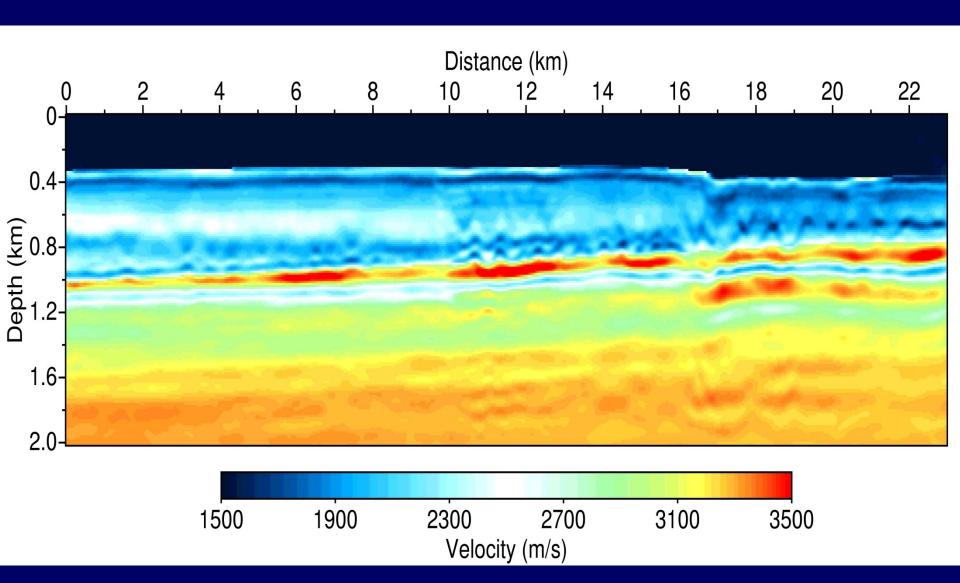
Tomography using frequencies of 6.9-30.0 Hz



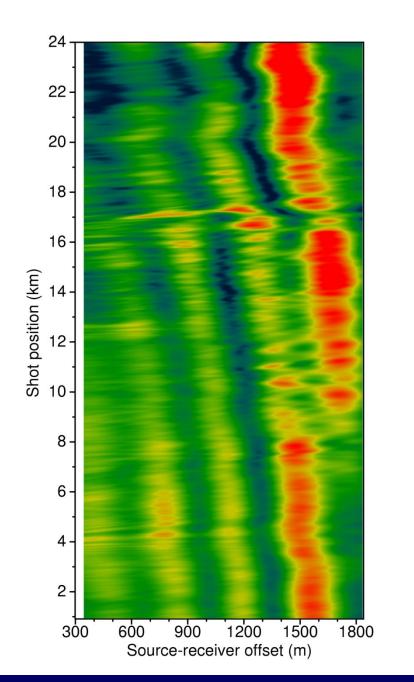
Stack section with amplitude and Phase compansation

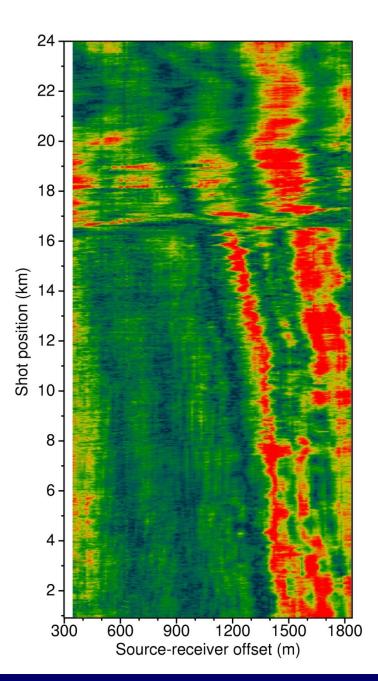


Traveltime tomography



Waveform tomography





Key points

Weighted updating scheme for inversion Pre-stack multiple attenuation Partial compensation of 3D to 2D Applicable to the exploration seismic data